# The Physical Oceanography of Bridport Inlet, N.W.T. Volume I of II. <br> Analysis and Interpretation 

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Canadian Contractor Report of Hydrography and Ocean Sciences No. 9

The Physical Oceanography of Bridport Inlet, N.W.T. Volume I of II. Analysis and Interpretation

by<br>Paul Greisman<br>Dobrocky Seatech Ltd. Sidney, B.C. V8L 3S1<br>for<br>Institute of Ocean Sciences<br>Department of Fisheries and Oceans<br>Sidney, B.C. V8L 4B2 Inlet in 1979 and 1980 by the Frozen Sea Research Group of the Institute of Ocean Sciences. Previous analyses were performed by R.A. Lake and by the present author and are referred to in the text. The 117 individual CTD profiles measured in 1980 are presented in Volume IT. The CTD and recorded current meter data can be obtained on computer tape from the Institute of Ocean Sciences and will be deposited with the Marine Envitonmental Data Services (DFO-OSS) in Ottawa.

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#### Abstract

Greisman, P. 1983. The Physical Oceanography of Bridport Inlet, N.W.T. Volume I of II: Analysis and Interpretation. Can. Contract. Rep. Hydrogr. Ocean Sci. No. 9.

Time series measurements of currents, water properties and water levels from 1979 and 1980 were analysed. The currents within the Inlet cannot be resolved adequately by direct measurement or by dynamic height computations, while the currents at the Inlet entrance rarely exceed $12 \mathrm{~cm} \mathrm{~s}^{-1}$. The exchange of salt through the Inlet entrance appears to be driven by densification of the water within the Inlet due to brine rejection from growing sea ice. The convective processes are not clearly indicated by the vertical density profiles, but rather the system is dominated by horizontal advection and probably by intermittent convection. There is a very pronounced interannual variability in the vertical stratification of the Inlet.


Key words: estuarine circulation, brine rejection, convection, Arctic Oceanography, thermohaline processes.

## RÉSUMÉ

Greisman, P. 1983. The Physical Oceanography of Bridport Inlet, N.W.T. Volume I of II: Analysis and Interpretafion. Can. Contract. Rep. Hydrogr. Ocean Sci. No. 9.

Le présent rapport analyse les mesures par séries chronologiques des courants et des propriétés et niveaux de 1'eau, en 1979 et 1980. Les courants à l'intérieur de l'inlet ne peuvent être déterminés par une mesure directs ni par le calcul de la hauteur dynamique; à l'entrée de l'inlet, la vitesse des courants dépasse rarement $12 \mathrm{~cm} \mathrm{~s}^{-1}$. L'échange de sel à l'entrée de l'inlet semble provenir de la densification de l'eau à l'intérieur de l'inlet, causée par le rejet de saumure contenue dans la glace de mer en croissance. Les processus de convection ne sont pas clairement montrés par les profits verticaux de la densité, mais le système est plutôt dominé par une advection horizontale et probablement par une convection intermittente. Il existe d'une année à l'autre une variabilité très prononcée de la stratification verticale dans 1'inlet.

Mots-clés: circulation estuarienne, rejet de saumure, convection, océanographie de l'Arctique, processus thermohalins.

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Bob Lake patiently corrected all my misapprehensions of the data set and suffered through several verbal drafts.

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Dennis Richards drew the figures, Wendy Holmes typed the draft manuscript and Lorena Quay typed the final report.

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## DISCLAIMER

This report was written under contract and its contents are the responsibility of the author.

## I INTRODUCTION

During March and April 1980 personnel of the Institute of Ocean Sciences (R.A. Lake, D. Richards, and R. Cooke) performed a physical oceanographic survey in Bridport Inlet, Melville Island, NWT where an LNG plant had been proposed for the Arctic Pilot Project. Previous surveys had been performed in 1978 (Greisman, 1978; Lake, 1979a) and in 1979 (Lake, 1979b, c). The data from the 1978 survey collected by Innovative Ventures of Calgary, were of doubtful quality which, in part, motivated the later program. In 1979, 2 current meters were installed at each of 3 locations, and CTD measurements were made at 8 locations. The results of the 1979 survey revealed that more intensive measurements at the entrance to Bridport Inlet were required as well as a more dense network of CTD surveys.

In 1980,9 current meters were deployed at the locations shown in Figure 1. Two current meters were emplaced at 12 m and 50 m depth at locations 7 and 8 , while at the other sites current meters were installed at 12 m depth. Table 1 lists the specifics of the current meter deployments, and Table 2 lists their geographic positions. It was hoped that the high density sampling at the Inlet entrance would resolve the details of the flow. In addition, CTD measurements were made at each of the current meter sites plus 4 additional locations (see Figure 1). The CTD measurements at the Inlet entrance wefe designed to determine the horizontal density gradients for dynamic computations of geostrophic flows. A total of 146 CTD measurements were made with multiple casts performed at each station. Casts were repeated at fixed time intervals at some stations to provide a time series over a tidal cycle. Table 3 lists the times and locations of the CTD casts performed. Two tide gauges were deployed; one just outside the Inlet and the other near the eastern shore. Finally, current profiles with an ultra-sonic current meter were measured at 8 locations.

## A. BATHYMETRY

The bottom topography of Bridport Inlet has been described elsewhere. Figure 2 is reproduced from Lake (1979,b).


FIGURE 1. Locations of current meter moorings, CTD Profiles and tide gauges in Bridport Inlet, March - April 1980.

TABLE 1
CURRENT METER DEPLOYMENTS

| Site | Meter Serial No. | Depth Metres | Deployed |  | Recovered |  | Record Length days-hrs;min:sec | Remarks |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Time | Date | Time | Date |  |  |
| CM-4 | 3228 | 12 | 00:45:00 | 27-03-80 | 13:15:45 | 28-04-80 | 32-12:30:45 |  |
| CM-5 | 1936 | 12 | 15:30;00 | 27-03-80 | 14:45:44 | 28-04-80 | 31-23:15:44 | Note 3 |
| CM-6 | 3223 | 12 | 19:16:00 | 27-03-80 | 16:45:50 | 28-04-80 | 31-21:29:50 | Note 4 |
| CM-7 | 1930 | 50 | 23:00:00 | 27-03-80 | 18:45:22 | 28-04-80 | 31-20:45:22 | : |
| CM-7 | 1932 | 12 | 02:45:00 | 28-03-80 | 18:30:42 | 28-04-80 | 31-15:45:42 |  |
| CM-8 | 2466 | 50 | 20:15:00 | 28-03-80 | 20:00:32 | 28-04-80 | 30-23:45:32 |  |
| CM-8 | 3387 | 12 | 18:45:00 | 28-03-80 | 20:15:46 | 28-04-80 | 31-01:30:46 |  |
| CM-9 | 1931 | 12 | 21:15:00 | 28-03-80 | 21:31:04 | 28-04-80 | 31-00:16:04 |  |
| CM-10 | 1939 | 12 | 00:45:00 | 29-03-80 | -15:01:09 | 29-04-80 | 31-14:16:09 |  |

Note 1: Dates and times are Universal Coordinated Time (G.M.T.).
Note 2: Times are given for first and last valid record.
Note 3: Direction vane gave unreliable results.
Note 4: Clock lost 4 hours during recording period of 4.72 seconds/scan.

TABLEE 2

LOCATIONS OF TIDE GAUGES, CURRENT METERS AND CTD SITES

| SITE | UTM COORDINATES |  | GEOGRAPHIC COORDINATES |  |
| :---: | :---: | :---: | :---: | :---: |
|  | ZONE |  |  |  |
|  | N | E | N | W |
| TG-1 | 8,324,300 | 561,050 | $74^{\circ} 59^{\prime \prime} 54^{\prime \prime}$ | $108^{\circ} 53^{\prime} 08^{\prime \prime}$ |
| TG-3 | 8,328,375 | 569,250 | $75^{\circ} 01^{\prime \prime} 54^{\prime \prime}$ | $108^{\circ} 35^{\prime} 50^{\prime \prime}$ |
| CM - 4 | 8,328,650 | 566,550 | $75^{\circ} 02^{\prime} 06^{\prime \prime}$ | $108^{\circ} 41^{\prime} 20^{\prime \prime}$ |
| - 5 | 8,325,300 | 566,250 | $75^{\circ} 00^{\prime} 18^{\prime \prime}$ | $108^{\circ} 42^{\prime} 20^{\prime \prime}$ |
| - 6 | 8,323,775 | 562,950 | $74^{\circ} 59^{\prime} 34^{\prime \prime}$ | $108^{\circ} 49^{\prime} 18^{\prime \prime}$ |
| - 7 | 8,323,950 | 562,525 | $74^{\circ} 59^{\prime} 40^{\prime \prime}$ | $108^{\circ} 50^{\prime} 03^{\prime \prime}$ |
| - 8 | 8,324,100 | 562,140 | 74*59'44' | $108^{\circ} 50^{\prime} 55^{\prime \prime}$ |
| -9 | 8,324,250 | 561,750 | 74059'50' | $108^{\circ} 51^{\prime} 40^{\prime \prime}$ |
| -10 | 8,326,650 | 561,100 | $75^{\circ} 01^{\prime} 08^{\prime \prime}$ | 108*52'52' |
| 80-1 | 8,330,075 | 563,025 | $75^{\circ} 02^{\prime} 56^{\prime \prime}$ | $108^{\circ} 48^{\prime} 40^{\prime \prime}$ |
| - 2 | 8,326,200 | 563,225 | $75^{\circ} 00^{\prime \prime} 51^{\prime \prime}$ | 108048'30' |
| -3 | 8,323,950 | 562,525 | $74^{\circ} 59^{\prime} 40^{\prime \prime}$ | $108^{\circ} 50^{\prime} 03^{\prime \prime}$ |
| -4 | 8,321,050 | 561,200 | $74^{\circ} 58^{\prime} 08^{\prime \prime}$ | $108^{\circ} 53^{\prime} 02^{\prime \prime}$ |
| - 5 | 8,323,525 | 568,600 | $74^{\circ} 59^{\prime} 18^{\prime \prime}$ | $108^{\circ} 37^{\prime} 36^{\prime \prime}$ |
| UCM - 1 | 8,325,500 | 561,375 | $75^{\circ} 00^{\prime} 30^{\prime \prime}$ | $108^{\circ} 52^{\prime} 20^{\prime \prime}$ |
| - 2 | 8,324,650 | 564,650 | $74^{\circ} 59^{\prime} 58^{\prime \prime}$ | 108\% $45^{\prime} 35^{\prime \prime}$ |
| - 3 | 8,326,650 | 569,225 | $75^{\circ} 00^{\prime} 58^{\prime \prime}$ | $108^{\circ} 36^{\prime} 00^{\prime \prime}$ |

## B. INSTRUMENTATION

## 1. Current Meters

The current meters used in the survey were Aanderaa model RCM-4 equipped with speed and direction sensors as well as temperature and conductivity sensors. Data were recorded internally every 15 minutes on magnetic tape. The instrument counts the number of rotor revolutions during the sampling interval so that an average current speed can be computed. On the other hand, direction, temperature and conductivity are instantaneous values. In the data processing the two direction measurements bracketing the speed averaging interval are averaged to produce a direction coincident with the mid-point of sampling interval.

The Savonius rotors have the unfortunate characteristic of stalling at a current speed of $2.2 \mathrm{~cm} \mathrm{~s}^{-1}$. While normally insignificant in most of the world's oceans, the current speeds encountered in the Canadian Archipelago often are too weak to be sensed with the Aanderaa instrument. The percentage of records for which the current speeds were below the stall speed of instrument are shown in Table 5 in section II.B. Particularly for current meters 4 and 10 , located within the Inlet, the speed records can convey no sense of the flow.

The standard Aanderaa direction vane was replaced with a Hydro Products vane mounted below the current meter pressure case in order to facilitate deployment through the ice. The vane responds to current direction at speeds considerably below the stall speed of the rotor.

CTD casts were taken at each mooring site to calibrate the current meter temperature and conductivity sensors.

The ranges, resolutions and accuracies for the variables measured by the Aanderaa instruments are shown in Table 4.

Moorings for the current meters were required to be torsionally rigid to provide a constant orientation. This is a standard procedure in proximity to the magnetic pole where compasses are useless. The 12 m depth meters were suspended from the sea ice by hydraulic hose


## Cross Section



FIGURE 2. Bottom topography of Bridport Inlet (from Lake, 1979 b ),

TABLE 3
CTD RECORDS - BRIDPORT INLET - APRIL 1980

| NO. | SITE | RANGE | DEPTH | G.M.T. | DATE | REMARKS |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 3300 | CM-10 | 2-61 m | 65 m | 18:06 | 30/03 | Bottle sample |
| 3301 | CM-10 | 2-60 | 65 | 20:35 | 30/03 | CM-10 Calibr. |
| 3302 | CM-10 | 2-61 | 65 | 20:41 | 30/03 | CM-10 Calibr. |
| 3303 | CM-9 | 2-45 | 50 | 23:00 | 30/03 | CM-9 Calibr. |
| 3304 | CM-9 | 2-45 | 50 | 23:15 | 30/03 | CM-9 Calibr. |
| 3305 | CM-9 | 2-45 | 50 | 23:30 | 30/03 | CM-9 Calibr. |
| 3306 | CM-8 | 2-80 | 80 | 13:20 | 31/03 | Ice in C cell |
| 3307 | CM-8 | 2-75 | 80 | 13:35 | 31/03 | CM-8 Calibr. |
| 3308 | CM-8 | 2-76 | 80 | 13:45 | 31/03 | CM-8 Calibr. |
| 3309 | CM-7 | 2-100 | 109 | 18:30 | 31/03 | Ice in C cell |
| 3310 | CM-7 | 2-100 | 109 | 18:45 | 31/03 | Bottle sample |
| 3311 | CM-7 | 2-100 | 109 | 19:00 | 31/03 | CM-7 Calibr. |
| 3312 | CM-6 | 2-20 | 26 | 20:00 | 31/03 | Ice in C cell |
| 3313 | CM-6 | 2-20 | 26 | 20:15 | 31/03 | Cm-6 Calibr. |
| 3314 |  |  |  |  |  | Delete |
| 3315 | CM-6 | 2-20 | 26 | 20:32 | 31/03 | CM-6 Calibr. |
| 3316 |  |  |  |  |  | Delete |
| 3317 | CM-5 | 2-25 | 28 | 17:32 | 01/04 | Ice in C cell |
| 3318 | CM-5 | 2-26 | 28 | 17:45 | 01/04 | Ice in C cell |
| 3319 | CM-5 | 2-25 | 28 | 17:50 | 01/04 | CM-5 Calibr. |
| 3320 | CM-5 | 2-25 | 28 | 17:52 | 01/04 | CM-5 Calibr. |
| 3321 | CM-5 | 2-25 | 28 | 18:00 | 01/04 | CM-5 Calibr. |
| 3322 | CM-4 | 2-59 | 62 | 15:30 | 02/04 | CM-4 Calibr. |
| 3323 | CM-4 | 2-58 | 62 | 15:45 | 02/04 | CM-4 Calibr. |
| 3324 | CM-4 | 2-58 | 62 | 16:00 | 02/04 | CM-4 Calibr. |
| 3325 | 80-1 | 2-110 | (120) | 01:45 | 06/04 | Ice in C cell |
| 3326 | 80-1 | 2-110 | (120) | 01:55 | 06/04 |  |
| 3327 | 80-1 | 2-110 | (120) | 02:05 | 06/04 |  |
| 3328 | 80-1 | 2-110 | (120) | 02:15 | 06/04 |  |
| 3329 | 80-2 | 2-100 | (137) | 17:05 | 07/04 | Ice in C cell |
| 3330 | 80-2 | 2-101 | (137) | 17:10 | 07/04 |  |
| 3331 | 80-2 | 2-125 | (137) | 17:25 | 07/04 |  |
| 3332 | 80-2 | 2-130 | (137) | 17:50 | 07/04 |  |
| 3333 | 80-2 | 2-130 | (137) | 19:20 | 07/04 |  |
| 3334 |  |  |  |  |  | Delete |
| 3335 | 80-2 | 2-131 | (137) | 20:23 | 07/04 |  |
| 3336 | 80-2 | 2-130 | (137) | 21:23 | 07/04 |  |
| 3337 | 80-2 | 2-130 | (137) | 21:40 | 07/04 |  |
| 3338 | 80-2 | 2-130 | (137) | 22:00 | 07/04 |  |
| 3339 | 80-2 | 2-130 | (137) | 22:20 | 07/04 |  |
| 3340 | 80-2 | 2-130 | (137) | 22:40 | 07/04 | Bottle sample |
| 3341 | 80-2 | 2-130 | (137) | 23:00 | 07/04 |  |
| 3342 | 80-2 | 2-130 | (137) | 23:20 | 07/04 |  |
| 3343 | 80-3 | 2-100 | 109 | 16:25 | 08/04 |  |
| 3344 | 80-3 | 2-100 | 109 | 16:40 | 08/04 |  |
| 3345 | 80-3 | 2-100 | 109 | 17:00 | 08/04 |  |
| 3346 | 80-3 | 2-100 | 109 | 17:20 | 08/04 |  |


| NO. | SITE | RANGE | DEPTH | G.M.T. | DATE | REMARKS |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 3347 | 80-3 | 2-101 m | 109 m | 17:40 | 08/04 |  | $\sim$ |
| 3348 | 30-3 | 2-100 | 109 | 18:00 | 08/04 |  |  |
| 3349 | 80-3 | 2-100 | 109 | 18:20 | 08/04 |  |  |
| 3350 | 80-3 | 2-100 | 109 | 18:40 | 08/04 |  | $\sim$ |
| 3351 | 80-3 | 2-100 | 109 | 19:00 | 08/04 |  |  |
| 3352 | 80-3 | 2-100 | 109 | 19:20 | 08/04 |  |  |
| 3353 | 80-3 | 2-100 | 109 | 19:40 | 08/04 |  |  |
| 3354 | 80-3 | 2-100 | 109 | 20:00 | 08/04 |  |  |
| 3355 | 80-3 | 2-100 | 109 | 20:20 | 08/04 |  |  |
| 3356 | 80-3 | 2-100 | 109 | 20:40 | 08/04 |  |  |
| 3357 | 80-3 | 2-100 | 109 | 21:00 | 08/04 |  |  |
| 3358 | 80-3 | 2-100 | 109 | 21:20 | 08/04 |  |  |
| 3359 |  |  |  |  |  | Delete |  |
| 3360 | 80-3 | 2-100 | 109 | 21:40 | 08/04 |  |  |
| 3361 | 80-3 | 2-100 | 109 | 22:00 | 08/04 |  |  |
| 3362 | 80-3 | 2-100 | 109 | 22:20 | 08/04 |  |  |
| 3363 | 80-3 | 2-100 | 109 | 22:40 | 08/04 |  |  |
| 3364 | 80-3 | 2-100 | 109 | 23:00 | 08/04 |  |  |
| 3365 | 80-3 | 2-100 | 109 | 23:20 | 08/04 |  |  |
| 3366 | 80-3 | 2-100 | 109 | 23:40 | 08/04 |  |  |
| 3367 | 80-3 | 2-100 | 109 | 00:00 | 09/04 |  |  |
| 3368 | 80-3 | 2-100 | 109 | 00:20 | 09/04 |  |  |
| 3369 | 80-3 | 2-100 | 109 | 00:40 | 09/04 |  |  |
| 3370 | 80-3 | 2-100 | 109 | 01:00 | 09/04 |  |  |
| 3371 | 80-3 | 2-101 | 109 | 01:20 | 09/04 |  |  |
| 3372 | 80-3 | 2-101 | 109 | 01:40 | 09/04 |  |  |
| 3373 | 80-3 | 2-100 | 109 | 02:00 | 09/04 |  |  |
| 3374 | 80-4 | 2-60 | 64 | 16:40 | 09/04 |  |  |
| 3375 | 80-4 | 2-60 | 64 | 17:40 | 09/04 |  |  |
| 3376 | 80-4 | 2-60 | 64 | 18:40 | $09 / 04$ |  |  |
| 3377 | 80-4 | 2-61 | 64 | 19:40 | 09/04 |  |  |
| 3378 | 80-4 | 2-60 | 64 | 20:40 | 09/04 |  |  |
| 3379 | 80-4 | 2-60 | 64 | 21:40 | $09 / 04$ |  |  |
| 3380 | 80-4 | 2-60 | 64 | 22:40 | 09/04 |  |  |
| 3381 | 80-4 | 2-60 | 64 | 23:40 | 09/04 |  |  |
| 3382 | 80-4 | 2-60 | 64 | 00:40 | 10/04 |  |  |
| 3383 | 80-4 | 2-60 | 64 | 01:40 | 10/04 |  |  |
| 3384 | 80-4 | 2-60 | 64 | 02:40 | 10/04 |  |  |
| 3385 | 80-3 | 2-100 | 109 | 15:49 | 10/04 |  |  |
| 3386 | 80-2 | 2-130 | (137) | 17:52 | 10/04 |  |  |
| 3387 | 80-1 | 2-100 | (120) | 20:30 | 10/04 |  |  |
| 3388 | 80-5 | 2-65 | 74 | 17:15 | 12/04 |  | $\cdots$ |
| 3389 | 80-5 | 2-65 | 74 | 17:55 | 12/04 |  | $\cdots$ |
| 3390 | 80-5 | 2-65 | 74 | 18:01 | 12/04 |  |  |
| 3391 | CM-4. | 2-58 | 62 | 17:00 | 22/04 |  |  |
| 3392 | CM-4 | 2-58 | 62 | 17:15 | . 22/04 |  | $\cdots$ |
| 3393 | CM-4 | 2-58 | 62 | 17:30. | 22/04 |  |  |
| 3394 | CM-5 | 2-25 | 28 | 16:00 | 23/04 | Ice in Cell |  |
| 3395 | CM-5 | 2-25 | 28 | 16:15 | 23/04 |  |  |
| 3396 | CM-5 | 2-25 | 28 | 16:30 | 23/04 |  |  |


| NO. | SITE | RANGE | DEPTH | G:M.T. | DATE | REMARKS |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 3397. | CM-5 | 2-25 m | 28 m | 16:45 | 23/04 |  |
| 3398 | CM-6 | 2-22 | 26 | 19:00 | 23/04 |  |
| 3399 | CM-6 | 2-22 | 26 | 19:15 | 23/04 |  |
| 3400 | CM-6 | 2-22 | 26 | 19:30 | 23/04 |  |
| 3401 | CM-7 | 2-100 | 109 | 21:00 | 23/04 |  |
| 3402 | CM-7 | 2-100 | 109 | 21:15 | 23/04 |  |
| 3403 | CM-7 | 2-100 | 109 | 21:30 | 23/04 |  |
| 3404 | CM-7 | 5-63 | 109 | 21:45 | 23/04 | Bottle Sample |
| 3405 | CM-7 | 2-100 | 109 | 22:30 | 23/04 | CTD No. 2 |
| 3406 | CM-7 | 2-100 | 109 | 22:45 | 23/04 | CTD No. 2 |
| 3407 | CM-7 | 2-100 | 109 | 23:15 | 23/04 |  |
| 3408 | CM-8 | 2-80 | 80 | 14:30 | 24/04 |  |
| 3409 | CM-8 | 2-80 | 80 | 14:45 | 24/04 |  |
| 3410 | CM-8 | 2-80 | 80 | 15:00 | 24/04 |  |
| 3411 | CM-9 | 2-45. | 50 | 16:00 | 24/04 |  |
| 3412 | CM-9 | 2-45 | 50 | 16:15 | 24/04 |  |
| 3413 | CM-9 | 2-45 | 50 | 16:30 | 24/04 |  |
| 3414 | CM-10 | 2-61 | 65 | 17:45 | 24/04 |  |
| 3415 : | CM-10 | 2-61 | 65 | 18:00 | 24/04 |  |
| 3416 | CM-10 | 2-61 | 65 | 18:15 | 24/04 |  |

