

Geological Survey
of Canada



Commission géologique
du Canada

**ORGANIC PETROGRAPHY AND KINETICS OF LIMESTONE
AND SHALE SOURCE ROCKS IN WELLS ADJACENT TO SABLE
ISLAND, NOVA SCOTIA AND THE INTERPRETATION ON OIL-OIL
OR OIL-SOURCE ROCK CORRELATION
AND BASIN MODELLING**

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SSC FILE NO. OSC93-00336-(014)
CONTRACT NO. 23420-3-C056/01-OSC
FINANCIAL CODE:230 120 0000720104 04B9

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JANUARY 31, 1994

OPEN FILE NO. 3167



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ABSTRACT

This report compared the hydrocarbon potential and maturity of fifty limestone (and calcareous shale) and fifty composite shale units based on organic facies, which uses maceral composition, fluorescence and oxidation criteria of macerals and minerals, hydrogen and oxygen indices (from Rock-Eval pyrolysis), and maturity (based on vitrinite reflectance).

Volumetrically, the comparable source rock potential of Limestone and calcareous shale are: 6% oil- and condensate-prone Kerogen Type IIA-IIB; 42% condensate- and gas-prone Kerogen Type IIB; 34% gas-prone Kerogen Type III; and 18% nonsource Kerogen Type IV. For Composite shale the potentials are: 20% oil- and condensate-prone Kerogen Type IIA-IIB; 40% condensate- and gas-prone Kerogen Type IIB; 40% gas-prone Kerogen Type III; and none of them are considered as nonsource for hydrocarbons.

The activation energy and Arrhenius Constant of Kerogen Type IIA-IIB source rock (N. Triumph B-52) suggest that 80% of the source rock is converted to oil and gas before a maturity of 0.8% R_o . On the other hand, Kerogen Type III source rock (Venture B-52) yields only 30-40% conversion at that maturity level. Activation energy distribution and Arrhenius Constant of two limestone source rocks indicate hydrocarbon generation at a late stage of maturity (0.9 to 1.1% R_o).

Based on aromatic GC-MS, the majority of petroleums and rock extracts in this study are at advanced levels of thermal maturity (middle to late oil window), as evidenced by the distribution of methylphenanthrene, dimethylphenanthrene, trimethylnaphthalene and tetramethylnaphthalene isomers. This is not surprising, given the great burial depth (4 to 6 km) of many of the samples.

Multiple linear regression and subsequent cluster analysis reveals a closely related family of oils ("Group A") including those from N. Triumph B-52, S. Venture O-59 (DST 10), Glenelg, Olympia, Thebaud, Penobscot, Bluenose, Arcadia, Venture H-22, Citnalta and Banquereau, all produced from Missisauga or Mic Mac reservoirs. The Cohasset oils form a second group and the Panuke oils do not seem to be related to each other or to any other oil. A group of rock extracts ("Cluster 1") from the lower and middle members of the Missisauga, along with the Penobscot Limestone and Cree Member samples, correlate well with each other, indicating similar organic facies. These samples also have reasonable affinities for the Group A oils, with

the Lower Mississauga samples from the Abenaki and S. Desbarres wells appearing to be the best candidate source rocks. A second cluster of extracts, including samples from the Verrill Canyon and Mississauga Formations are related to one another, but correlate neither with Group A oils nor Cluster 1 extracts.

To supplement the cluster analysis, a parameter related to the sulfur content of the samples (the ratio of the sum of C₂-alkyldibenzothiophenes to the sum of the C₂-alkylphenanthrenes) was established. It supports the statistical findings, further indicating that the oils and Cluster 1 extracts have a similarly low thiophenic sulfur content, in contrast to the relatively high sulfur Cluster 2 and unclassified extracts. Samples richer in sulfur derive from marine rocks, while the remainder are apparently of non-marine origin, a conclusion supported by the organic petrologic and stratigraphic data.

1-D basin modeling of two key wells (Abenaki J-56, carbonate-rich; W. Chebucto K-20, clastic-rich) showed that in W. Chebucto K-20 source rocks generated and migrated cumulative hydrocarbons five times more than Abenaki J-56 source rocks. Major hydrocarbons (both oil and gas) have been generated and expelled between 137 and 50 mybp from the Mississauga source rocks in the W. Chebucto K-20 well. Non-expelled oils in deeper source rocks have been cracked to gas and residue.