TABLE 4

## INSTRUMENT SPECIFICATIONS

## AANDERAA RCM-4

VARIABLE
RANGE
RESOLUT ION
ACCURACY

Speed
2.2 to $45 \mathrm{~cm} \mathrm{~s}^{-1}$
$0.04 \mathrm{~cm} \mathrm{~s}^{-1}$
Direction
0 to $360^{\circ}$
$0.35^{\circ}$
$\pm 7^{\circ}$
Temperature
$-2^{0}$ to $+3^{\mathrm{O}} \mathrm{C}$
$0.005^{\circ} \mathrm{C}$
$0.015^{\circ} \mathrm{C}$
Conductivity
23 to $30.7 \mathrm{mmho} /$
$0.007 \mathrm{mmho} / \mathrm{cm}$
~. $075 \mathrm{mmho} / \mathrm{cm}$ ( $\sim 0.1 \% / 00$ )

Sample interval $=15$ minutes

GUILDLINE CTD 8101A

| Salinity | 28 to 40 PPT | $0.001^{\circ} / 00$ | $\pm 0.005^{\circ} / \mathrm{oo}$ |
| :--- | ---: | :--- | :--- |
| Temperature | $-2^{\circ}$ to $+30^{\circ} \mathrm{C}$ | $0.001^{\circ} \mathrm{C}$ | $\pm 0.002^{\circ} \mathrm{C}$ |
| Pressure | 0 to 1000 db | 0.1 db | $\pm 1.0 \%$ of reading |

AANDERAA WLR-5 WATER LEVEL RECORDER

| Height of Water Column | 2.7 mm | $\pm 27 \mathrm{~mm}$ |
| :--- | :--- | :--- |
| Time | 0.5 s | $\pm 1 \mathrm{~s} \mathrm{~d}^{-1}$ |

Sample interval $=5$ minutes
while the instruments at 50 m depth were mounted on a 3 m long horizontal spreader bar which was suspended by two lines anchored 15 m apart in the ice.
2. CTDs

CTD profiles were measured with a Guildine Model 8101A instrument. Water samples were taken frequently and salinity determined with a Hytech 6220 bench salinometer in order to monitor the performance of the conductivity cell. In addition, in-situ calibrations of the CTD temperature sensor were performed using previously calibrated thermistors. Details of the CTD calibration procedure may be found in Lewis and Sudar (1972), while salinity and density computations are discussed in Volume II.
3. Tide Gauges

Two Aanderaa WLR 5 tide gauges were deployed for a one month period at the locations shown in Figure 1 and listed in Table 2. The sampling interval was 5 minutes. No attempt was made to correct the water level data for variations in the atmospheric pressure.

## 4. Ultra-sonic current meter

Measurements were made with a prototype instrument developed by the Christian Mikkelson Institute in Norway. This instrument measures the differences in travel time of an acoustic pulse between two orthogonal sets of transmitter receivers. The pulses are reflected off an acoustic mirror displaced vertically from the transmitter-receiver in order to remove the bias of a vertical flow component. The threshold of this instrument is about $1 \mathrm{~mm} \mathrm{~s}^{-1}$ while its accuracy varies with duration of deployment due to electronic drift. For a one hour deployment the accuracy is better than $\pm 3 \mathrm{~mm} \mathrm{~s}^{-1}$.

In the field, the instrument was "zeroed" in the hole drilled through the sea ice.with a cap covering the sensors for each profile. It was then lowered to predetermined depths and the induced turbulence permitted to dissipate. Readings were then made over about a 30 second time interval. Two profiles were measured with the instrument on a
torsionally-rigid aluminum pipe supplying directional data while an average of four profiles were measured at 8 sites with the instrument lowered on a cable. The latter profiles furnish speed versus depth
data only.

Current meters were deployed in Bridport Inlet at the end of March 1980; their locations are shown in Figure 1 and the details of the deployment in Table 1. The meters recorded, on the average, about 30 days of data. The data were plotted as time series and de-spiked. Progressive vector diagrams were constructed in order to aid in interpretation and mean flows over the period of deployment were computed.

Examination of the records revealed that the current speeds were often below the stall speed of the Aanderaa roters ( $2.2 \mathrm{~cm} \mathrm{~s}^{-1}$ ). Rather than assign the resolvable speed to every record below stall speed, it was considered more representative to assign a value of $1.1 \mathrm{~cm} \mathrm{~s}^{-1}$ (halfway between zero and stall speed) to these records. The resulting current vectors are shown in Figure 6 and tabulated in Table 5.

CTD measurements were made at all the current meter sites as well as at five additional locations denoted $80-1$ through 80-5. The locations of the CTD stations are shown in Figure 1 and listed in Table 2. At all the locations multiple CTD casts were taken. Time series measurements (e.g. CTD casts every twenty minutes for $9 \frac{1}{2}$ hours at site $80-3$ ) were performed at several locations over the period from the end of March to the end of April. A list of all the CTD measurements performed is presented as Table 3. The individual CTD profiles are reproduced in Volume 2 and the average CTD profiles at each location are reproduced in Appendix A.

## A. GEOSTROPHIC CURRENTS

In order to describe quantitatively the circulation in Bridport Inlet, the mass field measurements were used to compute geostrophic shears. In addition an attempt was made to reference the computed shears with the measured currents.

Dynamic heights were computed at each station using an average of all the CTD casts at that location. Geostrophic shears were then computed between each station pair. There are however, important limitations on the use of the dynamic method in Bridport Inlet.

By far the most important of these is the presence of relatively weak horizontal density gradients.

The geostrophic shear is computed from the vertical derivative of the horizontal equations of motion.

$$
\begin{equation*}
f \hat{\mathrm{k}} \times \overline{\mathrm{u}}=-\frac{1}{\rho} \nabla \mathrm{p} \tag{1}
\end{equation*}
$$

are the horizontal equations of motion of a geostrophic fluid, where $f$ is the Coriolis parameter, $\hat{k}$ is the vertical vector, $\bar{u}$ is the velocity vector, $\rho$ the fluid density and $\nabla p$ the horizontal pressure gradient. The $z$ derivative of (1) is

$$
\begin{equation*}
\frac{d \rho}{d z} f \hat{k} \times \bar{u}+\rho f \hat{k} \times \frac{d \bar{u}}{d z}=+g \nabla \rho \tag{2}
\end{equation*}
$$

The first term on the left hand side is at most $1 \%$ of the second term in the ocean, and when it is ignored the relationship used to compute geostrophic shears results:

$$
\begin{equation*}
\rho f \hat{k} \times \frac{d \bar{u}}{d z}=g \nabla \rho \tag{3}
\end{equation*}
$$

The precision with which (3) is evaluated is determined by the precision with which the horizontal density gradient and the station spacing can be measured. Within Bridport Inlet the station spacing is roughly 1 km , while the density can be resolved to about $\pm 2 \times 10^{-3}$ $\mathrm{kg} \mathrm{m}^{-3}$. Using typical numerical values,

$$
\begin{aligned}
& \frac{d \bar{u}}{d z}=\frac{10 \mathrm{~m} \mathrm{~s}^{-2}\left(0 \rho \pm 2 \times 10^{-3} \mathrm{~kg} \mathrm{~m}^{-3}\right)}{10^{3} \mathrm{~kg} \mathrm{~m}^{-3}\left(1.4 \times 10^{-4} \mathrm{~s}^{-1}\right)\left(1 \times 10^{3} \mathrm{~m}\right)} \\
& \frac{\mathrm{d} \bar{u}}{\mathrm{dz}}=7.2 \times 10^{-2}\left(\delta \rho \pm 2 \times 10^{-3}\right) \mathrm{s}^{-1}
\end{aligned}
$$



FIGURE 3. Surface geostrophic currents with precisions. The bottom is taken as the reference level where the velocity is zero.


FIGURE 4. Profiles of geostrophic shear at the Inlet entrance.

The precision of the computed vertical shear is therefore $\pm 1.4 \times 10^{-4} \mathrm{~s}^{-1}$. For a water column 100 m deep, the precision of the computed surface velocity is therefore

$$
\pm 1.4 \times 10^{-2} \mathrm{~m} \mathrm{~s}^{-1} \text { or } \pm 1.4 \mathrm{~cm} \mathrm{~s}^{-1}
$$

The computed geostrophic surface flows using the bottom as the level of no motion are shown in Figure 3. The precisions vary due to varying station spacings. With the exception of the flows at the Inlet entrance, the computed geostrophic current magnitudes are all smaller than the associated precisions. No conclusion should be drawn therefore from the currents computed within the Inlet. Not only are the magnitudes of these flows poorly determined, but their directions too may be in error.

Significant geostrophic shears were computed at the entrance between current meter sites 7,8 and 9 and these profiles are shown in Figure 4. The flow at the entrance is characterized by a vertical shear directed into the Inlet at the top of the water column. In order to satisfy the continuity constraint a barotropic (constant with depth) mean flow must be added. The vertical shear however, is concentrated in the interval between 10 and 25 m depth so that it is within this depth range where a reversal of the mean flow direction is expected. The total velocity difference through the water column is about $8 \mathrm{~cm} \mathrm{~s}^{-1}$.

The computations suggest inward flowing surface waters with a weak outward flow at depth. Such a flow is characteristic of a "negative estuary" or a basin in which densification takes place either due to evaporation (as in the Mediterranean Sea) or to brine rejection by growing sea ice.

## B. CURRENT METER MEASUREMENTS

The histograms for the 9 current meter velocity time series are shown in Figures 5a through i. The maximum speeds recorded are about $16 \mathrm{~cm} \mathrm{~s}^{-1}$ at the current meters at 12 m depth in the Inlet entrance. Most of the histograms show predominantly bi-directional flow suggestive of tidal oscillations. (Tidal flows are discussed in Section V.) The histogram for current meter 5 shows no currents flowing in any but the $070^{\circ}-130^{\circ}$ sextant. This narrow range of indicated direction is almost certainly due to a malfunctioning vane. In addition, a profile measured with an ultrasonic current meter at site 5 showed the flow direction to be about $230^{\circ} \mathrm{T}$, while the speeds were in agreement with the Aanderaa current meter. The data from C.M. 5 are, therefore, indicative of speed only.

Table 5 is a summary of the mean current vectors computed at each of the current meters. Included in the table is the percentage of speed records below the threshold speed of the Savonius rotor. This percentage can have a significant influence on the precision to which the average current vectors are known.

If p is the fraction of current records below stall speed, and $S$ is the stall speed, then the maximum value of the true average current speed (V̄max) is

$$
\bar{V}_{\max }=\overline{\mathrm{V}} \mathrm{a}(1-\mathrm{p})+\mathrm{Sp}
$$

where $\bar{V} a$ is the average of the speeds above stall speed. The minimum value is

$$
\overline{\mathrm{V}} \min =\overline{\mathrm{V}} \mathrm{a}(1-\mathrm{p}),
$$

since the speed may lie anywhere in the interval ( $0<\mathrm{V} \leq \mathrm{S}$ ). The mean expected speed is, therefore

$$
\overline{\text { Vimean }}=\overline{\mathrm{V} a}(1-\mathrm{p})+\frac{\mathrm{S}}{2} \mathrm{p} .
$$

If it is assumed that the speeds are normally distributed on the interval $(0, S)$ then $\pm 3$ standard deviations cover $98 \%$ of the interval. The $95 \%$ confidence interval for the speed in the range below rotor

FREDUENCY DISTRIBUTION OF DIRECTION ANO PATE
BRD 124500270380 GMT CM-4. 3123 ERIOPORT INLET 7502110841315 322GTCTVG


Figure 5a. Histogram for CM 4
BRŨ $12 \quad 3015270380$ EMT CM-5 3070 BRIDPORT INLET $75003108423151936 T C T V G$

$\qquad$ $\ldots 0$ $\qquad$ 0
NUMBER OF RECORES AT OR BELOW STALL SPEED 12.2 CM/SECI 1245

Figure 5b, Histogram for CM 5
BRD 121619270380 GMT CH-E 3040 BRIDPORT INLET $74596109493153223 T C T V 6$


NUMBER OF REGORDS AT OR BELOw STALL SPEED $12.2 \mathrm{CM} / 5 E C Y \quad 204$

Figure 5c. Histogram for CM 6.

## FREQUENCY DISTRIBUTION OF DIRECTION AND RATE

BRU $50 \quad 0023270380$ GMT CM-7 3056 BRIDPORT INLET $7459710850115193 U T C T V G$


Figure 5d. Histogram for CM 7 at 12 m depth.


Figure 5e. Histogram for CM 7 at 50 m depth.


NUMBER OF RECORDS AT OR BELOW STALL SPEED 12:2 CM/SEC) $15 B 7$

Figure 5f: Histogram for CM 8 at 12 m depth.

FREQUENCY DISTRIBUTION OF DIRECTION AND RATE BFO 12 451828038日 EMT CM-8 2983 BPIDPORT INLET $74597108504153387 T C T V 6$


Figure 5g. Histogram for CM 8 at 50 m depth.

FREQUENCY DISTRIBUTION OF DIRECTION AND RATE
BRO 1215212 PO3 12 GMT CM-9 2978 BPIDPORT INLET $74598108517151931 T C T V G$


NUMBER OF RECORDS AT OR BELO STALL SPEED (2.2 CM/SEC) 709

Figure 5h. Histogram for GM 9.

FREOUENCY DISTRIBUTION OF DIRECTION AND RATE
BPU 124500290380 GMT CMID 3034 BRIOPORT INLET $75011148529151939 T C T V 6$



NUMEEP OF RECOROS AT OR BFLOW STALL SPEEO 12.2 CM/SEC) 2978

Figure 5i. Histogram for CM 10.
stall speed corresponds with a range of $\pm 2$ standard deviations, and thus the $95 \%$ confidence interval is

$$
\pm \frac{2}{3} \frac{S}{2}= \pm \frac{S}{3}
$$

The final expression for the average speed is therefore

$$
\bar{V}=\bar{V} a(1-p)+\left(\frac{S}{2} \pm \frac{S}{3}\right) p .
$$

For a time series with $10 \%$ of the speeds below stall speed, the $95 \%$ confidence interval for the mean speed is $\pm s / 30$. If $90 \%$ of the speeds are below stall speed then the confidence interval is $\pm 95 / 30$.

The 5 th column in Table 5 lists the confidence intervals for the computed average speeds. The last column is the ratio of the uncertainty to the mean. It must be concluded from these calculations that the mean velocity vector from CM-4 is not significantly different from zero and that the vector at $\mathrm{CM}-10$ is only barely significant.

Figure 6 shows the statistically significant mean velocity vectors measured in Bridport. Inlet. It would be unwise to speculate on the circulation pattern within the Inlet from these measurements since all of the interior current meter records suffer flaws: either the data are dominated by sub-stall speed records (CM 4 and CM 10) or the vane is jammed (CM-5). At the Inlet entrance all the current meters at 12 m depth recorded mean flows into the Inlet while the flow at 50 m depth is out of the Inlet at CM-8 and appears to have a negligibly small component oriented along the entrance channel at CM-7. The principal direction of flow through the Inlet entrance is taken as $023^{\circ} \mathrm{T}$ below, where the mean components for flow directed into ( + ) and out of ( - ) the Inlet are summarized.


FIGURE 6. Monthly mean vectors recorded at the current meters. The dashed vectors refer to currents at 50 m depth. ${ }^{*}=$ no direction information; $* *=$ mean speed not significant.

TABLE 5

## MEAN CURRENT VECTORS AND THEIR ASSOCIATED PRECISIONS



[^0]Mean Flow Component Through the Inlet Entrance ( $\mathrm{cm} \mathrm{s}^{-1}$ )

$$
\left(+=023^{\circ} \mathrm{T}\right)
$$

Site Site Site Site

12 m
50 m

Site
8
$+3.9$
$-1.7$

Site
7
$+5.3$
+0. 6

Site
6
$+2.2$
--

The mean shear directed along $023^{\circ} \mathrm{T}$ is $4.7 \mathrm{~cm} \mathrm{~s}^{-1}$ at site 7 and $5.6 \mathrm{~cm} \mathrm{~s}^{-1}$ at site 8 between 12 and 50 m depth, in good agreement with the computed geostrophic shears (see Section II.A).

The general impression from the current meter data at the Inlet entrance is of an inward near-surface flow balanced by a slow return flow at depth and thus the circulation is oppositely directed from that obtaining in estuaries where the run-off and rainfall exceed the evaporation. The process which drives this negatively directed estuarine circulation is probably brine rejection by thickening sea ice.

In order to determine if the spatial sampling at the Inlet entrance was of sufficient density, cross correlation coefficients of the velocity component along $023^{\circ} \mathrm{T}$ were computed among the current meter data sets. The resulting correlation coefficient matrix is shown below (the $95 \%$ significance value is 0.07 ):

## CORRELATION COEFFICIENT MATRIX

CM6 CM7 (12) CM7 (50) CM8 (12) CM8 (50) CM9

| CM6 (12) | 1.00 | -0.382 | +0.461 | -0.310 | -0.070 | -0.467 |
| :--- | :---: | :---: | :---: | :---: | :---: | :---: |
| CM7 (12) |  | 1.00 | -0.547 | +0.751 | +0.445 | +0.765 |
| CM7 (50) |  |  | 1.00 | -0.579 | -0.342 | -0.672 |
| CM8 (12) |  |  |  | 1.00 | +0.404 | +0.773 |
| CM8 (50) |  |  |  |  | 1.00 | +0.473 |
| CM9 (12) |  |  |  |  |  | 1.00 |

The correlation coefficients are representative of vertical separations of 38 m and horizontal separations of 450,900 and 1350 m (the current meters were approximately evenly spaced 450 m apart horizontally). There are six categories of coefficients in the matrix:
$C(x, 0)$ representing 1 horizontal and no vertical separations
$C(2 x, 0) \quad 2$ horizontal and no vertical separations
$C(3 x, 0) \quad 3$ horizontal and no vertical separations
$\mathrm{C}(0, \mathrm{z})$ no horizontal and 1 vertical separations
$C(x, z) \quad 1$ horizontal and 1 vertical separations
$\mathrm{C}(2 \mathrm{x}, \mathrm{z}) \quad 2$ horizontal and 1 vertical separations

The values in these categories are listed below:

| $C(x, 0)$ | $C(2 x, 0)$ | $C(3 x, 0)$ | $C(0, z)$ | $C(x, z)$ | $C(2 x, z)$ |
| :--- | :--- | :--- | :--- | :--- | :--- |
| -0.38 | -0.31 | -0.46 | -0.54 | +0.44 | -0.07 |
| +0.75 | +0.76 |  | -0.40 | -0.57 | -0.67 |
| +0.77 |  |  |  | +0.46 |  |
| -0.34 |  |  |  | +0.77 |  |

There are no apparent significant differences among the average values of the 6 categories of correlation coefficients. Current meters 6 and 7 (50) are negatively correlated with all other current meters. In the case of current meter 6 this is probably due to the over-riding
effects of local topography (a long shallow sub-surface spit) and in the case of 7 (50) the outflow from the Inlet imposes a bias in direction. Generally, however, the magnitudes of the correlation coefficients are large indicating that little spatial variation exists, particularly between the current meters at sites 7,8 and 9 , at 12 m depth. One might have expected weaker correlations since the internal Rossby radius, $r_{i}$ is small, $\left(r_{i}=\frac{N H}{f}\right.$ where $N$ is the Vaisala frequency, $H$ the layer depth and $f$ the Coriolis parameter. For the Bridport Inlet entrance $r_{i}=470 \mathrm{~m}$ for $\mathrm{N}=4.4 \times 10^{-3} \mathrm{~s}^{-1}$ and $\mathrm{H}=15 \mathrm{~m}$ ).

Cross spectra among the current meters were also computed and strong coherences were found in the semi-diurnal tidal band. Only current meter $7(50)$ showed a significant phase difference ( $53^{\circ}$ ) with respect to current meters $7(12), 8(12), 8(50)$ and $9(12)$. The other current meters were in phase ( $\pm 8^{\circ}$ ) with each other in the semi-diurnal band.

## C. CURRENT PROFILES

Current speed profiles were measured at 8 locations with an ultrasonic current meter. Directions as well as speeds were determined at sites $80-5$ and CM-5. The ultrasonic current meter profiles were employed in two ways; to check the Aanderaa current meter operation and to obtain a detailed view of the vertical current profiles. The USCM profiles are reproduced in Appendix B.

As mentioned in Section II.B, the profiling current meter confirmed that erroneous directions were recorded by the moored Aanderaa instrument at CM-5 due to a malfunctioning vane. At the other current meter sites where intercomparisons were performed (CM-4, 9 and 10) the instantaneous speeds recorded by the two instruments were in good agreement when currents were above the Aanderaa stall speed of $2.2 \mathrm{~cm} \mathrm{~s}^{-1}$.

The ultrasonic current meter profiles confirm the vertical structure of the velocity field at the Inlet entrance derived from geostrophic and mass-balance considerations. That is, a minimum current speed is measured at about 25 m depth near where a flow reversal
was computed by indirect means. In addition, the ultrasonic data show a very slight decrease of velocity near the underside of the sea ice. Such a decrease is due to the interfacial stress between the land-fast ice and the fluid motion. These boundary layer effects are not noticeable at distances of more than about 5 m below the ice.

An attempt was made to fit the observed current profiles to the log-linear boundary layer structure applicable in convective regions (such as under growing sea ice) described by Turner (1973). This profile is

$$
\begin{equation*}
u=\frac{u_{t}}{k}\left(\ln \frac{\mathrm{z}}{\mathrm{z}_{\mathrm{o}}}+\alpha \frac{\mathrm{z}}{\mathrm{~L}}\right) \tag{4}
\end{equation*}
$$

where $u$ is the velocity at a distance $z$ below the ice, $u_{*}$ is the friction velocity equal to $\sqrt{\tau / \rho}$ where $\tau$ is the interfacial stress, $k$ is Von Karman's constant equal to about $0.4, z_{0}$ is the roughness length, $\alpha$ is a constant with a value of about 5.0 and L is the Obukhov length, the ratio of the kinetic energy flux due to boundary shear stress to that due to convection.

$$
\begin{equation*}
L=\frac{\bar{\rho} u_{t}^{3}}{k g \rho^{\prime} w^{\prime}} \tag{5}
\end{equation*}
$$

where $\bar{\rho}$ is the mean density in the convective regime, $g$ is the acceleration of gravity and $\overline{\rho^{\prime} w^{\top}}$ is the correlation of the density and vertical velocity fluctuations equivalent to the turbulent vertical mass transport. L is negative in unstable conditions.
$\overline{\rho^{\prime} w^{\prime}}$ can be estimated from the ice growth rate

$$
\begin{equation*}
\overline{\rho^{\prime} w^{\prime}} \simeq i \rho_{i} \Delta s \tag{6}
\end{equation*}
$$

where $i$ is the ice growth rate, $\rho_{i}$ is the density of the ice and $\Delta s$ is the difference in salinity between the ice and the sea water. $u_{*}$ and $a_{o}$ can be determined with the profile data.

The results of the fit showed the Obukhov length to be between 16 and 45 metres. Convective effects are considered to be important only when the mixed layer depth is comparable to $L$, and since the observed depth of the truly mixed layer is only about 2 or 3 m , the velocity profile is not greatly affected by convection.

Resulting values of $u_{*}$ are about $0.9 \mathrm{~cm} \mathrm{~s}^{-1}$ and the drag coefficient defined as

$$
\begin{equation*}
C_{D}=\frac{u_{t}^{2}}{U^{2}} \tag{7}
\end{equation*}
$$

where $U$ is the velocity far from the boundary, is about $3 \times 10^{-3}$. These values are in fair agreement with those of Shirasawa and Langleben (1976) who found $u_{*}=0.97 \mathrm{~cm} \mathrm{~s}^{-1}$ and $C_{D}=0.97 \times 10^{-3}$.

The implication of the strong vertical gradient of horizontal velocities concentrated near the boundary is that these effects may be ignored in mass transport computations. The observed profiles show that current speeds are not appreciably decreased by the ice-water stress at distances greater than 1 or 2 metres below the ice.

Both dynamic computations, whereby currents are computed from the distribution of mass, and direct current measurements within Bridport Inlet are inadequate to formulate a clear picture of the circulation. In the case of dynamic computations, the horizontal density gradients are too small to accurately determine the flow velocities and in the case of the direct current measurements the speeds are often so low as to be undetectable by the Aanderaa current meter. As an alternative to these methods, plots of the distribution of properties were constructed which ultimately pointed to a mechanism for driving the mean circulation.

Horizontal plots of temperature and salinity at $5,10,20,30$ and 40 m were constructed from the average CTD data at each site. (Below 40 m depth the horizontal gradients were too weak to yield any details of the property fields.) These plots are shown in Figures 7a-e.

The isohalines on surfaces shallower than 20 m imply an intrusion of lower salinity water from Viscount Melville Sound into Bridport Inlet. At 30 m depth and below, the salinity plots weakly indicate that the water of Bridport Inlet have a higher salinity than those of Viscount Melville Sound but the direction of flow is not so dramatically portrayed as at shallower depths. There is an indication of a pooling of higher salinity water in the central, deep region of the Inlet from which one might infer a slow cyclonic circulation at depth. The temperature and salinity signals below 30 m depth are, however, very weak, barely exceeding the precision of the measurements. The plots therefore are only suggestive of water motions.

A vertical section of salinity, running nearly north-south through the Inlet entrance was constructed and is shown in Figure 8. Station 80-3 is coincident with CM-7 located between the two spits at the entrance. The shapes of the isohalines suggest an inflow of fresher water near the surface, and an outflow of higher salinity water at depth. The nearly horizontal 32.74 isopyenal at the entrance is located at 24 m depth indicating that this depth is, roughly, the boundary between inflow and outflow. The water column near the north shore of the inlet is homogeneous ( $\pm 0.007 \%$ ) below 20 m depth and between 20 and 60 metres it is homogeneous within $\pm 0.003 \%$. This


FIGURE 7a. Horizontal plots of temperature and salinity at 5 m depth.


FIGURE 7b. Horizontal plots of temperature and salinity at 10 m depth.


FIGURE 7c. Horizontal plots of temperature and salinity at 20 m depth.


FIGURE 7d. Horizontal plots of temperature and salinity at 30 m depth.


FIGURE 7e, Horizontal plots of temperature and salinity at 40 m depth.


FIGURE 8. Vertical salinity section running nearly north south through the Inlet entrance.
structure is indicative of the presence of convective processes within the inlet, probably driven by the rejection of brine from growing sea ice.

The modification of the mass field due to brine rejection can be graphically portrayed by a method conceived by Melling (personal communication). The difference between the in-situ temperature and the surface freezing point at the in-situ salinity is contoured on a vertical section. Waters cooled to or below the freezing point at the surface are indicated by values of $T-T_{f} \leq 0$. The formula used for computation of the freezing point at 1 bar (the ocean surface) is taken from Millero (1977);

$$
\begin{equation*}
\mathrm{T}_{\mathrm{f}}=-.0575 \mathrm{~S}+1.710523 \times 10^{-3} \mathrm{~S}^{3 / 2}-2.154996 \times 10^{-4} \mathrm{~S}^{2} \tag{8}
\end{equation*}
$$

where $T_{f}$ is the freezing point and $S$ is the salinity.
A vertical section of $\mathrm{T}-\mathrm{T}_{\mathrm{f}}$ from Bridport Inlet in 1980 is shown in Figure 9. Except for a thin layer less than 2 metres thick below the ice, all the water at, or below its surface freezing point is found at depths greater than 10 m near the north shore, and at depths greater than 30 m at the entrance. There is however, a strong suggestion of the upper surface of the deep "cold" water mass shoaling toward the north shore and the $T-T_{f}=0$ isolines connecting in shallower water, just a few hundred metres north of station $80-1$.

The section indicates that water warmer than $0.06^{\circ} \mathrm{C}$ above its surface freezing point enters the inlet at 8 to 12 m depth. As this relatively warm water flows northward it loses heat and approaches the surface freezing point. In order to balance the intrusion of relatively warm water at shallow depths there must be an outflow of colder water below it.

The estuarine circulation described may be driven by rejection of brine and the establishment of denser waters within the Inlet than are found in Viscount Melville Sound. Such an estuarine circulation is supported by the geostrophic shear computed at the Inlet entrance.

If the salinity distribution shown in Figure 8 is due to a quasistationary process, then a balance between vertical diffusion and horizontal


FIGURE 9. Vertical north-south section of the difference between in-situ temperature ( $T$ ) and the freezing point temperature at atmospheric pressure and in-situ salinity ( $\mathrm{T}_{\mathrm{f}}$ ) for 1980 (Millidegrees).
advection might be present. Such a balance can be represented as

$$
\begin{equation*}
u \frac{\partial s}{\partial x}=K_{z} \quad \frac{\partial^{2} s}{\partial z^{2}} \tag{9}
\end{equation*}
$$

where $u$ is the horizontal velocity in the $x$ direction and $K_{z}$ is the eddy diffusivity of salt. If a velocity of $1 \mathrm{~cm} \mathrm{~s}^{-1}$ is assumed, then the necessary value of $K_{z}$ is approximately $10^{-3} \mathrm{~m}^{2} \mathrm{~s}^{-1}$. This value is an order of magnitude too high for the stratification present in Bridport Inlet $\left(\mathrm{N}^{2} \approx 5 \times 10^{-5} \mathrm{~s}^{-2}\right)$ when compared with values determined in Agfardiikavsa Fjord by the Danish Hydraulic Institute which was about $10^{-4} \mathrm{~m}^{2} \mathrm{~s}^{-1}$ (Lewis and Perkin, 1982). If the rejection of brine by growing sea ice is included in the processes which maintain the salinity field, then (9) may be rewritten as

$$
\begin{equation*}
u \frac{\partial s}{\partial x}=K_{z} \frac{\partial^{2} s}{\partial z^{2}}+i \frac{\rho_{i}}{\rho_{w}}\left(S_{w}-S_{i}\right) / h \tag{10}
\end{equation*}
$$

where $i$ is the growth rate of the ice $\left(=1 \mathrm{~cm} \mathrm{~d}^{-1}\right)$, the subscripts $i$ and $w$ signify ice and water respectively, $\rho$ is density and $h$ is the depth of the inward moving layer. The numerical values of the terms in (10) are

$$
1.5 \times 10^{-7} \% \mathrm{oo} \mathrm{~s}^{-1}=\mathrm{K}_{\mathrm{z}}\left(3 \times 10^{-4}\right)+1.1 \times 10^{-7}
$$

which yields a value of $K_{z}=1.3 \times 10^{-4}$, in better agreement with values from Agfardlakavsa Fjord.

In (10), the rejection of brine by ice growth is probably very nearly in balance with the advection but is almost certainly much larger than the diffusion of salt upward from the deeper layer. The system is thus advectively controlled.

The process by which the brine increases the salinity of the inward flowing surface layer is not revealed by the data. The mixed layer depths may be as small as 1 or 2 metres, while the classical concept of the density structure under growing sea ice has a mixed layer extending below the ice over the entire depth where convective processes dominate. In Bridport Inlet it is likely that the convection is an intermittent
process. Examination of all the instantaneous CTD profiles did not, however, yield any convincing evidence of intermittent convection.

The degree of stratification can also be scrutinized by evaluating the mass of salt required to be added at the surface to mix the water column. Figure 10 shows the average temperature, salinity and density profiles at Site 80-1 near the northern shore of the Inlet. The mixed layer is entirely absent immediately below the ice while the water column is nearly completely homogeneous below a depth of about 20 m ....The Väisälă frequency below 20 m depth is less than $2 \times 10^{-3} \mathrm{~s}^{-1} \pm 10^{-3} \mathrm{~s}{ }^{-1}$ (Figure 11). The mass of salt, $M_{S}$, which must be added to the water column by sea ice growth in order to cause convection to 50 m depth (approximate sill depth) can be computed from

$$
M_{S}=0 w \sum_{j=0}^{50} S_{50}-S_{j} \Delta z
$$

where $S_{50}$ is the salinity at 50 m depth, $\mathrm{S}_{\mathrm{j}}$ are the salinities at depths $j$ and $\Delta z$ is the depth interval taken to be one metre.

The required salt input from ice formation to destabilize the water column, and the time required (at an ice growth rate of $1 \mathrm{~cm} \mathrm{~d}{ }^{-1}$ ) to produce that amount of salt is graphed in Figure 12, versus site location. At the Inlet entrance approximately 8 days of ice growth would mix the water column while at site $80-1$ just over one day's ice growth would be sufficient. If this trend of increasing vertical homogeneity is extrapolated into the shallower waters north of site $80-1$, then one can conjecture that the water column is completely homogeneous about 1 km north of site $80-1$ in a depth of about 50 m . The apparent stability associated with the upper water column is therefore probably not reflective of the convective process which is known to be occurring in the Inlet.

In summary, the property distributions on a vertical section running nearly north-south through the Inlet entrance imply a mean circulation characteristic of a negative estuary. Flow appears to be inward in the upper 25 m and outward below that depth. The upper layer salinity is increased as the fluid moves northward by brine rejection until the water column is completely destabilized near the north shore in depths of about


FIGURE 10. Average $\mathrm{T}, \mathrm{S}$, and $\sigma_{\text {t }}$ profiles of $80-1$. The envelopes about each dotted mean profile span all the values recorded.


FIGURE 11. $\begin{aligned} & \text { Profile of the Vaisälä frequency } \\ & \text { horizontal bars indicate the pre- }\end{aligned}\left(\begin{array}{l}\frac{\sqrt{g}}{\rho} \frac{\partial \rho}{\partial z}\end{array}\right)$ at site $80-1$. The $\begin{aligned} & \text { cision of the }\end{aligned}$ computed values at several depths.


FIGURE 12. The amount of salt required for convection to 50 m depth and the time required for ice growth to reject sufficient brine for convection. Section north-south through the Inlet entrance.

50 m . At this location the relatively high salinity freezing point water is advected downward to depths below about 25 m and flows south (outward) through the Inlet entrance. The southerly flow is probably largely confined to the depths above sill depth which is approximately 50 m . The salt balance in the Inlet is advectively dominated.

## Cyclonic Flow?

The north-south section cannot address the possibility which is weakly indicated by the horizontal temperature and salinity plots, i.e. the presence of a slow cyclonic flow. A plot of $T-T_{f}$ at stations taken in series around the periphery of the Inlet in a cyclonic sense is presented in Figure 13. The flow is taken to begin in Viscount Melville Sound, follow a course past CM6 on the east side of the entrance steered by the Coriolis force, and then proceed cyclonically around the Inlet exiting towards the west of the entrance at CM9. The concept of weak cyclonic circulation is supported by Figure 13. The isoline $T-T_{f}=0$ generally shoals in the cyclonic direction, intersects the surface at CM-10 and the fluid hypothesized to leave the Inlet at CM9 is at or below its surface freezing point.

The weak cyclonic circulation is, however, most probably a response to the pressure gradient established between the Inlet entrance and its interior. The balance of forces maintaining the estuarine flow is analysed in Section IV where a vertical velocity profile and a salt balance is computed.

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FIGURE 13. Vertical section of $T-T_{f}$ counter clockwise around the periphery of the Inlet (millidegrees).

## A. DYNAMICS

The current meter measurements, the computed geostrophic currents at the entrance, and the property distributions suggest that the mean flow (over the period of the measurements) is into the inlet near the surface and outward at depth. The sense of this circulation is suggestive of a "negative" estuary in which processes of densification predominate. Evaporation is the usual agent for densification in lower latitudes, but in the Arctic brine rejection from growing sea ice will increase the density of the underlying waters. If a situation occurs where there is a horizontal difference in the rate of ice growth then a horizontal density gradient can be established and a resulting horizontal. pressure gradient. Flows would result near the surface from the less dense (less ice growth) region toward the more dense region, with oppositely directed return flows at depth. Depending upon the horizontal and vertical scales of the system, rotation and friction respectively, may become important.

In order to investigate the possibility of a negative estuarine circulation driven by ice growth, the density structures within Bridport Inlet and at the entrance (about 6 km away) were compared. The data from the average of 5 profiles at site $80-1$ within the Inlet and the average of 31 profiles at site $80-3$ at the Inlet entrance were used to compute dynamic heights at these stations. The differences between dynamic heights at each depth were used to compute the pressure difference at 1 m depth intervals between the stations.

If it is assumed for a first coarse approximation that friction can be ignored, that the estuarine circulation is non-rotational in character, and that the flow is steady, then the equation of motion for flow through the entrance reduces to

$$
\begin{equation*}
u \frac{\partial u}{\partial x}=-\frac{1}{\rho} \cdot \frac{\partial P}{\partial x} \tag{12}
\end{equation*}
$$

The balance of forces in (12) is between the non-linear acceleration and the pressure gradient. Such a balance is more commonly seen in flows with small radii of curvature such as tornadoes, than in estuaries. However, the situation in Bridport Inlet does intuitively justify this balance. The mean currents within the inlet are very weak ( $\sim 1 \mathrm{~cm} \mathrm{~s}^{-1}$ ) (see Section II) while they are about 3 to $8 \mathrm{~cm} \mathrm{~s}^{-1}$ at the entrance. One can therefore visualize stream lines widely spread within the inlet and converging at the entrance. The most important missing terms are friction and the Coriolis force but it can be estimated that frictional terms are small because of the relatively large depth and cross stream velocities are limited by the narrow entrance.

The relative size of the frictional term can be estimated by comparing it with the acceleration term. If friction is included in (12) then

$$
\begin{equation*}
u \frac{\partial u}{\partial x}=-\frac{1}{\rho} \frac{\partial P}{\partial x}+\frac{1}{\rho} \frac{\partial}{\partial z}{ }^{\tau} x z \tag{13}
\end{equation*}
$$

and we seek to compare the first and last terms

$$
\frac{u \frac{\partial u}{\partial x}}{\frac{1}{\rho} \frac{\partial}{\partial z}{ }^{\tau} x z} \sim \frac{u^{2 /} x}{\frac{\rho C_{D} U^{2}}{\rho H}}=\frac{H}{X C_{D}}
$$

where $\tau_{x z}$ is the stress on $z$ planes in the $x$ direction, $u$ is the velocity scale, $H$ is the depth and $X$ the horizontal scale.

$$
\frac{\mathrm{H}}{X C_{D}}=\frac{100 \mathrm{~m}}{6000 \mathrm{~m} 2 \times 10^{-3}}=\frac{100}{12}=8
$$

The acceleration is thus approximately an order of magnitude larger than the frictional forces.

While it is probably justifiable to ignore the cross channel velocities resulting from rotation at the Inlet entrance, the interior of the Inlet presents no such topographical constraint. If $x$ is the primary direction of flow, then the equation of motion in the crossstream direction would be for the above assumptions

$$
\begin{equation*}
f u=-\frac{1}{\rho} \frac{\partial p}{\partial y} \tag{14}
\end{equation*}
$$

The two equations (12) and (14) describe a flow in the $x$ direction, accelerated through the Inlet entrance by pressure gradients in the $x$ direction. The Coriolis force associated with this flow is balanced by cross stream pressure gradients. Thus the geostrophic shear at the Inlet entrance is expected to be and is in fairly good agreement with the current meter data. The cross stream pressure gradients are established in response to the density driven estuarine flow. A similar balance of forces was proposed by Garrett and Petrie (1981) for the Strait of Belle Isle where both tidal oscillations and mean flows were investigated.

The right hand side of equation (12) is known at each depth and if we assume that the flow is accelerated across the Inlet through the entrance then at any level

$$
\begin{equation*}
\frac{u^{2}}{2}=\Delta X\left(-\frac{1}{\rho} \frac{\partial p}{\partial x}\right) \tag{15}
\end{equation*}
$$

The velocities resulting from (15) were computed for each depth between 3 m and 95 m ( 3 m depth is approximately 1 m below the ice). The velocity profile resulting from this computation is shown in Figure 14.

In order to maintain a long-term constant sea level, continuity must be satisfied and the flow into the Inlet must balance the flow out of the Inlet. This condition is clearly not satisfied by the profile of Figure 14. If a constant sea surface slope is superimposed upon the internal pressure gradients due to horizontal density gradients


FIGURE 14. Computed thermohaline velocity profile $\left(u=\frac{\sqrt{2 \Delta x}}{\rho} \cdot \frac{\mathrm{~d} \rho}{\mathrm{dx}}\right)$ unbalanced
for continuity.


FIGURE 15. Continuity balanced thermohaline flow, $u_{\mathrm{T}}$.
then a velocity constant with depth will be added to the thermohaline flow. This constant flow $u_{o}$ was computed by integrating the thermohaline velocity $u$ with depth:

$$
\begin{equation*}
\sum_{j=3}^{95} u_{j} \Delta Z+u_{o} H=0 \tag{16}
\end{equation*}
$$

In (16) $H$ is the total depth, the $u_{j}$ are the thermohaline velocities computed at depth intervals ( $\Delta Z$ ) of 1 metre. $u_{o}$ resulting from (4) is $8.90 \mathrm{~cm} \mathrm{~s}^{-1}$ into the inlet.

The total velocity profile $u+u_{o}=u_{T}$ is plotted versus depth in Figure 15.