INTRODUCTION

Administrative Aspect

This research proposal was requested by Supply and Services Canada, Dartmouth, Nova Scotia at the initiation of the Basin Analysis Subdivision of the Atlantic Geoscience Centre, Geological Survey of Canada, Bedford Institute of Oceanography. Global Geoenergy Research Ltd. of Halifax, Nova Scotia, submitted a financial and work schedule for the research proposal. The proposal was accepted and the research work was started July 8, 1993. Canada Nova Scotia Offshore Petroleum Board (CNSOPB), Halifax, Nova Scotia, on our request, permitted us to collect four crude oil and condensate samples and 100 unwashed cuttings samples from selected Scotian Shelf wells at the CNSOPB Repository, Dartmouth, Nova Scotia. According to the contract, Rock-Eval pyrolysis of the 100 samples was done at the ISPG Laboratory at Calgary, Alberta, and kerogen isolation, smear slide preparation, kerogen plug preparation were done at

the palynology and coal laboratory at the AGC, Dartmouth, N. S. Bitumen extraction and liquid chromatography of 4 oil and 3 source rock samples was subcontracted to Dr. Michael Kruge, Southern Illinois University, Carbondale, Illinois, USA who also did special aromatic GC-MS analyses. Kinetic analysis of four source rock samples were subcontracted to Dan Jarvie, Humble Geochemical Services, Humble, Texas.

Scientific Aspect

During the past twenty years, significant gas and associated condensate or light oil were discovered in the Jurassic-Cretaceous reservoirs around the Sable Subbasin of the Scotian Basin. The major issues on hydrocarbon generation, migration, and entrapment, which still remained unsolved included: (a) the source rock types in various stratigraphic intervals; (b) regional maturity variation for both source rocks and crude oils and condensates; (c) the relation between hydrocarbon generation and overpressuring; and (d) possible oil-oil and oil-condensate-source rock correlation.

Since 1988, Basin Analysis Subdivision of the Atlantic Geoscience Centre of the Geological Survey of Canada, Dartmouth, Nova Scotia initiated systematic research projects to characterize and resolve these problems. These studies (as done by Global Geoenergy Research Ltd.) characterized some of the source rocks in various stratigraphic intervals and characterized the geochemical properties of some selected oils and condensates, possible oil-oil or oil-source rock correlation, and quantified the timing and amount of hydrocarbon generation from various source rock intervals using BASIN-MOD program (Mukhopadhyay, 1989; Mukhopadhyay and Wade, 1990; Mukhopadhyay, 1990a; Mukhopadhyay, 1990b; Mukhopadhyay, 1991, Mukhopadhyay, 1993). Mukhopadhyay (1991) from the extensive organic petrography and Rock-Eval analyses established the proportion of oil-, condensate-, and gas-prone source rocks in the various stratigraphic sequences. Mukhopadhyay (1991, 1990b, 1993) using the data from Kruge (1990, 1991, 1993) on the aromatic biomarkers and pyrolysis-gas chromatography of the asphaltene fractions of source rock extract or oils and isotope analyses indicated possible correlation of oil-oil and oil-source rock pairs. Accordingly, Mukhopadhyay (1991, 1993) indicated that the various crude oils or condensates may have been derived from three types of source rocks.

In the Scotian Basin, source rock studies during 1989-1993 were determined mainly in

selected clastic rock intervals. Carbonates and evaporites may also be a source rock for liquid hydrocarbons as seen in the Gulf Coast and various basins of the Middle East (Palacas, 1984). The geochemically unique Cohasset/Panuke/Balmoral oil discoveries are in close proximity to the Abenaki carbonate bank. This stratigraphic unit was not studied earlier as a possible source rock for this oil. Moreover, kinetics analysis suggests a diversity in kinetics properties in various source rocks intervals (Mukhopadhyay, 1993). Another systematic study on kinetics analyses from more samples is extremely important to achieve a mass balance and timing of hydrocarbon generation and expulsion. In order to establish a mass balance between hydrocarbon generation, expelled oil and reservoired petroleum, it is pertinent to compile a source rock characterization on the composite shale and limestone units which have already been indicated as good sources for liquid and gaseous hydrocarbons. Basin modeling studies from some additional wells may contribute to the mass balance of oil and gas generation and migration within the Scotian Basin. Considering all those unanswered questions, this research project was initiated by AGC.