The velocity profile of Figure 15 is in rough quantitative agreement with the mean currents measured at current meter stations 7 and 8 as well as with the geostrophic current shear computed from stations across the Inlet entrance. The validity of (15) is therefore relatively well supported. The concept of the process driving the mean circulation through the Inlet entrance presented in Section III is supported by these computations: Horizontal density differences between the Inlet and the Viscount Melville Sound accelerate the flow through the entrance. In response to this flow, geostrophic balance is established in the Inlet entrance with the vertical shear obeying the thermal wind relations.

The ratio of equation (14) to (12) yields


This comparison implies that in the absence of cross-channel flow, the cross-stream pressure gradient at the entrance is approximately 5 times as large as the along-stream pressure gradient.
B. SALT FLUX

Further confirmation of the concept of estuarine flow driven by brine rejection as well as information on the mechanisms of salt balance were obtained by computing the salt flux directed through the Inlet entrance. The velocity computed from (15) was multiplied by the difference in salinity between stations $80-1$ and $80-3$ at each depth. A profile of the salinity difference is shown in Figure 16 while a profile of the salt flux is shown in Figure 17.

In the upper 28 m of the water column the salinity is greater within the Inlet than at the entrance. Below 29 m the salinity is greater at the entrance than within. The salt flux profile of Figure 18, however, is more complicated; there are four flux conditions which apply in 4 depth intervals:

1. Positive salt flux (out of the Inlet) due to inward flow and higher salinity within the Inlet $3<d<17 \mathrm{~m}$
2. Negative salt flux due to outward flow and higher salinity within the Inlet $17<\mathrm{d} \leq 29 \mathrm{~m}$
3. Positive salt flux due to outward flow and lower salinity within the In1et $29<d \leq 67 \mathrm{~m}$
4. Negative salt flux due to inward flow and lower salinity within the Inlet $67<\mathrm{d} \leq 95 \mathrm{~m}$

It is clear, however, that the contribution to the salt flux from the flow below 30 or 40 metres depth is negligible.

The salt flux at each depth was multiplied by the entrance width at that depth to compute the total salt flux.

$$
\begin{equation*}
\text { Salt flux }=\frac{\rho}{1000} \sum_{3}^{95}\left(u_{j}+u_{o}\right) \Delta S_{j} W_{j} \Delta Z \tag{17}
\end{equation*}
$$

In (17) $\rho$ is the mean density of sea water, $\Delta S_{j}$ are the salinity difference at each depth between station $80-3$ and $80-1$, and $W_{j}$ are the entrance widths at each depth.


FIGURE 16. Salinity difference, $\Delta \mathrm{S}$, between station $80-1$ and $80-3$ versus depth.


FIGURE 17. Profile of computed salt flux, $u_{\mathrm{T}} \Delta \mathrm{S}$, versus depth.

The computed salt flux is $200 \mathrm{Kg} \mathrm{s}^{-1}$ out of Bridport Inlet. The sign of this result is expected since relatively shallow areas in the Arctic are thought to be sources of higher salinity water for the rest of the Arctic Basin (Melling $\&$ Lewis, 1982). The source of the high salinity water is brine rejection by growing sea ice during the winter months. The amount of salt rejected by the growing sea ice can also be estimated:

$$
\begin{equation*}
\text { Salt rejection }=\mathrm{dA} \rho_{i}\left(S_{W}-S_{i}\right) \tag{18}
\end{equation*}
$$

where $d$ is the thickness of the ice, $A$ is the surface area of the Inlet, $\rho_{i}$ is the density of ice, $S_{W}$ is the salinity of the water from which the ice congeals, and $S_{i}$ is the salinity of the sea ice. (18) yields a figure of $4.52 \times 10^{9} \mathrm{Kg}$ of salt rejected by growing ice if the ice salinity is $5 \%$ (Greisman, 1978). If we assume that ice formation began in Bridport Inlet on 1 September, then the time elapsed from the beginning of the ice formation to the time of the measurements was 222 days or $1.92 \times 10^{7} \mathrm{~s}$. The average rate of salt rejected as brine is therefore

$$
\text { Average salt rejection rate }=\frac{4.52 \times 10^{9} \mathrm{Kg}}{1.92 \times 10^{7} \mathrm{~s}}=236 \mathrm{Kg} \mathrm{~s}^{-1}
$$

This figure agrees with the computed salt balance to within $7 \%$ ! The very close agreement is somewhat fortuitous, however, it lends credence to the dynamical assumption of the advective acceleration balancing the pressure gradient force.

## C. COMPARISON WITH 1979

Data collected in 1979 (Lake $1979 \mathrm{~b}, \mathrm{c}$ ) were examined to determine if the ice growth driven circulation was present in that year. The situation in 1979 was, however, very different from 1980. Figure 18 is a vertical section of $T-T_{f}$ for the 1979 data. There is little resemblance to Figure 9 from the 1980 data. The layer which is below the surface freezing point is 10 to 20 m deep in 1979 compared with


FIGURE 18. Vertical north-south section of $T-T_{f}$, 1979 (Millidegrees).
less than 4 m deep in 1980. In 1979 the deep water of the Inlet (below 30 m ) was about $0.2^{\circ} \mathrm{C}$ above the surface freezing point while in 1980 the deep water was about $0.005^{\circ}$ below the surface freezing point. Apparently convective processes had penetrated to the bottom in 1980 but not in 1979. This supposition is confirmed by comparing the CTD profiles from site 80-2 (1980) with that from site 1-79 (1979) in Figure 19. These casts were both made at the centre of the Inlet. The mixed layer depth was clearly 25 m in 1979, while in 1980, if strong intermittent convection is assumed, the mixed layer extends to the bottom and if the assumption is not made, the mixed layer depth is about 4 m . The salinity at 130 m depth in 1979 is 32.616 which is lower than the salinity found at the surface in 1980. Finally, the overall stability (in terms of the Vaisala frequency, $N$ ) of the water column between 3 m and 130 m depth is

$$
\mathrm{N}^{2}=-\frac{\mathrm{g}}{\rho} \frac{\rho_{130}-\rho_{3}}{127 \mathrm{~m}}
$$

and was $2.1 \times 10^{-6} \mathrm{~s}^{-2}$ in 1979 and $0.9 \times 10^{-6} \mathrm{~s}^{-2}$ in 1980.
The salt flux computation described in the previous section was performed with the 1979 data and yielded a figure of $130 \mathrm{Kg} \mathrm{s}^{-1}$ of outward transport of salt; roughly half the value computed with the 1980 data.

The difference in the year-to-year pictures is due to processes in Viscount Melville Sound which ultimately determine the mass structure of Bridport Inlet. At stations outside the Inlet the surface salinity was $0.35^{\circ} / 00$ ligher in 1980 than in 1979 and the vertical density gradient was about half as large in 1980 as in 1979.

It would seem that inter-annual differences are so large as to eclipse the variations over a several month interval. The important point here is that a single year's data would not have been representative of the system. This last point is gradually being accepted as pertinent to most areas of the Archipelago and, indeed the Arctic Ocean.


FIGURE 19. Mean CTD profiles from sites $80-2$ and 1-79.

## v TIDAL OSCILLATIONS

The tidal heights in Bridport Inlet had been observed in 1978 (Greisman, 1978) and in 1979 (Lake, 1979 c).

During the 1980 field program two tide gauges were installed: one near the inlet entrance and the other near the eastern shore of the inlet (see Figure 1 for tide gauge locations). The two tide gauge records are in good agreement with each other as well as with the record from 1978. There is an indication of a phase lag between the two gauges which indicates that the tide wave has a progressive component and that some tidal energy is transported into and dissipated within the inlet. Table 6 lists the major tidal constituents, their amplitude and Greenwich phase as measured in 1980.

The current meter data were subjected to tidal stream analyses, and the results for the major constituents are shown in Table 7. The larger tidal signal at the Inlet entrance makes these data sets more attractive than the interior current data for scrutiny and comparison with tidal height oscillations.

At current meter sites 7 and 8 there is significant difference in the amplitude of the tidal currents measured at depths of 10 and 50 metres. In addition, there is a nearly $180^{\circ}$ phase difference between the 10 m and 50 m currents at site 7 while the currents at site 8 are nearly in phase. Foreman (1978) explains that the tidal stream analysis routine can be biased by $180^{\circ}$ when the computed phases are near $0^{\circ}$ or $180^{\circ}$. Since a $180^{\circ}$ phase difference between shallow and deep currents might be ascribed to the presence of internal waves, this possibility was investigated.

It was first assumed that there was a $180^{\circ}$ error in the phase of the semi-diurnal components at the 50 m depth current meter at site 8. Applying this correction to the analysis resulted in a coherent picture of the tidal oscillation at the entrance: stronger shallow currents in phase across the entrance with weaker deep currents also roughly in phase with each other and out of phase with the shallow flow.

Tidal current profiles may be regarded as the sum of barotropic and baroclinic components, the former being uniform through the water column and the latter being determined by the vertical stratification, the latitude and

TABLE 6a
TIDAL HEIGHT ANALYSIS
TIDE GAUGE WEST OF ENTRANCE SPIT

| Name | Speed | Amp | G |
| :---: | :---: | :---: | :---: |
| Z0 | 0.0000000000 | 36.1816 | . 00 |
| MSF | 0.0028219327 | 0.0170 | 212.98 |
| 2Q1 | 0.0367063506 | 0.0029 | 164.10 |
| Q1 | 0.0372185027 | 0.0138 | 196.98 |
| 01 | 0.0387306544 | 0.0383 | 266.59 |
| N01 | 0.0402685944 | 0.0013 | 241.17 |
| K1 | 0.0417807461 | 0.0586 | 359.94 |
| J1 | 0.0432928982 | 0.0088 | 60.85 |
| 001 | 0.0448308382 | 0.0058 | 170.65 |
| UPS 1 | 0.0463429899 | 0.0048 | 228.46 |
| N2 | 0.9789992493 | 0.0683 | 207.74 |
| M2 | 0.0805114005 | 0.3673 | 251.25 |
| S2 | 0.0833333330 | 0.2084 | 304.36 |
| ETA2 | 0.0850736443 | 0.0094 | 315.38 |
| MO3 | 0.1192420553 | 0.0040 | 83.14 |
| Mc | 0.1207671007 | 0.0023 | 308.42 |
| MK3 | 0.1222921470 | 0.0004 | 245.53 |
| SK3 | 0.1251140796 | 0.0009 | 160.81 |
| MN4 | 0.1595106497 | 0.0010 | 270.10 |
| M4 | 0.1610228010 | 0.0003 | 53.06 |
| MS4 | 0.1638447344 | 0.0005 | 326.20 |
| S4 | 0.1666666660 | 0.0004 | 315.35 |
| 2MK5 | 0.2028035484 | 0.0003 | 230.97 |
| 2SK5 | 0.2084474135 | 0.0014 | 30.61 |
| 2MN6 | 0.2400220502 | 0.0006 | 157.01 |
| M6 | 0.2415342014 | 0.0007 | 216.07 |
| 2MS6 | 0.2443561349 | 0.0013 | 294.40 |
| 2SM6 | 0.2471780665 | 0.0005 | 300.53 |
| 3MK7 | 0.2833149470 | 0.0008 | 97.68 |
| M8 | 0.3220456019 | 0.0003 | 299.29 |

TABLE 6b
TIDAL HEIGHT ANALYSIS
INTERIOR TIDE GAUGE
Name Speed $\quad \therefore \quad$ Amp $\quad$ G

| zo | 0.000000000 | 49.7912 | . 00 |
| :---: | :---: | :---: | :---: |
| MM | 0.0015121518 | 0.0331 | 225.91 |
| MSF | 0.0028219327 | 0.0159 | 205.71 |
| ALPI | 0.0343965697 | 0.0039 | 131.05 |
| 2 Q1 | . 0.0357063506 | 0.0041 | 161.03 |
| Q1 | 0.0372185027 | 0.0139 | 202.02 |
| 01 | 0.0387306544 | 0.0373 | 268.53 |
| NO1 | 0.0402685944 | 0.0024 | 273.44 |
| K1 | 0.0417807461 | 0.0571 | 1.38 |
| J1 | 0.0432928982 | 0.0059 | 38.79 |
| 001 | 0.0448308382 | 0.0058 | 160.76 |
| UPS1 | 0.0463429899 | 0.0050 | 225.35 |
| EPS2 | 0.0761773158 | 0.0012 | 193.85 |
| MU2 | 0.0776894679 | 0.0159 | 201.2] |
| N2 | 0.0789992493 | 0.0684 | 211.63 |
| M2 | 0.0805114005 | 0.3638 | 251.66 |
| L2 | 0.0820235526 | 0.0202 | 322.95 |
| S2 | 0.0833333330 | 0.2111 | 305.84 |
| ETA2 | 0.0850736443 | 0.0092 | 299.45 |
| M03 | 0.1192420553 | 0.0038 | 81.95 |
| M3 | 0.1207671007 | 0.0017 | 290.29 |
| MK3 | 0.1222921470 | 0.0007 | 249.33 |
| SK3 | 0.1251140796 | 0.0012 | 169:02 |
| MN4 | 0.1595106497 | 0.0007 | 357.06 |
| M4 | 0.1610228010 | 0.0011 | 53.30 |
| SN4 | 0.1623325814 | 0.0010 | 222.27 |
| MS4 | 0.1638447344 | 0.0006 | 348.83 |
| S4 | 0.1666666660 | 0.0012 | 108.10 |
| $2 \mathrm{MK5}$ | 0.2028035484 | 0.0007 | 357.41 |

TABLE 6b continued

| Name | Speed | Amp | $G$ |
| :--- | :---: | :---: | :---: |
| 2SK5 | 0.2084474135 | 0.0013 | 155.58 |
| 2MN6 | 0.2400220502 | 0.0007 | 176.91 |
| M6 | 0.2415342014 | 0.0009 | 192.63 |
| 2MS6 | 0.2443561349 | 0.0012 | 277.82 |
| 2SM6 | 0.2471780665 | 0.0007 | 282.82 |
| 3MK7 | 0.2833149470 | 0.0006 | 251.86 |
| M8 | 0.3220456019 | 0.0001 | 102.88 |

TABLE 7

| Current <br> Meter | Depth | Constituent | Speed <br> (cyc/hr) | Major Axis $\mathrm{cm} \mathrm{s}^{-1}$ | Minor Axis $\mathrm{cm} \mathrm{s}^{-1}$ | Inclination * | Greenwich Phase |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| CM4 | 12 | No tidal constituents $>0.5 \mathrm{~cm} \mathrm{~s}^{-1}$ |  |  |  |  |  |
| CM5 | 12 | Direction vane malfunctioned |  |  |  |  |  |
| CM6 | 12 | MSF | . 002822 | 1.6 | -0.3 | $44^{\circ}$ | $196^{\circ}$ |
|  |  | N2 | . 078999 | 2.2 | 0.1 | $57^{\circ}$ | $359{ }^{\circ}$ |
|  |  | M2 | . 080511 | 4.2 | 0.6 | $62^{\circ}$ | $17^{\circ}$ |
|  |  | S2 | . 083333 | 2.3 | 0.4 | $58^{\circ}$ | $96^{\circ}$ |
|  |  | ETA. | . 085074 | 2.1 | -0.2 | $54^{\circ}$ | $359{ }^{\circ}$ |
|  |  | M4 | . 161023 | 1.0 | 0.0 | $25^{\circ}$ | $278{ }^{\circ}$ |
| CM7 | 12 | MSF |  | 1.2 | -0.3 | $16^{\circ}$ | $214{ }^{\circ}$ |
|  |  | M2 |  | 5.4 | -0.4 | $56^{\circ}$ | $165^{\circ}$ |
|  |  | S2 |  | 3.2 | -0.4 | $56^{\circ}$ | $225^{\circ}$ |
|  |  | M4 |  | 1.0 | -0.0 | $13^{\circ}$ | $201{ }^{\circ}$ |
|  |  | MS4 | . 163845 | 1.0 | -0.2 | $12^{\circ}$ | $270^{\circ}$ |
| CM7 | 50 | M2 |  | 2.3 | -0.2 | $116^{\circ}$ | $356{ }^{\circ}$ |
|  |  | SZ |  | 1.3 | -0.2 | $111^{\circ}$ | $51^{\circ}$ |
| CM* | 12. | MSF |  | 1.5 | -0.3 | $108^{\circ}$ | $39^{\circ}$ |
|  |  | M2 |  | 3.9 | -1.1 | $45^{\circ}$ | $162{ }^{\circ}$ |
|  |  | S2 |  | 2.2 | -0.7 | $58^{\circ}$ | $229{ }^{\circ}$ |
| CM8 | 50 | M2 |  | 2.5 | -0.6 | $10^{\circ}$ | $168{ }^{\circ}$ |
|  |  | S2 |  | 1.9 | -0.5 | $16^{\circ}$ | $215^{\circ}$ |
| CM9 | 12 | MSF |  | 1.1 | 0.3 | $117^{\circ}$ | $47^{\circ}$ |
|  |  | M2 |  | 4.2 | -0.6 | $75^{\circ}$ | $164{ }^{\circ}$ |
|  |  | S2 |  | 2.4 | -0.1 | $83^{\circ}$ | $221{ }^{\circ}$ |
| CM. 10 | 12 | No tidal | constituen | ts $>0.2$ | $\mathrm{cm} \mathrm{s}^{-1}$ |  |  |

* Inclination in degreesmeasured counter-clockwise from east (mathematical convention)


FIGURE 20, Profiles of the first internal mode at tidal frequency at the Inlet entrance.
the frequency of oscillation. In order to separate the two components, current measurements are required at several depths along with the mass structure which is obtainable from CTD casts.

Internal wave mode profiles were computed for several conditions with the Coriolis parameter set to zero. This limitation is justified in narrow channels where rotational effects are negligible but is questionable in the entrance to Bridport Inlet where, although only 2 km wide, the weak stratification yields an internal Rossby radius of deformation of only about 500 m . However, no internal waves of tidal period on a rotating earth can theoretically exist at this latitude so rotational effects cannot be included in any case in mode profile computations.

Several mode structure computations were performed using the averages of all the CTD casts made at sites 7 and 8 . It was found that the computations were extremely insensitive to varying the tidal frequency from semidiurnal to diurnal. On the other hand the computations were extremely sensitive to the magnitude of the mean shear. Two mean shear profiles were used: the first reflecting the average currents recorded at 10 and 50 m depth and the second, more detailed profile, was the geostrophic shear computed by the dynamic method.

The resulting mode structures for sites 7 and 8 in the absence of shear and with geostrophic shear are shown in Figure 20. These profiles represent the normalized amplitude of the first mode velocity oscillations.

The magnitude of the first mode baroclinic and barotropic oscillations can be computed at each tidal frequency by solving a pair of simultaneous equations:

$$
\begin{align*}
& X \sin \left(\omega t-\phi_{0}\right)+\alpha_{1} Y \sin \left(\omega t-\phi_{1}\right)=R_{1} \sin \left(\omega t-\phi_{R_{1}}\right)  \tag{19a}\\
& X \sin \left(\omega t-\phi_{0}\right)+\alpha_{2} Y \sin \left(\omega t-\phi_{1}\right)=R_{2} \sin \left(\omega t-\phi_{R_{2}}\right) \tag{19b}
\end{align*}
$$

where $\mathrm{X}=$ the ampitude of the barotropic tidal velocity oscillation (uniform with depth)
$Y=$ the amplitude of the first mode baroclinic velocity oscillation

$$
\begin{aligned}
\mathrm{R}_{1,2}= & \text { the measured tidal oscillation at the current meters } \\
\alpha_{1,2}= & \text { the magnitude of the mode structure at the depth of the current } \\
& \text { meters } \\
\omega= & \text { the frequency of the tidal oscillation } \\
\phi_{0}= & \text { the Greenwich phase of the barotropic component } \\
\phi_{1}= & \text { the Greenwich phase of the baroclinic component } \\
\phi_{R_{1,2}}= & \text { the Greenwich phase of the measured tidal oscillations at the }
\end{aligned}
$$

The results of the modal decomposition for the semi-diurnal components at site 7 are presented below:

|  | lst baroclinic |  | Barotropic |  | Tidal Height |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | AMP $\left(\mathrm{cm} \mathrm{s}^{-1}\right)$ | Phase | AMP $\left(\mathrm{cm} \mathrm{s}^{-1}\right)$ | Phase | AMP $(\mathrm{cm})$ | Phase |
| $\mathrm{M}_{2}$ | 6.4 | $174^{\circ}$ | 1.1 | $012^{\circ}$ | 36.7 | $251^{\circ}$ |
| $\mathrm{S}_{2}$ | 4.1 | $227^{\circ}$ | 1.0 | $040^{\circ}$ | 20.8 | 304 |

It is immediately apparent that the baroclinic component as computed is substantially larger than the barotropic. The amplitude of the barotropic component however, can be independently checked. Since the internal waves at no time are responsible for the horizontal transport of fluid, it is the barotropic mode which must drive the oscillating water level in the Inlet. The transport at any instant is equal to the barotropic tidal current integrated over the cross-sectional area of the Inlet entrance. This mass transport must be equal to the rate of change of the water level integrated over the surface area of Bridport Inlet and the mass transport over half a tidal cycle must be equal to the volume change of the Inlet, or

$$
\begin{equation*}
\mathrm{CS}_{0} \int^{\pi} \mathrm{X} \sin \check{\omega t}=2 \mathrm{H}_{0} \mathrm{~A} \tag{20}
\end{equation*}
$$

where $C S=$ the cross-sectional area of the Inlet entrance $=1.2 \times 10^{5} \mathrm{~m}^{2}$
$\mathrm{H}_{\mathrm{o}}=$ the tidal amplitude so that $2 \mathrm{H}_{\mathrm{o}}=$ tidal range
$A=$ the surface area of the inlet $=9.2 \times 10^{7} \mathrm{~m}^{2}$

The above computation reveals that a barotropic tidal current amplitude of $4.1 \mathrm{~cm} \mathrm{~s}^{-1}$ would be required to force the observed tidal range for the $M_{2}$ component.

The barotropic currents computed from the modal decomposition are therefore too small to explain the observed tidal range. In addition the phase difference between the barotropic currents and height should be approximately $\pi / 2$ close to that for a standing wave. The computed phase difference for the baroclinic components relative to the tidal heights are about $\pi / 2$ but the computed barotropic currents lead the tidal heights by about $3^{\pi} / 2$.

The conclusion must be that the modal decomposition is in some way inadequate to unify the observations. Phillips (1969) develops the theory of internal waves in the presence of a weak mean shear. In order for his development (and the methods of mode structure computation available) to apply, the shear must be much smaller than the Vaisald frequency or

$$
\frac{\partial u}{\partial z} \ll N
$$

where $\partial u / \partial z$ is the mean shear and $N$ is the Väisäla frequency $=\left(-\frac{g}{\rho} \frac{\partial p}{\partial z}\right)^{\frac{1}{2}}$. at the entrance to Bridport Inlet, the following approximate values apply

$$
\begin{aligned}
& N=4 \times 10^{-3} s^{-1} \\
& \partial u / \partial z=6 \times 10^{-3} s^{-1}
\end{aligned}
$$

Since $\partial u / \partial z$ is not much smaller than $N$, the observed oscillatory flow may be a manifestation of internal waves in a strong mean shear. This phenomenon may be of interest to fluid dynamicists working in this field, but the problem is not pursued further here.

The presence of a relatively strong mean shear can be represented in terms of an overall Richardson number, which is

$$
\mathrm{Ri}_{\mathrm{o}}=\frac{\mathrm{N}^{2} \mathrm{H}^{2}}{\mathrm{U}^{2}}
$$

where $R i_{o}$ is the overall Richardson number, $N$ is the Väisälä frequency, $H$ is the field depth and $U$ is the maximum velocity (such that $U H \simeq \partial u / \partial z$ ). For the entrance to Bridport Inlet, the overall Richardson number is 20. According to Turner (1973) the oscillating flow becomes critical when the overall Richardson number is less than $\pi^{2}$. This flow is therefore stable, but barely.

Maximum tidal currents at the entrance reach about $10 \mathrm{~cm} \mathrm{~s}^{-1}$ at spring tide while they are less than $2 \mathrm{~cm} \mathrm{~s}{ }^{-1}$ within the Inlet. The energy contained in the tidal oscillations is therefore relatively small in Bridport Inlet and it is not surprising that no events of instability such as lee waves of internal hydraulic jumps were apparent from the data set.

SOME COMMENTS ON THE EFFECTS OF AN LNG TERMINAL

From three years' of oceanographic data collection in Bridport Inlet the picture has emerged of an area characterized by relatively slow flows and dominated by convective processes due to brine rejection from growing sea ice. Currents rarely exceed $10 \mathrm{~cm} \mathrm{~s}{ }^{-1}$ at the entrance and are not discernable in the interior of the Inlet. The stratification is extremely weak, so much so that convection to the bottom could be triggered by just a few days' ice growth.

In the advent of the development of a natural gas liquifaction plant and an LNG tanker terminal at Bridport Inlet, both ship operations and the discharge of heated cooling water would contribute discernably to the energy balance of the Inlet. The heated cooling water is to be discharged in a confined region to minimize the ice thickness in the berthing area. Most of the heat derived from the liquifaction of the natural gas will therefore melt ice in a confined area and eventually be lost to the atmosphere. The motion of the ships themselves, however, will probably affect, the density structure within the Inlet.

By computing the potential energy difference between a fully mixed water column and that observed in 1980 and multiplying this difference by the Inlet surface area and it was found that $1.5 \times 10^{12}$ joules would have been required to completely mix the Inlet. This amount of energy could be added to the system by dissipating about 17 Mega Watts continously over a one day period.

The LNG carriers proposed are to have approximately 100 MW power plants. It is not unreasonable to assume that at least 20 MW would be required for maneuvering within the Inlet. 24 hours of such maneuvering would provide enough kinetic energy (through dissipation of the propellor wash) to fully mix Bridport Inlet.

It is the combination of an extremely weak stratification of the Inlet with the very large ship power plants which make this situation possible. Induced mixing would increase the apparent diffusive exchange of salt and decrease the advective exchange. However, the impact (if any) of the
altered mass structure on the biological community is out of the realm of expertise of this writer.

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APPENDIX A**
MEAN CTD PROFILES


CRUISE 15-80-020 ERIDPORT INLET-80
STATION 80-1 EXPERIMENT NO.S 3325-28,3387
LAT. 75-02.9 N LONG. 108-48.7 W
WATER DEPTH 120.0 M
pressure $T$ mean $T$ range $S$ mean $S$ range sigmat sigmat (DBARS) (DEG.C) (DEG.C) MEAN RANGE

| 2.0 | -1.800 | . 007 |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 3.0 | $-1.793$ | . 009 | 32.673 | . 027 | 26.289 | . 022 |
| 4.0 | -1.786 | . 008 | 32.682 | . 026 | 26.296 | . 021 |
| 5.0 | -1.784 | . 005 | 32.712 | .038 | 26.321 | . 031 |
| 6.0 | -1.790 | . 010 | 32.719 | . 048 | 26.326 | . 039 |
| 7.0 | -1.787 | .047 | 32.726 | . 053 | 26.332 | . 044 |
| 8.0 | -1.788 | . 043 | 32.733 | . 036 | 26.337 | . 030 |
| 9.0 | -1.791 | . 029 | 32.736 | . 027 | 26.340 | . 022 |
| 10.0 | $-1.794$ | . 021 | 32.740 | . 017 | 26.343 | . 014 |
| 11.0 | -1.795 | . 014 | 32.742 | . 013 | 26.345 | . 011 |
| 12.0 | -1.797 | . 009 | 32.744 | . 010 | 26.346 | . 008 |
| 13.0 | -1.797 | . 007 | 32.744 | . 008 | 26.346 | . 007 |
| 14.0 | $-1.798$ | . 005 | 32.745 | . 008 | 26.348 | . 007 |
| 15.0 | -1.797 | . 005 | 32.745 | . 007 | 26.347 | . 006 |
| 16.0 | -1.798 | . 007 | 32.746 | -D11 | 26.348 | . 009 |
| 17.0 | -1.798 | . 005 | 32.747 | . 009 | 26.349 | -007 |
| 18.0 | -1.798 | . 005 | 32.747 | . 009 | 26.349 | . 007 |
| 19.0 | -1.798 | . 004 | 32.748 | . 010 | 26.350 | . 009 |
| 20.0 | -1.798 | . 004 | 32.749 | . 007 | 26.351 | . 006 |
| 21.0 | -1.798 | . 007 | 32.749 | . 011 | 26.351 | . 009 |
| 22.0 | $-1.799$ | . 005 | 32.749 | . 009 | 26.351 | . 007 |
| 23.0 | -1.799 | . 004 | 32.750 | . 009 | 26.351 | . 007 |
| 24.0 | -1.799 | . 005 | 32.750 | . 006 | 26.352 | . 0005 |
| 25.0 | -1.798 | . 005 | 32.750 | . 008 | 26.351 | . 007 |
| 26.0 | -1.799 | . 005 | 32.751 | .007 | 26.352 | .006 |
| 27.0 | -1.799 | . 005 | 32.751 | . 009 | 26.352 | . 007 |
| 28.0 | -1.798 | . 005 | 32.749 | . 006 | 26.351 | . 005 |
| 29.0 | -1.799 | . 004 | 32.750 | . 006 | 26.351 | . 005 |
| 30.0 | -1.799 | . 002 | 32.750 | . 002 | 26.352 | . 002 |
| 31.0 | -1.799 | . 005 | 32.751 | . 007 | 26.353 | . 006 |
| 32.0 | -1.799 | . 003 | 32.751 | . 005 | 26.352 | . 004 |
| 33.0 | -1.799 | . 005 | 32.751 | - D06 | 26.352 | . 005 |
| 34.0 | -1.799 | . 001 | 32.751 | . 001 | 26.352 | . 001 |
| 35.0 | -1.799 | . 003 | 32.751 | . 003 | 26.353 | . 002 |
| 36.0 | -1.799 | .001 | 32.750 | . 001 | 26.352 | . 001 |
| 37.0 | -1.800 | . 003 | 32.752 | . 001 | 26.353 | . 001 |
| 38.0 | -1.799 | . 001 | 32.751 | . 001 | 26.353 | . 001 |
| 39.0 | -1.799 | . 003 | 32.752 | . 002 | 26.353 | .002 |
| 40.0 | -1.799 | . 002 | 32.752 | . 002 | 26.353 | . 001 |
| 41.0 | -1.800 | . 000 | 32.752 | . 001 | 26.354. | . 001 |
| 42.0 | -1.799 | . 005 | 32.752 | .002 | 26.353 | . 002 |
| 43.0 | -1.800 | . 003 | 32.753 | -005 | 26.354 | . 004 |
| 44.0 | -1.800 | . 002 | 32.753 | . 003 | 26.354. | . 002 |


| PRESSURE | T MEAN | t range | 5 MEAN | $S$ RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DBARS) | (DEG.C) | (DEG.C) |  |  | MEAN | range |
| 45.0 | -1.799 | . 004 | 32.753 | . 003 | 26.354 | . 002 |
| 46.0 | -1.799 | . 004 | 32.753 | . 003 | 26.354 | . 002 |
| 47.0 | -1.801 | . 002 | 32.754 | . 002 | 26.355 | . 002 |
| 48.0 | -1.800 | . 003 | 32.753 | . 004 | 26.354 | .003 |
| 49.0 | -1.800 | . 003 | 32.754 | . 003 | 26.355 | .002 |
| 50.0 | -1.799 | . 003 | 32.754 | . 003 | 26.355 | . 003 |
| 51.0 | -1.800 | . 002 | 32.754 | . 002 | 26.354 | . 002 |
| 52.0 | -1.799 | . 002 | 32.753 | .003 | 26.354 | . 002 |
| 53.0 | -1.794 | . 001 | 32.753 | . 002 | 26.354 | . 001 |
| 54.0 | -1.799 | . 003 | 32.753 | . 003 | 26.354 | . 002 |
| 55.0 | -1.799 | . 002 | 32.753 | . 005 | 26.354 | . 004 |
| 56.0 | -1.799 | . 002 | 32.753 | . 006 | 26.354 | . 005 |
| 57.0 | -1.800 | . 001 | 32.754 | . 003 | 26.355 | . 002 |
| 58.0 | -1.799 | . 002 | 32.752 | . 005 | 26.353 | . 004 |
| 59.0 | -1.799 | . 003 | 32.753 | . 004 | 26.354 | . 003 |
| 60.0 | -1.799 | .001 | 32.753 | . 006 | 26.354 | . 005 |
| 61.0 | -1.799 | . 001 | 32.753 | . 004 | 26.354 | . 003 |
| 62.0 | -1.799 | . 001 | 32.752 | . 004 | 26.353 | . 003 |
| 63.0 | -1.800 | . 000 | 32.754 | . 004 | 26.354 | . 003 |
| 64.0 | -1.799 | . 001 | 32.753 | . 006 | 26.354 | . 005 |
| 65.0 | -1.798 | . 003 | 32.753 | .003 | 26.354 | . 002 |
| 66.0 | -1.798 | . 002 | 32.753. | . 002 | 26.354 | . 002 |
| 67.0 | -1.798 | . 002 | 32.753 | . 004 | 26.354 | . 003 |
| 68.0 | -1.799 | . 002 | 32.754 | .005 | 26.354 | . 004 |
| 69.0 | -1.799 | . 002 | 32.754 | . 003 | 26.355 | . 002 |
| 70.0 | -1.799 | . 002 | 32.753 | . 004 | 26.354 | . 003 |
| 71.0 | -1.798 | . 001 | 32.753 | . 005 | 26.354 | . 004 |
| 72.0 | -1.799 | . 001 | 32.755 | . 004 | 26.356 | . 004 |
| 73.0 | -1.798 | . 002 | 32.754 | . 005 | 26.355 | . 004 |
| 74.0 | -1.799 | . 001 | 32.754 | . 003 | 26.355 | . 002 |
| 75.0 | -1.798 | . 001 | 32.755 | .004 | 26.356 | . 003 |
| 76.0 | -1.798 | .003 | 32.756 | . 002 | 26.356 | .002 |
| 77.0 | -1.798 | . 002 | 32.757 | . 0.01 | 26.357 | . 001 |
| 78.0 | -1.797 | . 002 | 32.756 | . 001 | 26.356 | . 001 |
| 79.0 | -1.798 | . 002 | 32.757 | . 002 | 26.357 | . 001 |
| 80.0 | -1.798 | . 002 | 32.758 | . 003 | 26.358 | . 003 |
| 81.0 | -1.798 | .001 | 32.757 | . 004 | 26.357 | . 003 |
| 82.0 | -1.798 | . 002 | 32.758 | . 003 | 26.358 | . 002 |
| 83.0 | -1.798 | . 002 | 32.759 | . 003 | 26.359 | . 002 |
| 84.0 | -1.799 | . 001 | 32.758 | . 001 | 26.358 | . 001 |
| 85.0 | -1.798 | . 001 | 32.758 | . 003 | 26.358 | . 002 |
| 86.0 | -1.798 | .002 | 32.758 | . 004 | 26.358 | . 003 |
| 87.0 | -1.798 | . 002 | 32.759 | . 000 | 26.358 | . 000 |
| 88.0 | $-1.798$ | . 002 | 32.759 | . 003 | 26.359. | .003 |
| 89.0 | -1.798 | . 001 | 32.759 | . 001 | 26.359 | . 001 |
| 90.0 | -1.798 | . 002 | 32.758 | .003 | 26.358 | . 002 |
| 91.0 | -1.798 | . 002 | 32.759 | . 002 | 26.359 | .002 |
| 92.0 | -1.798 | . 001 | 32.759 | . 001 | 26.359 | -001 |

$\left.\begin{array}{ccccccc}\text { PRESSURE } & \text { TMEAN } & \text { T RANGE } & S \text { MEAN } & \text { S RANGE } & \text { SIGMAT } & \text { SIGMAT } \\ \text { (DGARS) } & \text { (DEG.C) } \\ \text { (DEG.C) }\end{array}\right]$


CRUISE 15-80-020 BRIDPORT INLET-80
STATION 80-2 EXPERIMENT NO.S 3329-42,3386 LAT. 75-00.8 N LONG. 108.48.5 W WATER DEPTH 137.D M
PRESSURE T MEAN T RANGE S MEAN S RANGE SIGMAT SIGMAT
(DGARS) MEAN RANGE

| 2.0 | -1.789 | .045 |  |  |  |  |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 3.0 | -1.778 | .035 | 32.622 | .037 | 26.247 | .030 |
| 4.0 | -1.777 | .034 | 32.620 | .038 | 26.245 | .031 |
| 5.0 | -1.773 | .033 | 32.626 | .064 | 26.250 | .052 |
| 0.0 | -1.770 | .034 | 32.637 | .064 | 26.259 | .052 |
| 7.0 | -1.765 | .041 | 32.648 | .051 | 26.268 | .041 |
| 6.0 | -1.757 | .042 | 32.657 | .044 | 26.275 | .035 |
| 9.0 | -1.752 | .035 | 32.664 | .032 | 26.280 | .026 |
| 10.0 | -1.747 | .033 | 32.671 | .021 | 26.286 | .017 |
| 11.0 | -1.744 | .028 | 32.684 | .016 | 26.297 | .014 |
| 12.0 | -1.750 | .031 | 32.696 | .018 | 26.306 | .015 |
| 13.0 | -1.757 | .030 | 32.702 | .016 | 26.311 | .014 |
| 14.0 | -1.765 | .027 | 32.708 | .014 | 26.317 | .012 |
| 15.0 | -1.770 | .016 | 32.714 | .018 | 26.322 | .015 |
| 16.0 | -1.772 | .011 | 32.716 | .016 | 26.323 | .013 |
| 17.0 | -1.773 | .014 | 32.717 | .017 | 26.324 | .014 |
| 18.0 | -1.774 | .020 | 32.720 | .019 | 26.326 | .016 |
| 19.0 | -1.776 | .021 | 32.721 | .021 | 26.328 | .017 |
| 20.0 | -1.781 | .015 | 32.727 | .015 | 26.332 | .012 |
| 21.0 | -1.785 | .011 | 32.732 | .014 | 26.336 | .012 |
| 22.0 | -1.788 | .010 | 32.735 | .012 | 26.339 | .010 |
| 23.0 | -1.790 | .012 | 32.737 | .014 | 26.341 | .012 |
| 24.0 | -1.791 | .011 | 32.738 | .012 | 26.342 | .010 |
| 25.0 | -1.791 | .012 | 32.740 | .011 | 26.343 | .009 |
| 26.0 | -1.793 | .009 | 32.742 | .009 | 26.345 | .008 |
| 27.0 | -1.794 | .007 | 32.744 | .007 | 26.347 | .006 |
| 26.0 | -1.795 | .006 | 32.7 .45 | .007 | 26.348 | .006 |
| 29.0 | -1.795 | .006 | 32.746 | .007 | 26.348 | .006 |
| 30.0 | -1.796 | .004 | 32.747 | .005 | 26.349 | .004 |
| 31.0 | -1.796 | .003 | 32.747 | .006 | 26.349 | .004 |
| 32.0 | -1.796 | .003 | 32.748 | .004 | 26.350 | .003 |
| 33.0 | -1.797 | .003 | 32.748 | .006 | 26.350 | .005 |
| 34.0 | -1.797 | .002 | 32.748 | .005 | 26.350 | .004 |
| 35.0 | -1.796 | .003 | 32.748 | .005 | 26.350 | .004 |
| 36.0 | -1.797 | .003 | 32.749 | .005 | 26.351 | .004 |
| 37.0 | -1.797 | .003 | 32.750 | .005 | 26.351 | .004 |
| 38.0 | -1.797 | .002 | 32.750 | .005 | 26.351 | .004 |
| 39.0 | -1.797 | .003 | 32.751 | .005 | 26.352 | .004 |
| 40.0 | -1.797 | .002 | 32.750 | .005 | 26.352 | .004 |
| 41.0 | -1.797 | .003 | 32.751 | .003 | 26.352 | .002 |
| 42.0 | -1.798 | .002 | 32.752 | .003 | 26.353 | .002 |
| 43.0 | -1.798 | .002 | 32.752 | .005 | 26.353 | .004 |
| 44.0 | -1.798 | .004 | 32.752 | .005 | 26.353 | .004 |


| PRESSURE | t mean | $t$ range | 5 MEAN | S RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DEARS) | (DEG.C) | (DEG.C) |  |  | MEAN | range |
| 45.0 | -1.798 | . 002 | 32.752 | .003 | 26.353 | . 002 |
| 46.0 | -1.798 | .003 | 32.753 | . 005 | 26.354 | . 004 |
| 47.0 | -1.798 | . 002 | 32.753 | . 003 | 26.354 | . 002 |
| 48.0 | -1.798 | . 003 | 32.753 | . 006 | 26.354 | . 005 |
| 49.0 | -1.798 | . 003 | 32.754 | . 007 | 26.355 | . 006 |
| 50.0 | -1.798 | . 003 | 32.753 | . 006 | 26.354 | . 005 |
| 51.0 | -1.799 | . 003 | 32.755 | . 004 | 26.355 | . 003 |
| 52.0 | -1.799 | . 003 | 32.755 | . 006 | 26.355 | . 005 |
| 53.0 | -1.798 | . 003 | 32.754 | . 005 | 26.355 | . 004 |
| 54.0 | -1.799 | . 003. | 32.755 | . 006 | 26.356 | . 005 |
| 55.0 | -1.799 | . 002 | 32.756 | . 003 | 26.356 | . 002 |
| 56.0 | -1.799 | . 002 | 32.756 | . 005 | 26.356 | . 004 |
| 57.0 | -1.799 | . 002 | 32.756. | . 004 | 26.357 | . 004 |
| 58.0 | -1.799 | . 005 | 32.756 | .006 | 26.356 | .005 |
| 59.0 | -1.799 | . 002 | 32.756 | . 006 | 26.357 | . 005 |
| 60.0 | -1.799 | . 003 | 32.757 | . 006 | 26.357 | . 005 |
| 61.0 | -1.798 | . 002 | 32.756 | . 005 | 26.357 | . 004 |
| 62.0 | -1.799 | . 003 | 32.757 | . 004 | 26.358 | . 003 |
| 63.0 | -1.799 | . 003 | 32.757 | . 005 | 26.357 | . 004 |
| 64.0 | -1.798 | . 003 | 32.757 | . 003 | 26.357 | .003 |
| 65.0 | -1.799 | . 002 | 32.758 | . 005 | 26.358 | .004 |
| 66.0 | -1.799 | . 001 | 32.757 | . 004 | 26.358 | . 003 |
| 67.0 | $-1.799$ | . 005 | 32.758 | . 007 | 26.358 | . 006 |
| 68.0 | -1.798 | . 003 | 32.758 | . 005 | 26.358 | . 004 |
| 69.0 | -1.799 | . 003 | 32.759 | . 005 | 26.359 | . 004 |
| 70.0 | -1.798 | . 004 | 32.758 | . 006 | 26.358 | . 005 |
| 71.0 | -1.798 | . 003 | 32.758 | . 003 | 26.358 | . 003 |
| 72.0 | -1.798 | . 003 | 32.758 | . 005 | 26.358 | .004 |
| 73.0 | -1.798 | . 005 | 32.758 | . 005 | 26.358 | . 004 |
| 74.0 | -1.798 | . 003 | 32.758 | .005 | 26.358 | . 004 |
| 75.0 | -1.798 | . 002 | 32.758 | . 005 | 26.358 | . 004 |
| 76.0 | $-1.798$ | . 003 | 32.758 | .003 | 26.358 | . 003 |
| 77.0 | -1.798 | . 002 | 32.759 | .005 | 26.359 | .004 |
| 78.0 | -1.798 | . 004 | 32.758 | . 005 | 26.358 | . 004 |
| 79.0 | -1.798 | . 003 | 32.759 | . 005 | 26.359 | . 004 |
| 80.0 | -1.798 | . 004 | 32.759 | . 006 | 26.359 | . 005 |
| 81.0 | -1.798 | . 003 | 32.758 | . 005 | 26.358 | . 004 |
| 82.0 | -1.798 | . 003 | 32.759 | . 003 | 26.359 | . 003 |
| 83.0 | -1.798 | . 004 | 32.759 | . 005 | 26.359 | . 004 |
| 84.0 | -1.798 | . 002 | 32.759 | . 004 | 26.359 | . 004 |
| 85.0 | -1.798 | . 002 | 32.759 | . 004 | 26.359 | . 004 |
| 86.0 | -1.798 | . 004 | 32.759 | . 005 | 26.359 | . 004 |
| 87.0 | -1.799 | . 003 | 32.760 | . 004 | 26.359 | . 003 |
| 88.0 | -1.798 | . 002 | 32.759 | . 005 | 26.359 | . 004 |
| 89.0 | -1.798 | . 003 | 32.760 | . 005 | 26.359 | . 004 |
| 90.0 | -1.798 | . 002 | 32.759 | . 004 | 26.359 | . 003 |
| 91.0 | -1.798 | .004 | 32.760 | . 009 | 26.359 | .008 |
| 92.0 | -1.798 | . 005 | 32.760 | . 007 | 26.360 | . 006 |