The objectives of this study are: (a) to evaluate the Abenaki Formation limestone and calcareous shale as a source rock for the Cohasset - Panuke or any other oils; (b) to evaluate the kerogen type and hydrocarbon potential of composite shale sections from the Lower Cretaceous and Jurassic sequences of the Scotian Basin; (c) to determine the activation energies and frequency factors of some selected samples; (d) to use various aromatic biomarkers and their ratios from more petroleum and source rock samples to determine the level of maturity and oil-oil or oil-source rock correlation; and (e) to evaluate mass balance of hydrocarbon generation of source rocks and reservoired petroleum by using numerical basin modeling (by BASIN-MOD™ of Platt River Associates, 1992), similar to the work done by Mukhopadhyay (1993) and Mukhopadhyay et al., (in press).

SAMPLES AND ANALYTICAL PROCEDURES

Samples

Fifty limestone and calcareous shale samples from the Abenaki Formation were chosen for Rock-Eval pyrolysis, organic facies, and source rock potential. Those samples are selected from the following wells: Abenaki J-56 (10,700-14,450'), Cohasset D-42 (10,000-14,525'), Cohasset L-97 (3100-4900 m), Demoscota G-43 (11,400-15,300'), Penobscot L-30 (11,200-

14,000'), Uniacke G-72 (4200-5745 m).

Similarly, fifty composite shale samples from the Logan Canyon, Missisauga, Verrill Canyon and Mic Mac Formations were chosen from the following wells for Rock-Eval pyrolysis, maceral analysis, and vitrinite reflectance: Alma F-67 (3945-4890 m), Chebucto K-90 (2450-5090 m), Glenelg J-48 (2850-5025 m), N. Triumph B-52 (2660-3715 m), N. Triumph G-43 (2660-4755 m), S. Desbarres O-76 (2345-5705 m), Venture B-43 (5310-5475 m), Venture B-52 (2310-5960 m), Venture H-22 (3860-4710 m), and W. Chebucto K-20 (2770-5000 m).

Kinetics analysis by Rock-Eval pyrolysis methods using Rock-Eval 5 on two selected composite shale (N. Triumph B-52, 3560-3715 m and Venture B-52, 3680-3740 m) and two Abenaki Formation calcareous shale limestone (Abenaki J-56, 10090-10130' and Demoscota G-43, 11840-11880') were done using the methods as described by Mukhopadhyay (1993).

Four light oil/condensate samples (Balmoral M-32, DST 3, Panuke PP3 J-99, DST 1, Uniacke G-72, DST 6, Venture H-22, DST 7) and three limestone source rock sample (Abenaki J-56, 10090-10130'; Demoscota G-43, 11840-11880'; and Penobscot L-30, 13810-13840') were selected for maturity determination and oil-oil or oil-source rock correlation, using aromatic biomarkers in GC-MS.

Analytical Procedures

For the determination of kerogen type by organic petrography, three types of sample preparation were used: kerogen smear slide, whole rock polished pellet, and kerogen polished pellet. We used incident and transmitted white and blue light excitation. The terminologies used for maceral composition and kerogen type determination are from Stach et al. (1982), Mukhopadhyay et al., (1985), Senftle et al., (1987), Teichmuller (1986), Hutton, (1987), and Mukhopadhyay (1989). Details on source-rock characterization using organic facies are shown in Mukhopadhyay and Wade (1990).

Vitrinite reflectance was measured using both whole rock and kerogen pellets and Zeiss Axioskop with MPM 21 Controller for MPM 03 Photomultiplier. Except for samples from W. Chebucto K-20, vitrinite reflectance on other samples were measured by M. P. Avery of the Atlantic Geoscience Centre. Rock-Eval pyrolysis was carried out on selected washed cuttings, washed/hand-picked cuttings, and from the conventional cores using the Rock-Eval II equipment. For details of Rock-Eval instrumentation and evaluation of various parameters, see Espitalie et

al. (1985).

Bitumen extraction was performed using Soxhlet for 24 hours with dichloromethane. Liquid chromatography was done using an open column packed with activated silica, eluting saturate, aromatic and two polar fractions with *n*-hexane, 9:1 *n*-hexane:CH₂Cl₂, CH₂Cl₂, and 1:1 methanol:CH₂Cl₂, respectively.

For analytical procedures for kinetic analysis, see Mukhopadhyay (1993). Basin modeling was done using the BASIN-MOD package provided at the AGC. For details on BASIN-MOD program, see Mukhopadhyay (1993).

For analytical methods using aromatic biomarker GC-MS, see the chapter on oil-oil or oil-source rock correlation chapter.