| PRESSURE | T MEAN | T Range | S MFAN | S Range | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DBARS) | (DEG.C) | (DEG.C) |  |  | MEAN | RANGE |
| 93.0 | -1.798 | . 002 | 32.760 | . 005 | 26.360 | .004 |
| 94.0 | -1.798 | . 002 | 32.760 | . 004 | 26.360 | . 003 |
| 95.0 | -1.798 | . 002 | 32.760 | . 005 | 26.360 | . 004 |
| 96.0 | -1.798 | . 003 | 32.761 | .004 | 26.360 | . 003 |
| 97.0 | $-1.799$ | . 002 | 32.761 | . 005 | 26.361 | . 004 |
| 98.0 | -1.798 | . 003 | 32.761 | . 005 | 26.361 | . 004 |
| 99.0 | -1.798 | . 003 | 32.761 | . 003 | 26.360 | . 003 |
| 100.0 | -1.798 | . 003 | 32.761 | . 004 | 26.361 | . 004 |
| 101.0 | -1.798 | . 003 | 32.761 | . 005 | 26.361 | . 004 |
| 102.0 | -1.798 | . 002 | 32.762 | . 005 | 26.361 | . 004 |
| 103.0 | -1.798 | . 003 | 32.761 | . 005 | 26.360 | . 004 |
| 104.0 | -1.798 | . 002 | 32.762 | . 004 | 26.361 | . 003 |
| 105.0 | -1.798 | . 005 | 32.762 | . 005 | 26.361 | . 005 |
| 106.0 | -1.798 | .003 | 32.762 | . 005 | 26.361 | . 004 |
| 107.0 | -1.798 | .002 | 32.762 | . 004 | 26.362 | . 003 |
| 108.0 | -1.798 | . 002 | 32.762 | . 004 | 26.361 | . 003 |
| 109.0 | -1.798 | . 002 | 32.763 | . 004 | 26.362 | . 003 |
| 110.0 | -1.798 | . 002 | 32.763 | . 005 | 26.362 | . 004 |
| 111.0 | -1.798 | . 003 | 32.762 | . 004 | 26.362 | . 003 |
| 112.0 | -1.798 | . 001 | 32.763 | . 004 | 26.362 | . 004 |
| 113.0 | -1.798 | . 003 | 32.763 | . 005 | 26.362 | . 004 |
| 114.0 | -1.798 | . 002 | 32.764 | . 003 | 26.362 | . 002 |
| 115.0 | -1.798 | . 002 | 32.763 | . 005 | 26.362 | . 004 |
| 116.0 | -1.798 | .002 | 32.763 | . 005 | 26.362 | . 004 |
| 117.0 | -1.797 | .005 | 32.763 | . 005 | 26.362 | . 004 |
| 118.0 | -1.798 | . 002 | 32.764 | . 007 | 26.363 | . 005 |
| 119.0 | -1.798 | .003 | 32.763 | . 008 | 26.362 | . 006 |
| 120.0 | -1.798 | . 003 | 32.764 | .003 | 26.363 | . 003 |
| 121.0 | -1.798 | . 002 | 32.764 | . 005 | 26.363 | . 004 |
| 122.0 | -1.797 | . 003 | 32.764 | . 004 | 26.363 | . 003 |
| 123.0 | $-1.798$ | . 003 | 32.765 | . 002 | 26.363 | . 002 |
| 124.0 | -1.797 | . 002 | 32.764 | . 003 | 26.363 | . 002 |
| 125.0 | -1.798 | .002 | 32.765 | . 003 | 26.364 | . 003 |
| 126.0 | -1.798 | . 002 | 32.765 | . 003 | 26.364 | . 003 |
| 127.0 | -1.797 | .001 | 32.765 | . 003 | 26.364 | . 003 |
| 128.0 | -1.798 | . 003 | 32.766 | . 003 | 26.364 | -002 |
| 129.0 | -1.797 | . 003 | 32.765 | . 005 | 26.364 | .004 |
| 130.0 | -1.797 | . 002 | 32.765 | . 005 | 26.364 | . 004 |
| 131.0 | -1.797 | . 002 | 32.765 | . 004 | 26.364 | . 003 |
| 132.0 | -1.797 | .003 | 32.766 | . 004 | 26.364 | . 003 |



CRUISE 15-80-020 BRIDPORT INLET-80
STATION 80-3 EXPERIMENT NO.S 3343-73,3385
LAT. 74-59:7.N LONG. 108-50.0 W
WATER DEPTH 109.0 M

PRESSURE T MEAN T RANGE S MEAN S RANGE SIGMAT SIGMAT
(DBARS) (DEG.C) (DEG.C) MEAN RANGE

| 2.0 | -1.784 | . 043 |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 3.0 | -1.762 | . 051 | 32.609 | . 072 | 26.236 | -058 |
| 4.0 | -1.756 | .066 | 32.606 | . 071 | 26.234 | .057 |
| 5.0 | -1. 1.745 | . 058 | 32.609 | .047 | 26.236 | .037 |
| 6.0 | -1.727 | . 055 | 32.620 | . 044 | 26.244 | . 036 |
| 7.0 | -1.723 | . 060 | 32.631 | . 061 | 26.253 | . 050 |
| 8.0 | -1.733 | . 059 | 32.640 | . 032 | 26.260 | . 028 |
| 9.0 | -1.741 | . 051 | 32.645 | . 022 | 26.265 | . 018 |
| 10.0 | -1.738 | .036 | 32.649 | . 013 | 26.268 | .011 |
| 11.0 | $-1.737$ | . 037 | 32.653 | . 014 | 26.272 | .012 |
| 12.0 | -1.738 | .031 | 32.658 | . 017 | 26.275 | . 014 |
| 13.0 | $-1.742$ | . 020 | 32.662 | . 018 | 26.279 | . 015 |
| 14.0 | -1.744 | . 022 | 32.665 | .013 | 26.282 | -011 |
| 15.0 | -1.747 | .010 | 32.669 | . 016 | 26.284 | .013 |
| 16.0 | -1.750 | . 019 | 32.673 | . 017 | 26.288 | . 014 |
| 17.0 | -1.752 | . 019 | 32.678 | . 026 | 26.292 | . 022 |
| 18.0 | -1.755 | .017 | 32.684 | . 034 | 26.297 | . 028 |
| 19.0 | -1.759 | .020 | 32.691 | . 034 | 26.302 | . 028 |
| 20.0 | -1.763 | . 025 | 32.698 | . 039 | 26.309 | .032 |
| 21.0 | -1.766 | . 026 | 32.703 | . 041 | 26.312 | .034 |
| 22.0 | -1.770 | . 026 | 32.708 | . 038 | 26.317 | . 032 |
| 23.0 | -1.775 | . 023 | 32.714 | . 030 | 26.322 | .025 |
| 24.0 | -1.779 | . 022 | 32.721 | . 030 | 26.327 | . 025 |
| 25.0 | -1.783 | .019 | 32.727 | . 027 | 26.332 | . 022 |
| 26.0 | $-1.786$ | . 024 | 32.731 | .033 | 26.336 | . 027 |
| 27.0 | -1.790 | .019 | 32.737 | . 028 | 26.340 | . 023 |
| 28.0 | -1.792 | . 018 | 32.740 | . 031 | 26.344 | . 025 |
| 29.0 | -1.794 | . 015 | 32.745 | . 026 | 26.348 | . 021 |
| 30.0 | -1.796 | . 014 | 32.747 | - 025 | 26.349 | . 020 |
| 31.0 | $-1.797$ | . 011 | 32.750 | . 020 | 26.351 | .016 |
| 32.0 | -1.797 | . 008 | 32.751 | .015 | 26.352 | .012 |
| 33.0 | -1.798 | -008 | 32.752 | . 015 | 26.353 | . 013 |
| 34.0 | -1.798 | . 006 | 32.753 | . 011 | 26.354 | . 009 |
| 35.0 | -1.799 | . 004 | 32.754 | .011 | 26.355 | . 009 |
| 36.0 | -1.799 | . 005 | 32.754 | .011 | 26.355 | .009 |
| 37.0 | -1.799 | . 004 | 32.755 | .008 | 26.356 | . 007 |
| 38.0 | -1.799 | .003 | 32.756 | . 008 | 26.356 | . 007 |
| 39.0 | -1.799 | . 003 | 32.756 | .009 | 26.356 | .007 |
| 40.0 | -1.800 | .005 | 32.757 | .010 | 26.357 | .008 |
| 41.0 | -1.800 | . 004 | 32.756 | . 007 | 26.357 | . 006 |
| 42.0 | -1.800 | .004 | 32.757 | .008 | 26.357 | . 007 |
| 43.0 | -1.800 | . 005 | 32.757 | .010 | 26.358 | .008 |
| 44.0 | -1.800 | .004 | 32.758 | . 009 | 26.358 | . 008 |


| PRESSURE | MEAN | range | $S$ MEAN | 5 RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DBARS) | (DEG.C) | (DEG.C) |  |  | MEAN | Range |
| 45.0 | -1.799 | . 003 | 32.757 | . 012 | 26.357 | . 010 |
| 46.0 | -1.800 | . 003 | 32.758 | .006 | 26.358 | . 005 |
| 47.0 | -1.800 | . 003 | 32.758 | . 005 | 26.358 | . 004 |
| 48.0 | -1.799 | . 004 | 32.758 | . 007 | 26.358 | . 006 |
| 49.0 | -1.799 | . 003 | 32.759 | . 006 | 26.359 | . 005 |
| 50.0 | -1.799 | . 003 | 32.759 | .007 | 26.359 | .005 |
| 51.0 | -1.800 | . 003 | 32.759 | . 006 | 26.359 | . 005 |
| 52.0 | -1.800 | . 004 | 32.759 | . 006 | 26.359 | . 005 |
| 53.0 | -1.799 | . 005 | 32.759 | .009 | 26.359 | . 007 |
| 54.0 | -1.799 | . 003 | 32.759 | . 006 | 26.359 | .005 |
| 55.0 | -1.799 | . 003 | 32.760 | . 006 | 26.360 | .005 |
| 56.0 | -1.799 | . 004 | 32.760 | . 006 | 26.360 | . 005 |
| 57.0 | -1.799 | . 003 | 32.760 | . 006 | 26.360 | . 005 |
| 58.0 | $-1.799$ | . 003 | '32.760 | . 006 | 26.360 | . 005 |
| 59.0 | -1.799 | . 004 | 32.760 | . 007 | 26.360 | . 006 |
| 60.0 | -1.799 | . 004 | 32.761 | .007 | 26.360 | . 006 |
| 61.0 | -1.799 | . 004 | 32.760 | .005 | 26.360 | .004 |
| 62.0 | -1.799 | . 004 | 32.761 | . 006 | 26.361 | . 005 |
| 63.0 | -1.799 | . 002 | 32.761 | .005 | 26.360 | . 004 |
| 64.0 | $-1.799$ | . 003 | 32.762 | . 005 | 26.361 | .004 |
| 65.0 | -1.799 | . 004 | 32.761 | . 005 | 26.361 | . 004 |
| 66.0 | -1.799. | . 004 | 32.762 | . 008 | 26.361 | .007 |
| 67.0 | -1.799 | . 004 | 32.762 | .009 | 26.362 | . 007 |
| 68.0 | -1.799 | . 005 | 32.762 | . 008 | 26.361 | . 006 |
| 69.0 | -1.799 | . 004 | 32.762 | .007 | 26.361 | . 006 |
| 70.0 | -1.799 | . 003 | 32.762 | . 008 | 26.362 | . 006 |
| 71.0 | -1.799 | . 004 | 32.763 | . 006 | 26.362 | . 005 |
| 72.0 | $-1.799$ | . 003 | 32.763 | . 008 | 26.362 | . 006 |
| 73.0 | -1.799 | - 003 | 32.763 | .007 | 26.362 | . 006 |
| 74.0 | -1.799 | . 004 | 32.763 | . 006 | 26.362 | . 005 |
| 75.0 | -1.799 | .003 | 32.763 | . 006 | 26.362 | . 005 |
| 76.0 | -1.799 | . 004 | 32.763 | . 006 | 26.362 | . 005 |
| 77.0 | -1.799 | . 002 | 32.763 | .005 | 26.362 | . 004 |
| 78.0 | -1.799 | . 003 | 32.763 | . 007 | 26.362 | . 006 |
| 79.0 | -1.799 | . 004 | 32.763 | . 006 | 26.362 |  |
| 80.0 | -1.799 | . 003 | 32.763 | . 006 | 26.362 | . 005 |
| 81.0 | -1.799 | . 004 | 32.764 | . 006 | 26.363 | . 005 |
| 82.0 | -1.799 | . 004 | 32.764 | . 007 | 26.363 | . 006 |
| 83.0 | -1.799 | . 005 | 32.764 | .007 | 26.363 | . 006 |
| 84.0 | -1.799 | . 003 | 32.764 | . 009 | 26.363 | . 007 |
| 85.0 | -1.799 | . 003 | 32.764 | . 006 | 26.363 | . 005 |
| 86.0 | -1.799 | . 004 | 32.765 | . 007 | 26.363 | . 005 |
| 87.0 | -1.799 | . 004 | 32.765 | . 007 | 26.363 | . 006 |
| 88.0 | -1.799 | . 002 | 32.765 | .005 | 26.363 | . 004 |
| 89.0 | -1.799 | . 002 | 32.765 | . 004 | 26.364 | . 003 |
| 90.0 | -1.799 | . 004 | 32.765 | . 005 | 26.364 | . 004 |
| 91.0 | -1.799 | . 003 | 32.765 | . 003 | 26.363 | . 003 |
| 92.0 | -1.799 | . 003 | 32.765 | .003 | 26.364 | .003 |


| PRESSURE | TMEAN | T RANGE | S MEAN | S RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DGARS) | (DEG.C) | (UEG.C) |  |  | MEAN | RANGE |
| 93.0 | -1.799 | .004 | 32.765 | .005 | 26.364 | .004 |
| 94.0 | -1.800 | .005 | 32.765 | .005 | 26.364 | .004 |
| 95.0 | -1.799 | .005 | 32.765 | .005 | 26.364 | .004 |
| 96.0 | -1.799 | .004 | 32.765 | .005 | 26.364 | .004 |
| 97.0 | -1.799 | .003 | 32.765 | .005 | 26.364 | .004 |
| 98.0 | -1.800 | .005 | 32.765 | .006 | 26.364 | .005 |
| 99.0 | -1.800 | .004 | 32.766 | .006 | 26.364 | .005 |
| 100.0 | -1.799 | .003 | 32.766 | .006 | 26.364 | .005 |
| 101.0 | -1.799 | .004 | 32.766 | .005 | 26.364 | .004 |



CRUISE 15-80-020 BRIDPORT. INLET-80 STATION 8U-4 EXPERIMENT NO.S 3374-3384 LAT. 74-58.1 N LONG. 108-53.0 W WATER DEPTH 64.0 M

PRESSURE TMEAN T RANGE S MEAN S RANGE SIGMAT SIGMAT (DEARS) (DEG.C) (DEG.C) MEAN. RANGE

| 2.0 | -1.787 | .007 |  |  |  |  |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 3.0 | -1.787 | .007 | 32.500 | .015 | 26.148 | .012 |
| 4.0 | -1.786 | .016 | 32.503 | .054 | 26.151 | .044 |
| 5.0 | -1.769 | .083 | 32.511 | .062 | 26.157 | .049 |
| 6.0 | -1.729 | .095 | 32.544 | .073 | 26.183 | .057 |
| 7.0 | -1.702 | .091 | 32.571 | .042 | 26.204 | .032 |
| 8.0 | -1.691 | .063 | 32.591 | .045 | 26.220 | .035 |
| 9.0 | -1.693 | .032 | 32.609 | .043 | 26.234 | .035 |
| 10.0 | -1.699 | .033 | 32.621 | .030 | 26.244 | .025 |
| 11.0 | -1.708 | .038 | 32.631 | .026 | 26.253 | .022 |
| 12.0 | -1.721 | .031 | 32.641 | .014 | 26.262 | .012 |
| 13.0 | -1.734 | .032 | 32.648 | .010 | 26.268 | .009 |
| 14.0 | -1.739 | .024 | 32.652 | .010 | 26.270 | .008 |
| 15.0 | -1.741 | .017 | 32.654 | .011 | 26.272 | .009 |
| 16.0 | -1.746 | .018 | 32.658 | .012 | 26.276 | .010 |
| 17.0 | -1.748 | .010 | 32.661 | .012 | 26.278 | .010 |
| 18.0 | -1.748 | .010 | 32.664 | .014 | 26.280 | .011 |
| 19.0 | -1.747 | .011 | 32.667 | .017 | 26.283 | .014 |
| 20.0 | -1.746 | .013 | 32.670 | .015 | 26.285 | .012 |
| 21.0 | -1.747 | .014 | 32.674 | .023 | 26.289 | .019 |
| 22.0 | -1.755 | .031 | 32.687 | .034 | 26.299 | .028 |
| 23.0 | -1.760 | 001 | 32.697 | .025 | 26.307 | .021 |
| 24.0 | -1.762 | .026 | 32.706 | .017 | 26.315 | .014 |
| 25.0 | -1.763 | .028 | 32.712 | .014 | 26.320 | .012 |
| 26.0 | -1.769 | .032 | 32.720 | .014 | 26.326 | .012 |
| 27.0 | -1.780 | .027 | 32.729 | .018 | 26.334 | .015 |
| 28.0 | -1.787 | .023 | 32.737 | .023 | 26.340 | .019 |
| 29.0 | -1.792 | .012 | 32.743 | .017 | 26.346 | .014 |
| 30.0 | -1.794 | .009 | 32.747 | .015 | 26.349 | .012 |
| 31.0 | -1.796 | .007 | 32.750 | .012 | 26.351 | .010 |
| 32.0 | -1.797 | .005 | 32.752 | .007 | 26.353 | .006 |
| 33.0 | -1.797 | .004 | 32.752 | .008 | 26.353 | .007 |
| 34.0 | -1.798 | .004 | 32.753 | .005 | 26.354 | .004 |
| 35.0 | -1.798 | .004 | 32.754 | .006 | 26.355 | .005 |
| 36.0 | -1.799 | .004 | 32.756 | .004 | 26.356 | .004 |
| 37.0 | -1.799 | .003 | 32.756 | .005 | 26.357 | .004 |
| 38.0 | -1.799 | .004 | 32.756 | .005 | 26.357 | .004 |
| 39.0 | -1.799 | .003 | 32.758 | .005 | 26.358 | .005 |
| 40.0 | -1.800 | .003 | 32.758 | .006 | 26.358 | .005 |
| 41.0 | -1.799 | .002 | 32.758 | .004 | 26.358 | .003 |
| 42.0 | -1.800 | .002 | 32.760 | .005 | 26.359 | .004 |
| 43.0 | -1.799 | .002 | 32.759 | .003 | 26.359 | .003 |
| 44.0 | -1.800 | .004 | 32.760 | .005 | 26.360 | .004 |


| Pressure | t MEAN | T range | 5 MEAN | S RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DBARS) | (DEG.C) | (dEG.C) |  |  | MEAN | Range |
| 45.0 | -1.800 | . 003 | 32.760 | . 004 | 26.360 | . 003 |
| 46.0 | -1.800 | . 002 | 32.761 | .003 | 26.360 | .003 |
| 47.0 | -1.800 | . 002 | 32.761 | . 002 | 26.361. | . 002 |
| 48.0 | -1.800 | . 003 | 32.761 | . 003 | 26.361 | . 003 |
| 49.0 | -1.799 | . 002 | 32.761 | . 005 | 26.360 | . 004 |
| 50.0 | -1.800 | . 003 | 32.762 | . 003 | 26.361 | . 003 |
| 51.0 | $-1.800$ | . 002 | 32.762 | . 004 | 26.361 | . 003 |
| 52.0 | -1.800 | . $0033^{\circ}$ | 32.763 | . 005 | 26.362 | . 004 |
| 53.0 | -1.800 | . 003 | 32.762 | . 004 | 26.362 | . 003 |
| $54=0$ | -1.800 | . 002 | 32.763 | .003 | 26.362 | . 003 |
| 55.0 | -1.800 | . 002 | 32.763 | . 003 | 26.362 | . 002 |
| 56.0 | -1.800 | . 0004 | 32.764 | . 005 | 26.363 | . 004 |
| 57.0 | -1.800 | . 003 | 32.764 | . 004 | 26.363 | . 005 |
| 58.0 | -1.800 | . 003 | 32.764 | . 004 | 26.363 | . 003 |
| 59.0 | -1.800 | . 003 | 32.764 | . 002 | 26.363 | . 002 |
| 60.0 | -1.799. | . 002 | 32.763 | . 004 | 26.362 | .003 |
| 61.0 | -1.799 | . 002 | 32 | 003 | 26.362 | . 002 |

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| PRESSURE | T MEAN | TRANGE | SMEAN | S RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DGARS) | (DEG.C) | (DEG.C) |  |  | MEAN | RANGE |
| 45.0 | -1.794 | .001 | 32.737 | .003 | 26.341 | .002 |
| 46.0 | -1.794 | .001 | 32.738 | .004 | 26.342 | .003 |
| 47.0 | -1.794 | .001 | 32.738 | .001 | 26.342 | .001 |
| 48.0 | -1.794 | .001 | 32.738 | .001 | 26.342 | .001 |
| 49.0 | -1.794 | .000 | 32.739 | .002 | 26.342 | .002 |
| 50.0 | -1.794 | .000 | 32.740 | .000 | 26.343 | .000 |
| 51.0 | -1.794 | .002 | 32.740 | .003 | 26.344 | .002 |
| 52.0 | -1.794 | .001 | 32.741 | .001 | 26.344 | .001 |
| 53.0 | -1.794 | .001 | 32.741 | .001 | 26.344 | .000 |
| 54.0 | -1.794 | .002 | 32.743 | .001 | 26.346 | .001 |
| 55.0 | -1.794 | .000 | 32.744 | .000 | 26.347 | .000 |
| 56.0 | -1.794 | .002 | 32.743 | .002 | 26.346 | .002 |
| 57.0 | -1.793 | .000 | 32.743 | .000 | 26.346 | .000 |
| 58.0 | -1.794 | .003 | 32.743 | .003 | 26.346 | .003 |
| 59.0 | -1.794 | .002 | 32.744 | .004 | 26.347 | .003 |
| 60.0 | -1.794 | .001 | 32.745 | .002 | 26.348 | .002 |
| 61.0 | -1.794 | .000 | 32.745 | .000 | 26.347 | .000 |
| 62.0 | -1.795 | .000 | 32.746 | .002 | 26.348 | .001 |
| 63.0 | -1.794 | .001 | 32.744 | .002 | 26.347 | .002 |
| 64.0 | -1.794 | .001 | 32.745 | .002 | 26.347 | .002 |
| 65.0 | -1.794 | .001 | 32.747 | .001 | 26.349 | .001 |
| 66.0 | -1.794 | .002 | 32.746 | .002 | 26.349 | .002 |

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CRUTSE 15-80-020 BRIDPORT INLET-80 STATION CM-4 EXPERIMENT NO.S 3322-24,3391-93 LAT. 75-02.1 N LONG. 108-41.3 W WATER DEPTH 62.0 M

| PRESSURE | T MEAN | T RANGE | S MEAN | S | RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DBARS) | (DEG.C) | (DEG.C) |  |  |  | MEAN | RANGE |
| 2.0 | -1.813 | . 0.17 |  |  |  |  |  |
| 3.0 | -1.794 | .023 |  |  |  |  |  |
| 4.0 | -1.792 | . 018 | 32.704 |  | . 097 | 26.314 | . 080 |
| 5.0 | -1.792 | . 019 | 32.703 | " | . 097 | 26.313 | . 080 |
| 6.0 | -1.793 | .017 | 32.706 |  | . 089 | 26.316 | . 072 |
| 7.0 | -1.793 | . 014 | 32.710 |  | . 072 | 26.319 | . 059 |
| 8.0 | -1.793 | . 017 | 32.713 |  | .067 | 26.321 | . 055 |
| 9.0 | -1.793 | . 018 | 32.715 |  | .067 | 26.323 | .055 |
| 10 | $-1.792$ | . 019 | 32.717 |  | .061 | 26.325 | . 050 |
| 11.0 | -1.791 | .021 | 32.718 |  | . 059 | 26.326 | . 048 |
| 12.0 | -1.789 | .027 | 32.719 |  | . 055 | 26.326 | . 045 |
| 13.0 | $-1.789$ | . 024 | 32.719 |  | .053 | 26.326 | . 044 |
| 14.0 | -1.789 | . 026 | 32.720 |  | . 055 | 26.327 | . 045 |
| 15.0 | -1.788 | . 024 | 32.720 |  | .050 | 26.327 | . 042 |
| 16.0 | -1.788 | .024 | 32.721 |  | . 048 | 26.328 | . 040 |
| 17.0 | $-1.791$ | .019 | 32.724 |  | .043 | 26.330 | . 035 |
| 18.0 | -1.792 | . 01.7 | 32.726 |  | . 040 | 26.332 | .033 |
| 19.0 | -1.792 | . 017 | 32.728 |  | . 036 | 26.334 | .030 |
| 20.0 | $-1.793$ | . 016 | 32.730 |  | . 033 | 26.335 | . 027 |
| 21.0 | -1.792 | .014 | 32.728 |  | . 034 | 26.334 | . 028 |
| 22.0 | -1.794 | .014 | 32.733 |  | .033 | 26.338 | . 027 |
| 23.0 | -1.793 | . 014 | 32.731 |  | . 031 | 26.336 | . 026 |
| 24.0 | -1.793 | .013 | 32.733 |  | . 028 | 26.338 | .023 |
| 25.0 | -1.7.94 | .013 | 32.735 |  | .023 | 26.340 | .019 |
| 26.0 | $-1.793$ | . 013 | $32.736^{\prime}$ |  | .022 | 26.340 | .018 |
| 27.0 | -1.794 | . 011 | 32.739 |  | .017 | 26.342 | . 014 |
| 28.0 | $-1.795$ | . 011 | . 32.740 |  | .015 | 26.343 | .013 |
| 29.0 | -1.796 | .007 | 32.742 |  | .011 | 26.345 | . 009 |
| 30.0 | -1.796 | .007 | 32.743 |  | .007 | 26.346 | .005 |
| 31.0 | $-1.796$ | .006 | 32.743 |  | .007 | 26.346 | . 006 |
| 32.0 | -1.797 | . 007 | 32.745 |  | .009 | 26.347 | .007 |
| 33.0 | -1.798 | .003 | 32.745 |  | .003 | 26.348 | -002 |
| 34.0 | -1.7.9.9 | .003 | 32.746 |  | .006 | 26.348 | . 005 |
| 35.0 | -1.798 | .003 | 32.746 |  | .004 | 26.348 | .003 |
| 36.0 | -1.799 | .001 | 32.745 |  | .002 | 26.348 | . 001 |
| 37.0 | -1.798 | .003 | 32.745 |  | .006 | 26.347 | . 005 |
| 38.0 | -1.798 | . 002 | 32.745 |  | .002 | 26.348 | .002 |
| 39.0 | $-1.798$ | -001 | 32.746 |  | .003 | 26.348 | .002 |
| 40.0 | $-1.798$ | . 003 | 32.746 |  | .006 | 26.348 | .005 |
| 41.0 | -1.799 | -003 | 32.7 .47 |  | . 004 | 26.349 | .003 |
| 42.0 | -1.799 | .001 | 32.747 |  | .002 | 26.349 | . 001 |
| 43.0 | $-1.799$ | . 002 | 32.747 |  | . 004 | 26.349 | .003 |
| 44.0 | -1.799 | . 002 | 32.747 |  | . 004 | 26.349 | .003 |


| PRESSURE | 1 MEAN | 1 Range | 5 MEAN | S RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DBARS) | (DEG.C) | (DEG.C) |  |  | MEAN | RANGE |
| 45.0 | -1.799 | . 002 | 32.747 | .001 | 26.349 | .001 |
| 46.0 | -1.799 | . 001 | 32.748 | .003 | 26.350 | . 002 |
| 47.0 | -1.799 | . 002 | 32.748 | . 005 | 26.350 | . 004 |
| 48.0 | -1.799 | - 002 | 32.748 | . 005 | 26.350 | . 004 |
| 49.0 | -1.799 | .003 | 32.747 | .004 | 26.349 | . 003 |
| 50.0 | -1.799 | . 001 | 32.749 | . 002 | 26.351 | . 002 |
| 51.0 | -1.799 | .002 | 32.750 | .005 | 26.351 | . 004 |
| 52.0 | -1.799 | .001 | 32.749 | . 001 | 26.351 | .001 |
| 53.0 | -1.799 | .002 | 32.750 | . 005 | 26.351 | .004 |
| 54.0 | -1.799 | .002 | 32.750 | .002 | 26.351 | . 001 |
| 55.0 | -1.799 | .002 | 32.749 | .006 | 26.351 | .005 |
| 56.0 | -1.799 | .002 | 32.751 | .003 | 26.352 | .003 |
| 57.0 | -1.799 | .003 | 32.750 | .003 | 26.352 | . 002 |
| 58.0 | -1.798 | .003 | 32.750 | . 002 | 26.352 | . 002 |
| 59.0 | $-1.799$ | .003 | 32.750 | .002 | 26.352 | -002 |

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CRUISE 15-80-020 ERIDPORT INLET-80
STATION CM-5 EXPERIMENT NO.S 3317-21,3394-97
LAT. 75-00.3 N LONG. 108-42.3 W
WATER DEPTH 28.D M
```

| PRESSURE | TMEAN | T RANGE | SMEAN | S RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DBARS) | (DEG.C) | IDEG.C) |  |  | MEAN | RANGE |
| 2.0 | -1.801 | .020 |  |  |  |  |
| 3.0 | -1.787 | .018 | 32.681 | .031 | 26.295 | .026 |
| 4.0 | -1.785 | .018 | 32.670 | .034 | 26.286 | .028 |
| 5.0 | -1.786 | .019 | 32.670 | .037 | 26.286 | .030 |
| 6.0 | -1.785 | .017 | 32.670 | .035 | 26.286 | .028 |
| 7.0 | -1.784 | .017 | 32.671 | .040 | 26.287 | .032 |
| 8.0 | -1.784 | .017 | 32.677 | .058 | 26.292 | .048 |
| 9.0 | -1.785 | .017 | 32.683 | .063 | 26.297 | .051 |
| 10.0 | -1.784 | .016 | 32.684 | .065 | 26.298 | .053 |
| 11.0 | -1.784 | .018 | 32.684 | .069 | 26.297 | .056 |
| 12.0 | -1.783 | .020 | 32.685 | .070 | 26.298 | .057 |
| 13.0 | -1.784 | .018 | 32.686 | .070 | 26.299 | .057 |
| 14.0 | -1.783 | .019 | 32.686 | .070 | 26.299 | .057 |
| 15.0 | -1.785 | .019 | 32.687 | .071 | 26.300 | .058 |
| 16.0 | -1.785 | .017 | 32.688 | .071 | 26.301 | .058 |
| 17.0 | -1.785 | .021 | 32.689 | .073 | 26.302 | .060 |
| 18.0 | -1.784 | .025 | 32.690 | .076 | 26.302 | .062 |
| 19.0 | -1.783 | .031 | 32.691 | .075 | 26.303 | .062 |
| 20.0 | -1.783 | .031 | 32.692 | .071 | 26.304 | .058 |
| 21.0 | -1.785 | .028 | 32.695 | .071 | 26.306 | .058 |
| 22.0 | -1.784 | .029 | 32.698 | .070 | 26.309 | .058 |
| 23.0 | -1.784 | .031 | 32.703 | .064 | 26.313 | .052 |
| 24.0 | -1.784 | .034 | 32.706 | .061 | 26.315 | .050 |
| 25.0 | -1.788 | .030 | 32.712 | .053 | 26.320 | .043 |



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CRUISE 15-80-020 BRIDPORT INLET-80
STATION CM-G EXPERIMENT NO.S 3312-15,3398-3400
LAT. 74-59.6 N LONG. 108-49.3 W
WATER DEPTH 26.0 M
```

| PRESSURE | TMEAN | TRANGE | $S$ MEAN | S RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DBARS) | (DEG.C) | (DEG.C) |  |  | MEAN | RANGE |
| 3.0 | -1.782 | .008 | 32.627 | .026 | 26.251 | .021 |
| 4.0 | -1.779 | .008 | 32.630 | .017 | 26.254 | .014 |
| 5.0 | -1.779 | .009 | 32.635 | .010 | 26.258 | .008 |
| 6.0 | -1.776 | .017 | 32.636 | .005 | 26.258 | .004 |
| 7.0 | -1.773 | .019 | 32.637 | .009 | 26.260 | .007 |
| 8.0 | -1.774 | .021 | 32.640 | .009 | 26.261 | .007 |
| 9.0 | -1.772 | .025 | 32.642 | .018 | 26.263 | .014 |
| 10.0 | -1.772 | .026 | 32.644 | .023 | 26.265 | .019 |
| 11.0 | -1.773 | .030 | 32.654 | .052 | 26.273 | .042 |
| 12.0 | -1.776 | .020 | 32.657 | .061 | 26.275 | .049 |
| 13.0 | -1.778 | .011 | 32.661 | .060 | 26.279 | .048 |
| 14.0 | -1.779 | .009 | 32.666 | .071 | 26.283 | .058 |
| 15.0 | -1.777 | .015 | 32.669 | .079 | 26.286 | .064 |
| 16.0 | -1.776 | .032 | 32.681 | .090 | 26.295 | .074 |
| 17.0 | -1.774 | .034 | 32.685 | .082 | 26.299 | .067 |
| 18.0 | -1.775 | .033 | 32.691 | .075 | 26.303 | .062 |
| 19.0 | -1.778 | .034 | 32.697 | .066 | 26.308 | .054 |
| 20.0 | -1.780 | .033 | 32.702 | .060 | 26.312 | .050 |



CRUISE 15-80-020 BRIDPORT INLET-80 STATION CM-7 EXPERIMENT NO.S 3309-11.3401-07 LAT. 74-59.7 N LONG. 108-50.0 W WATER DEPTH 109.D M

| PRESSURE | $T$ MEAN | T RANGE | $S$ MEAN | S | RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DBARS) | (DEG.C) | (DEG.C) |  |  |  | MEAN | RANGE |
| 2.0 | -1.783 | . 028 |  |  |  |  |  |
| 3.0 | -1.775 | -025 | 32.618 |  | . 025 | 26.244 | . 021 |
| 4.0 | -1.774 | . 025 | 32.616 |  | . 021 | 26.242 | . 018 |
| 5.0 | $-1.773$ | .022 | 32.620 | + | . 022 | 26.245 | .018 |
| 0.0 | -1.775 | . 019 | 32.626 |  | . 031 | 26.250 | . 025 |
| 7.0 | $-1.780$ | .006 | 32.637 |  | . 024 | 26.259 | .020 |
| 8.0 | -1.781 | .005 | 32.640 |  | . 029 | 26.262 | .024 |
| 9.0 | -1.780 | . 012 | 32.641 |  | . 031 | 26.263 | . 025 |
| 10.0 | -1.781 | .010 | 32.644 |  | . 039 | 26.265 | . 031 |
| 11.0 | $-1.780$ | . 013 | 32.645 |  | . 042 | 26.266 | .034 |
| 12.0 | -1.778 | .016 | 32.646 |  | . 042 | 26.267 | . 034 |
| 13.0 | -1.776 | . 022 | 32.648 |  | . 043 | 26.268 | . 035 |
| 14.0 | -1.775 | . 023 | 32.649 |  | . 044 | 26.269 | . 035 |
| 15.0 | -1.770 | . 023 | 32.652 |  | .044 | 26.271 | . 036 |
| 16.0 | -1.763 | . 027 | 32.656 |  | . 038 | 26.275 | . 031 |
| 17.0 | -1.759 | . 022 | 32.663 |  | . 037 | 26.280 | .030 |
| 18.0 | -1.758 | . 017 | 32.668 |  | . 041 | 26.284 | . 033 |
| 19.0 | $-1.762$ | . 020 | 32.676 |  | . 037 | 26.291 | . 031 |
| 20.0 | -1.765 | . 019 | 32.683 |  | . 035 | 26.296 | .028 |
| 21.0 | -1.771 | . 020 | 32.693 |  | . 033 | 26.305 | .027 |
| 22.0 | -1.778 | . 019 | 32.702 |  | . 027 | 26.312 | . 024 |
| 23.0 | -1.781 | . 018 | 32.705 |  | . 025 | 26.314 | . 021 |
| 24.0 | $-1.787$ | . 012 | 32.713 |  | . 022 | 26.321 | . 018 |
| 25.0 | -1.792 | - 005 | 32.720 |  | . 016 | 26.327 | . 013 |
| 26.0 | -1.792 | . 004 | 32.725 |  | .014 | 26.331 | . 012 |
| 27.0 | -1.792 | -008 | 32.727 |  | . 012 | 26.333 | . 009 |
| 28.0 | -1.791 | - 008 | 32.728 |  | . 009 | 26.333 | . 008 |
| 29.0 | -1.792 | . 009 | 32:730 |  | . 011 | 26.335 | . 009 |
| 30.0 | $-1.792$ | . 007 | 32.730 |  | . 010 | 26.335 | . 008 |
| 31.0 | -1.793 | . 008 | 32.730 |  | . 007 | 26.335 | - 006 |
| 32.0 | -1.793 | -009 | 32.731 |  | . 010 | 26.336 | -008 |
| 33.0 | -1.792 | . 010 | 32.731 |  | . 010 | 26.336 | . 008 |
| 34.0 | -1.792 | - 011 | 32.732 |  | .011 | 26.337 | -009 |
| 35.0 | -1.792 | . 012 | 32.733 |  | .010 | 26.338 | -008 |
| 30.0 | -1.792 | - 010 | 32.734 |  | .007 | 26.338 | . 006 |
| 37.0 | -1.792 | . 009 | 32.733 |  | -004 | 26.338 | . 003 |
| 38.0 | $-1.793$ | $\therefore .010$ | 32.736 |  | . 006 | 26.340 | . 005 |
| 39.0 | -1.793 | -009 | 32.736 |  | . 005 | 26.340 | .005 |
| 40.0 | $-1.793$ | . 010 | 32.737 |  | -005 | 26.341 | . 004 |
| 41.0 | $-1.794$ | -009 | 32.737 |  | . 006 | 26.341 | - 005 |
| 42.0 | -1.794 | - 008 | 32.738 |  | . 008 | 26.341 | . 006 |
| 43.0 | -1.795 | - 007 | 32.739 |  | -009 | 26.343 | .007 |
| 44.0 | -1.795 | . 008 | 32.740 |  | . 008 | 26.343 | . 007 |


| PRESSURE | T MEAN | 1 range | 5 MEAN | $s$ range | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DEARS) | (DEG.C) | (DEG.C) |  |  | MEAN | Range |
| 45.0 | -1.793 | . 002 | 32.742 | . 009 | 26.345 | . 007 |
| 46.0 | -1.796 | . 006 | 32.742 | .009 | 26.345 | . 008 |
| 47.0 | -1.796 | . 006 | 32.745 | . 008 | 26.348 | . 006 |
| 48.0 | $-1.796$ | . 007 | 32.744 | . 008 | 26.347 | . 006 |
| 49.0 | -1.796 | . 005 | 32.746 | . 011 | 26.348 | . 009 |
| 50.0 | -1.797 | . 006 | 32.746 | . 008 | 26.348 | . 006 |
| 51.0 | -1.797 | . 008 | 32.747 | . 009 | 26.349 | . 008 |
| 52.0 | -1.797 | . 006 | 32.748 | . 012 | 26.350 | . 010 |
| 53.0 | -1.797 | . 005 | 32.747 | . 013 | 26.349 | .010 |
| 54.0 | -1.797 | . 006 | 32.749 | . 010 | 26.351 | . 008 |
| 55.0 | $-1.798$ | . 004 | 32.750 | . 009 | 26.352 | . 008 |
| 56.0 | -1.797 | . 007 | 32.751 | . 009 | 26.352 | . 007 |
| 57.0 | -1.798 | .003 | 32.752 | . 008 | 26.353 | . 006 |
| 58.0 | -1.798 | .006 | 32.753 | . 011 | 26.354 | . 008 |
| 59.0 | -1.798 | . 005 | 32.753 | . 008 | 26.354 | . 006 |
| 60.0 | -1.798 | . 005 | . 32.754 | . 009 | 26.355 | . 007 |
| 61.0 | -1.798 | . 004 | 32.754 | . 013 | 26.355 | . 010 |
| 62.0 | -1.798 | . 005 | 32.755 | . 010 | 26.355 | . 008 |
| 63.0 | $-1.798$ | . 005 | 32.755 | .011 | 26.356 | . 009 |
| 64.0 | -1.799 | . 005 | 32.756 | . 011 | 26.356 | . 009 |
| 65.0 | -1.798 | . 003 | 32.756 | . 012 | 26.356 | . 009 |
| 66.0 | -1.798 | . 004 | 32.756 | . 011 | 26.356 | . 009 |
| 67.0 | -1.798 | . 005 | 32.757 | . 010 | 26.357 | . 008 |
| 68.0 | -1.799 | . 005 | 32.757 | .010 | 26.357. | . 008 |
| 69.0 | -1.799 | . 005 | 32.757 | . 010 | 26.357 | . 008 |
| 70.0 | -1.798 | . 005 | 32.756 | . 010 | 26.357 | . 008 |
| 71.0 | $-1.798$ | . 006 | 32.757 | . 012 | 26.357 | . 009 |
| 72.0 | -1.799 | . 005 | 32.758 | . 009 | 26.358 | . 008 |
| 73.0 | $-1.799$ | . 004 | 32.758 | . 012 | 26.358 | . 010 |
| 74.0 | -1.799 | . 005 | 32.758 | . 010 | 26.358 | . 008 |
| 75.0 | -1.798 | . 005 | 32.758 | . 111 | 26.358 | -009 |
| 76.0 | -1.799 | . 006 | 32.759 | . 011 | 26.359 | .009 |
| 77.0 | $-1.798$ | .007 | 32.758 | . 013 | 26.358 | . 011 |
| 78.0 | -1.798 | . 005 | 32.759 | .012 | 26.359 | . 009 |
| 79.0 | -1.798 | . 005 | 32.759 | .013 | 26.359 | . 011 |
| 80.0 | -1.798 | .005 | 32.759 | . 012 | 26.358 | .010 |
| 81.0 | $-1.798$ | . 005 | 32.759 | .013 | 26.359 | . 010 |
| 82.0 | $-1.799$ | . 005 | 32.760 | . 011 | 26.359 | . 009 |
| 83.0 | -1.798 | . 006 | 32.759 | . 009 | 26.359 | . 008 |
| 84.0 | -1.798 | . 004 | 32.759 | . 013 | 26.359. | . 011 |
| 85.0 | -1.798 | . 004 | 32.760 | . 012 | $26.359^{\circ}$ | . 009 |
| 86.0 | -1.798 | . 004 | 32.760 | . 015 | 26.360 | . 012 |
| 87.0 | -1.798 | . 005 | 32.760 | . 013 | 26.359 | . 010 |
| 88.0 | $-1.798$ | . 006 | 32.760 | . 013 | 26.359 | . 011 |
| 89.0 | -1.798 | .006 | 32.760 | . 011 | 26.360 | . 009 |
| 90.0 | -1.798 | . 005 | 32.760 | . 010 | 26.359 | . 008 |
| 91.0 | $-1.798$ | . 006 | 32.761 | . 012 | 26.360 | . 010 |
| 92.0 | $-1.798$ | . 006 | 32.760 | . 012 | 26.360 | -009 |


| PRESSURE | TMEAN | T RANGE | $S$ MEAN | S RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DBARS) | (DEG.C) | IDEG.C) |  |  | MEAN | RANGE |
| 93.0 | -1.798 | .005 | 32.761 | .010 | 26.360 | .008 |
| 94.0 | -1.798 | .006 | 32.762 | .012 | 26.361 | .010 |
| 95.0 | -1.798 | .007 | 32.761 | .010 | 26.360 | .008 |
| 96.0 | -1.798 | .006 | 32.761 | .011 | 26.361 | .009 |
| 97.0 | -1.798 | .006 | 32.761 | .014 | 26.360 | .011 |
| 98.0 | -1.798 | .007 | 32.761 | .010 | 26.360 | .008 |
| 99.0 | -1.797 | .007 | 32.761 | .011 | 26.360 | .009 |
| 100.0 | -1.798 | .007 | 32.761 | .009 | 26.360 | .007 |
| 101.0 | -1.797 | .007 | 32.762 | .009 | 26.361 | .007 |