SOURCE ROCK CHARACTERIZATION

Rock-Eval Pyrolysis

Table 1A illustrates the data of various Rock-Eval pyrolysis parameters (S_1 , S_2 , S_1+S_2 , S_3 , PI, HI, OI, T_{max} , and TOC) for whole rock and isolated kerogens from 50 limestone and calcareous shale samples. The total organic carbon of the whole rock limestone and calcareous shale samples varies between 0.06 (Cohasset D-42) to 2.39% (Abenaki J-56). Some samples from Uniacke G-72 well which have >8% TOC, are considered as contaminated with drilling additives, which is based on organic petrographic study. Except for a few samples from both Cohasset wells, most samples have >0.3% TOC (minimum threshold for potential source rock for liquid hydrocarbons; Tissot and Welte, 1984). Abenaki J-56 and Penobscot L-30 samples, on an average, have more than 0.7% TOC. High TOC content of the isolated kerogens suggest better separation of mineral matter of limestone and calcareous shale to isolate kerogen. The total organic carbon content of all composite shales is >0.5% and most samples from Alma F-67, N. Truimph B-52/G-43 and S. Desbarres O-76 have TOC >1.0% (Table 1B).

Considering the depth and total organic carbon content of the samples, the amount of S_1 (thermal extraction below 300°C) fraction in most limestone samples, is considered as extremely high (Table 1A)(Tissot and Welte, 1984). Anomalously high S_1 content may be caused by: (a)contamination from drilling mud additives, especially hydrocarbons; (b) early generation of hydrocarbons; and (c) soaking of allochthonous bitumen or migrated petroleum. At this stage, it

is difficult to predict the cause of this anomaly. The S_1 and PI of all composite shale samples were also found to be anomalously high (Table 1B).

Except for a few samples from Abenaki J-56 well, the S_2 fraction of most samples have <1 mg HC/g of rock suggesting extremely low potential for crude oil (Table 1A). The low S_2 can also be caused by the retention of liquid hydrocarbons (derived from cracking of kerogen) within the mineral matrix (Espitalie et al., 1985; Mukhopadhyay, 1989 and references therein). The high S_2 values of the isolated kerogens also support the idea. Samples from Alma F-67, N. Triumph B-52 and G-43, and S. Desbarres O-76 have significantly high S_2 content in the composite shale units suggesting presence of abundant potential oil-prone source rocks compared to limestone and calcareous shale samples.

Except in the Penobscot L-30 and Abenaki J-56 wells, T_{max} (maturation parameter) values in both limestone and composite shale samples do not show any increasing trend with increasing depth of the samples. The suppression of T_{max} may be caused by the anomalously high S_1 (free hydrocarbons) content of the samples.

Plots of HI vs. OI and HI vs. T_{max} (Tissot and Welte, 1984; Espitalie et al., 1985) show the maturation path of Kerogen Types I, II, III, and IV (Figs. 1A through 1C) and Kerogen Types I, II, and III (Figs. 2A through 2C). Figures 1A and 1B or 2A and 2B show comparable data for the whole rock and kerogen concentrate of the limestone and calcareous shale samples. Figures 1C and 2C show the plot of the composite shale whole rock samples. Comparing Fig. 1A and 1B, about 50% of the samples are falling within the Kerogen Types II and III, whereas a vast majority of kerogen concentrate samples lie on the Kerogen Type I and II path. The grouping of kerogen samples within Kerogen Type I and II are due to selective removal of oxygen-functional groups during the kerogen isolation procedure. Comparing Figs. 2A and 2B, a vast majority of whole rock samples are below Kerogen Type III, whereas most kerogen samples are within Kerogen Types II and III or close to Kerogen Type III suggesting a mineral matrix effect, Mukhopadhyay, 1989). A majority of the composite shale whole rock samples, however, lie within Kerogen Type II and III maturation path (Figs. 1C and 2C).

Maturation Based on Vitrinite Reflectance

Vitrinite reflectance data (Tables 1A and 1B) of limestone samples suggest that all samples are considered as mature (between 0.5 to 1.3% R_o) to overmature ($>1.4\%$ R_o). In most

cases, vitrinite reflectance values increase with the depth of the sediment. The vitrinite reflectance values in the deeper sediments, in most cases, do not correlate with the T_{max} values which are possibly caused by bitumen absorption within the whole rock matrix. In Penobscot L-30, however, the increase in vitrinite reflectance (0.61 to 1.02% R_o) correlates well with T_{max} values (439 to 452°C). In the composite shale samples, the maturity follows the stratigraphic depth. The lowest maturity (0.44% R_o) is observed at N. Triumph B-52 at 2860-3115 m, whereas the highest maturity is in S. Desbarres O-76 at 5590-5715 m (1.78% R_o).