CRUISE 15-8U-O20 BRIDPORT INLET-80
STATION CM-8 EXPERIMENT NO.S 3306-08, 3408-10 LAT. 74-59.7 N LONG. 108-50.9 W WATER DEPTH 80.0 M

| PRESSURE | T MEAN | T RANGE | 5 MEAN | $S$ RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DGARS) | (DEG.C) | (DEG.C) |  |  | MEAN | RANGE |
| 3.0 | -1.781 | . 029 | 32.618 | . 041 | 26.244 | .033 |
| 4.0 | -1.777 | .020 | 32.623 | . 058 | 26.248 | . 047 |
| 5.0 | -1.780 | . 023 | 32.630 | . 064 | 26.253 | . 052 |
| 6.0 | -1.779 | . 039 | 32.635 | . 058 | 26.258 | . 04.7 |
| 7.0 | -1.777 | .040 | 32.642 | . 051 | 26.263 | . 041 |
| 8.0 | -1.777. | . 034 | 32.655 | .055 | 26.274 | . 044 |
| 9.0 | -1.778 | . 025 | 32.665 | . 037 | 26.282 | . 030 |
| 10.0 | -1.773 | .022 | 32.672 | . 034 | 26.288 | -027 |
| 11.0 | -1.774 | . 015 | 32.678 | . 037 | 26.292 | . 030 |
| 12.0 | -1.777 | . 010 | 32.683 | .038 | 26.297 | . 031 |
| 13.0 | -1.779 | . 011 | 32.690 | . 036 | 26.302 | . 029 |
| 14.0 | -1.780 | . 012 | 32.694 | . 027 | 26.306 | . 022 |
| 15.0 | -1.782 | . 015 | 32.703 | .020 | 26.313 | . 016 |
| 16.0 | -1.785 | . 015 | 32.709 | .006 | 26.318 | . 005 |
| 17.0 | -1.785 | . 014 | 32.711 | -005 | 26.319 | . 005 |
| 18.0 | -1.785 | . 015 | 32.712 | . 003 | 26.321 | . 002 |
| 19.0 | -1.786 | .014 | 32.716 | . 004 | 26.324 | .003 |
| 20.0 | -1.787 | .010 | 32.718 | . 008 | 26.325 | . 006 |
| 21.0 | -1.787 | . 007 | 32.720 | . 009 | 26.327 | . 007 |
| 22.0 | -1.788 | .008 | 32.722 | .010 | 26.328 | . 009 |
| 23.0 | -1.789 | . 007 | 32.723 | -009 | 26.329 | . 008 |
| 24.0 | -1.791 | . 006 | 32.724 | . 011 | 26.330 | . 009 |
| 25.0 | -1.791 | . 006 | 32.724 | . 010 | 26.330 | . 008 |
| 26.0 | -1.792 | - 007 | 32.725 | . 011 | 26.331 | . 009 |
| 27.0 | -1.791 | . 008 | 32.725 | . 012 | 26.331 | - 010 |
| 28.0 | -1.792 | . 008 | 32.726 | . 012 | 26.332 | .010 |
| 29.0 | .-1.793 | -007 | 32.727 | . 011 | 26.333 | . 009 |
| 30.0 | -1.793 | .007 | 32.728 | . 013 | 26.333 | . 010 |
| 31.0 | -1.794 | . 007 | 32.729 | - 013 | 26.334 | . 010 |
| 32.0 | -1.79,4 | . 010 | 32.730 | . 015 | 26.335 | .013 |
| 33.0 | -1.796 | -008 | 32.730 | . 014 | 26.335 | . 012 |
| 34.0 | -1.795 | . 009 | 32.729 | .015 | 26.335 | .013 |
| 35.0 | -1.795 | . 007 | 32.729 | . 010 | 26.334 | . 009 |
| 36.0 | -1.795 | . 010 | 32.730 | . 016 | 26.335 | .013 |
| 37.0 | $-1.796$ | .009 | 32.732 | .014 | 26.337 | . 012 |
| 38.0 | -1.795 | . 008 | 32.731 | . 015 | 26.336 | . 012 |
| 39.0 | -1.797 | . 010 | 32.732 | . 014 | 26.337 | . 012 |
| 40.0 | -1.796 | .010 | 32.731 | .015 | 26.336 | . 013 |
| 41.0 | -1.795 | .007 | 32.726 | .001 | 26.332 | .001 |
| 42.0 | -1.796 | .010 | 32.732 | . 018 | 26.337 | . 015 |
| 43.0 | -1.795 | . 0.11 | 32.732 | . 015 | 26.337 | . 013 |
| 44.0 | -1.798 | . 010 | 32.734 | .018 | 26.339 | .015 |
| 45.0 | -1.796 | .011 | 32.734 | .020 | 26.338 | . 016 |


|  |  |  |  |  |  |  |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| PRESSURE | TMEAN | T RANGE | S MEAN | S RANGE | SIGMAT | SIGMAT |
| (OEARS) | (DEG.C) | IDEG.C) |  |  | MEAN | RANGE |
| 46.0 | -1.795 | .011 | 32.730 | .018 | 26.335 | .015 |
| 47.0 | -1.794 | .010 | 32.732 | .017 | 26.337 | .014 |
| 48.0 | -1.798 | .009 | 32.738 | .018 | 26.342 | .015 |
| 49.0 | -1.793 | .010 | 32.733 | .016 | 26.337 | .013 |
| 50.0 | -1.795 | .011 | 32.736 | .019 | 26.340 | .016 |
| 51.0 | -1.796 | .008 | 32.738 | .018 | 26.342 | .015 |
| 52.0 | -1.796 | .010 | 32.738 | .018 | 26.341 | .015 |
| 53.0 | -1.795 | .009 | 32.738 | .017 | 26.341 | .014 |
| 54.0 | -1.797 | .008 | 32.740 | .016 | 26.333 | .013 |
| 55.0 | -1.797 | .006 | 32.742 | .013 | 26.345 | .010 |
| 56.0 | -1.797 | .008 | 32.742 | .013 | 26.345 | .011 |
| 57.0 | -1.797 | .006 | 32.742 | .010 | 26.345 | .009 |
| 58.0 | -1.797 | .007 | 32.742 | .011 | 26.345 | .009 |
| 59.0 | -1.797 | .006 | 32.743 | .010 | 26.346 | .009 |
| 60.0 | -1.797 | .005 | 32.745 | .007 | 26.347 | .006 |
| 61.0 | -1.799 | .004 | 32.747 | .007 | 26.349 | .006 |
| 62.0 | -1.799 | .003 | 32.749 | .004 | 26.351 | .003 |
| 63.0 | -1.799 | .004 | 32.750 | .004 | 26.352 | .003 |
| 64.0 | -1.800 | .003 | 32.750 | .002 | 26.352 | .002 |
| 65.0 | -1.800 | .002 | 32.752 | .004 | 26.353 | .003 |
| 66.0 | -1.800 | .004 | 32.753 | .002 | 26.354 | .001 |
| 67.0 | -1.799 | .002 | 32.752 | .006 | 26.353 | .005 |
| 68.0 | -1.800 | .003 | 32.753 | .005 | 26.354 | .004 |
| 69.0 | -1.800 | .002 | 32.753 | .006 | 26.354 | .005 |
| 70.0 | -1.799 | .001 | 32.753 | .008 | 26.354 | .006 |
| 71.0 | -1.799 | .002 | 32.753 | .009 | 26.354 | .007 |
| 72.0 | -1.800 | .003 | 32.753 | .005 | 26.354 | .004 |
| 73.0 | -1.799 | .003 | 32.754 | .007 | 26.354 | .005 |
| 74.0 | -1.799 | .003 | 32.754 | .007 | 26.354 | .005 |
| 75.0 | -1.799 | .003 | 32.754 | .007 | 26.355 | .006 |
| 76.0 | -1.799 | .001 | 32.754 | .009 | 26.355 | .007 |
| 77.0 | -1.799 | .001 | 32.756 | .007 | 26.357 | .006 |

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CRUISE 15-80-020 8RIDPORT INLET-80
STATION CM-9 EXPERIMENT NO.S 3303-D5,3411-13 LAT. 74-59.8 N LONG. 108-51.7.W WATER DEPTH 50.0 M

| PRESSURE |  |
| :--- | :--- | :--- | :--- |
| (DBARS) MEAN | (DEG.C) (DEG.C) |


| 2.0 | -1.807 | . 028 |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 3.0 | -1.793 | .007 | 32.691 | . 042 | 26.304 | . 035 |
| 4.0 | -1.789 | .017 | 32.679 | . 045 | 26.294 | . 037 |
| 5.0 | -1.784 | . 055 | 32.680 | . 044 | 26.294 | . 036 |
| 6.0 | -1.782 | .054 | 32.679 | . 044 | 26.293 | . 036 |
| 7.0 | -1.781 | . 055 | 32.679 | .043 | 26.294 | . 036 |
| 8.0 | $-1.774$ | .053 | 32.681 | . 040 | 26.295 | . 034 |
| 9.0 | -1.768 | . 059 | 32.686 | . 034 | 26.299 | . 028 |
| 10.0 | -1.770 | . 061 | 32.691 | . 019 | 26.303 | . 016 |
| 11.0 | $-1.774$ | . 046 | 32.696 | .010 | 26.307 | .009 |
| 12.0 | -1.780 | .031 | 32.703 | .007 | 26.313 | -005 |
| 13.0 | -1.782 | . 026 | 32.706 | . 010 | 26.315 | . 008 |
| 14.0 | -1.784 | . 022 | 32.710 | .013 | 26.319 | . 011 |
| 15.0 | -1.785 | . 018 | 32.712 | . 013 | 26.321 | . 011 |
| 16.0 | -1.787 | . 014 | 32.714 | . 010 | 26.322 | . 009 |
| 17.0 | $-1.788$ | .013 | 32.715 | . 012 | 26.323 | .010 |
| 18.0 | -1.788 | . 013 | 32.716 | . 012 | 26.323 | . 010 |
| 19.0 | -1.789 | .016 | 32.718 | . 015 | 26.326 | . 012 |
| 20.0 | -1.789 | . 014 | 32.721 | . 012 | 26.328 | . 010 |
| 21.0 | -1.788 | . 010 | 32.722 | . 011 | 26.329 | . 009 |
| 22.0 | -1.789 | . 007 | 32.725 | .005 | 26.331 | . 004 |
| 23.0 | -1.791 | . 008 | 32.726 | . 017 | 26.332 | - 014 |
| 24.0 | -1.792 | . 005 | 32.729 | .013 | 26.334 | . 010 |
| 25.0 | -1.792 | .005 | 32.728 | . 011 | 26.333 | .009 |
| 26.0 | -1.793 | .002 | 32.728 | .013 | 26.334 | .011 |
| 27.0 | -1.792 | . 005 | 32.728. | . 016 | 26.334 | .013 |
| 28.0 | -1.793 | . 004 | 32.729 | . 015 | 26.335 | .013 |
| 29.0 | -1.794 | . 005 | 32.729 | . 015 | 26.334 | -013 |
| 30.0 | -1.793 | .003 | 32.729 | . 014 | 26.334 | . 011 |
| 31.0 | -1.794 | .004 | 32.731 | .016 | 26.336 | .013 |
| 32.0 | -1.794 | .003 | 32.730 | . 014 | 26.335 | . 012 |
| 33.0 | $-1.794$ | .005 | 32.730 | .013 | 26.336 | .011 |
| 34.0 | $-1.794$ | .004 | 32.729 | . 017 | 26.335 | .014 |
| 35.0 | -1.794 | . 003 | 32.730 | .014 | 26.335 | . 012 |
| 36.0 | -1.794 | . 004 | 32.730 | -018 | 26.335 | . 014 |
| 37.0 | -1.794 | -004 | 32.734 | . 018 | 26.339 | . 015 |
| 38.0 | -1.793 | . 005 | 32.729. | .017 | 26.334 | .014 |
| 39.0 | -1.796 | .005 | 32.736 | . 018 | 26.340 | -014 |
| 40.0 | -1.793 | .006 | 32.731 | . 021 | 26.336 | -017 |
| 41.0 | -1.797 | . 013 | 32.733 | . 019 | 26.338 | .015 |
| 42.0 | -1.794 | . 007 | 32.733 | .020 | 26.338 | .016 |
| 43.0 | -1.799 | . 011 | 32.737 | . 020 | 26.341 | . 016 |
| 44.0 | -1.795 | - 007 | $32.735^{\circ}$ | . 018 | 26.339 | .015 |



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CRUISE 15-80-020 BRIDPORT INLET-8D
STATION CM-1O EXPERIMENT NO.S 3300-02.3414-16 LAT. 75-01.1 N LONG. 108-52.9 W WATER DEPTH 65.0 M
PRESSURE T MEAN T RANGE $S$ MEAN $S$ RANGE SIGMAT SIGMAT

| 2.0 | -1.805 | .029 |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 3.0 | -1.791 | .010 | 32.708 | . 035 | 26.317 | . 029 |
| 4.0 | -1.791 | . 010 | 32.706 | . 034 | 26.316 | . 028 |
| 5.0 | $-1.787$ | . 008 | $32.707 \ldots$ | . 032 | 26.316 | . 026 |
| 6.0 | -1.790 | .006 | 32.714 | .024 | 26.322 | . 020 |
| 7.0 | $-1.7 .94$ | . 012 | 32.721 | .009 | 26.328 | .008 |
| 8.0 | -1.795 | . 012 | 32.726 | .013 | 26.332 | . 011 |
| 9.0 | -1.795 | . 012 | 32.727 | . 016 | 26.333 | . 013 |
| 10.0 | $-1.795$ | . 012 | 32.728 | .017 | 26.333 | . 014 |
| 11.0 | -1.795 | .012 | 32.728 | . 019 | 26.334 | . 016 |
| 12.0 | -1.795 | .013 | 32.729 | . 020 | 26.334 | . 017 |
| 13.0 | -1.795 | .011 | 32.730 | .019 | 26.335 | .016 |
| 14.0 | -1.795 | .013 | 32.731 | .020 | 26.336 | . 017 |
| 15.0 | -1.7.95 | .011 | 32.730 | . 021 | 26.336 | . 018 |
| 16.0 | -1.795 | . 011 | 32.730 | .020 | 26.335 | . 017 |
| 17.0 | -1.795 | . 013 | 32.731 | . 021 | 26.336 | . 018 |
| 18.0 | $-1.795$ | .012 | 32.731 | . 022 | 26.336 | . 018 |
| 19.0 | -1.795 | . 012 | 32.732 | .023 | 26.336 | . 019 |
| 20.0 | -1.794 | .013 | 32.730 | . 024 | 26.336 | .020 |
| 21.0 | -1.795 | . 012 | 32.732 | . 023 | 26.337 | . 019 |
| 22.0 | -1.795 | .010 | 32.731 | . 022 | 26.336 | . 018 |
| 23.0 | -1.794 | . 011 | 32.731 | .025 | 26.336 | . 020 |
| 24.0 | -1.794 | .012 | 32.730 | . 024 | 26.335 | .020 |
| 25.0 | -1.794 | . 011 | 32.731 | .024 | 26.336 | . 020 |
| 26.0 | -1.795 | .011 | 32.732 | . 024 | 26.337 | .020 |
| 27.0 | -1.795 | . 013 | 32.732 | . 025 | 26.337 | .020 |
| 28.0 | -1.794 | . 011 | 32.732 | .024 | 26.337 | .020 |
| 29.0 | -1.794 | .013 | 32.733 | . 025 | 26.337 | . 021 |
| 30.0 | -1.794 | . 01.4 | -32.731 | .027 | 26.336 | .022 |
| 31.0 | -1.795 | .012 | 32.733 | . 024 | 26.337 | .020 |
| 32.0 | -1.7.95 | .013 | 32.734 | .023 | 26.339 | .019 |
| 33.0 | -1.791 | .013 | 32.728 | . 023 | 26.334 | .019 |
| 34.0 | $-1.796$ | .013 | 32.737 | .025 | 26.341 | . 021 |
| 35.0 | -1.792 | . 014 | 32.729 | . 025 | 26.335 | . 020 |
| 36.0 | -1.797 | . 016 | 32.739 | . 022 | 26.343 | .019 |
| 37.0 | -1.795 | . 015 | 32.735 | . 021 | 26.339 | -018 |
| 38.0 | -1.793 | .013 | 32.733 | . 021 | 26.338 | .017 |
| 39.0 | -1.794 | .012 | 32.733 | . 019 | 26.337 | -016 |
| .40 .0 | -1.795 | . 013 | . 32.736 | .020 | 26.340 | .017 |
| 41.0 | -1.794 | . 013 | 32.734 | . 022 | 26.338 | . 018 |
| 42.0 | -1.794 | . 013 | 32.735 | .020 | 26.339 | .016 |
| 43.0 | $-1.795$ | .013 | 32.736 | . 020 | 26.340 | . 016 |
| 4.4.0 | -1.796 | .013 | 32.738 | . 020 | 26.341 | .017 |


| PRESSURE | TMEAN | T RANGE | SMEAN | S RANGE | SIGMAT | SIGMAT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (DGARS) | (DEG.C) | (DEG.C) |  |  | MEAN | RANGE |
| 45.0 | -1.793 | .013 | 32.735 | .022 | 26.339 | .018 |
| 46.0 | -1.795 | .013 | 32.738 | .021 | 26.342 | .017 |
| 47.0 | -1.796 | .011 | 32.739 | .018 | 26.343 | .015 |
| 48.0 | -1.795 | .012 | 32.738 | .017 | 26.342 | .014 |
| 49.0 | -1.795 | .011 | 32.739 | .017 | 26.343 | .014 |
| 50.0 | -1.797 | .008 | 32.744 | .008 | 26.347 | .007 |
| 51.0 | -1.796 | .007 | 32.745 | .007 | 26.347 | .006 |
| 52.0 | -1.798 | .007 | 32.746 | .007 | 26.348 | .006 |
| 53.0 | -1.798 | .007 | 32.746 | .005 | 26.348 | .004 |
| 54.0 | -1.798 | .007 | 32.748 | .003 | 26.349 | .003 |
| 55.0 | -1.799 | .007 | 32.748 | .005 | 26.350 | .004 |
| 56.0 | -1.799 | .004 | 32.749 | .003 | 26.351 | .002 |
| 57.0 | -1.799 | .005 | 32.749 | .004 | 26.351 | .003 |
| 58.0 | -1.799 | .005 | 32.750 | .004 | 26.351 | .003 |
| 59.0 | -1.799 | .004 | 32.751 | .006 | 26.352 | .005 |
| 60.0 | -1.799 | .004 | 32.750 | .008 | 26.352 | .006 |
| 61.0 | -1.799 | .004 | 32.751 | .007 | 26.352 | .005 |
| 62.0 | -1.798 | .003 | 32.752 | .006 | 26.353 | .005 |

APPENDIX B
ULTRA SONIC CURRENT METER PROFILES


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