Organic Facies and Source Rock Potential

The overwhelming majority of the limestone and calcareous shale samples have abundant refractory macerals (fusinite, semifusinite, clustered micrinite, inertinitized algae, and inertodetrinite) indicating oxidation at the depositional interface (Table 2A). Those samples also contain other terrestrial macerals like vitrinite, exinite (resinite, sporinite etc.) and AOM 3. This suggest a terrestrial influx within the Abenaki carbonate platform. However, there is also abundant algal derived organic matter (telalginite, lamalginite, exinite [algal spore], AOM 2). The algal organic matter is mostly oxidized and has lost its fluorescence. Within the exinite group, oxidized algal spore are common (example: Demoscota G-43, 11840'). As a result, most limestones and calcareous shale samples formed either Kerogen Types III or IIB. Based on organic facies, only a few samples from Abenaki J-56 well can be considered as Kerogen Type IIA-IIB having fluorescent lamalginite and AOM 2. On the other hand, composite shale samples from Alma F-67, N. Triumph B-52 or G-43, and W. Chebucto K-20 contain more marine macerals (lamalginite, telalginite, and AOM 2) which have higher fluorescence than most limestone samples (Table 2B).

Volumetrically, the comparable source rock potential between limestone and composite shale samples can be illustrated as follows:

Limestone and calcareous shale: 6% are oil- and condensate-prone Kerogen Type IIA-IIB; 42% are condensate- and gas-prone Kerogen Type IIB; 34% are gas-prone Kerogen Type III; and 18% are nonsource Kerogen Type IV.

Composite shale: 20% are oil- and condensate-prone Kerogen Type IIA-IIB; 40% condensate- and gas-prone Kerogen Type IIB; 40% are gas-prone Kerogen Type III. None of the shale samples are considered as nonsource for hydrocarbons.

KINETIC PROPERTIES OF SELECTED SOURCE ROCKS

Table 3 shows the activation energy distribution related to the percentages of hydrocarbon generated and their corresponding Arrhenius Constant from the selected four source rock samples. N. Triumph B-52 (Kerogen Type IIA-IIB) sample is similar to the N. Triumph G-43 sample which was analyzed in 1993 (Mukhopadhyay, 1993). The limestone samples from Abenaki and Demoscota show extremely slow reaction rates. More than 80% of the hydrocarbons from those two samples are converted between 49 and 52 kcal/mole, whereas 80% of the hydrocarbons in the N. Triumph B-52 sample are converted below 47 kcal/mole. The kinetic data of the Kerogen Type III source rock from Venture B-52 are similar to the default values for Kerogen Type III in the BASIN-MOD program. The activation energies and Arrhenius Constant values of these samples are comparable to kinetics data of Monterrey, Phosphoria, Alum and Woodford shales determined by Hunt et al. (1991) using hydrous pyrolysis. Accordingly, the reaction rates for hydrocarbon generation from N. Triumph B-52 sample are considered as medium-fast (similar to Phosphoria Shale) and limestone samples are slow. These values are also different from the default values of Kerogen Type II used in the BASIN-MOD program.

Figures 3b, 3c, 3d and 3e show the distribution of activation energies when plotted against % HC conversion for Abenaki J-56, Demoscota G-43, N. Triumph B-52, and Venture B-52 wells, respectively. Figure 3d (N. Triumph B-52, Kerogen Type IIA-IIB) shows smooth peaks of activation energy around 44 and 48 kcal/mole. In contrast to that Venture B-52 (kerogen Type III) sample shows a sharp peak of activation energy between 46 and 50 kcal/mole (Fig. 3e). Neither of them are comparable to the standard Kerogen Type I Green River Shale (Fig. 3a). Both Abenaki J-56 and Demoscota G-43 limestone samples show sharp peaks of activation energies between 50 and 52 kcal/mole (Figs. 3b and 3c). They show some similarity of activation energy distributions of Green River Shale. This possibly suggests an algal affinity for those samples.

OIL-SOURCE ROCK CORRELATION AND MATURITY DETERMINATION USING AROMATIC MOLECULAR MARKERS

In the last two decades, great advances have been made in the application of biological marker compounds to the determination of maturation and organic facies of sedimentary organic matter.

However, these compounds, generally polycyclic alkanes, cannot always present a complete picture. For example, maturation indicators based on the saturate hydrocarbons are ineffective in samples at mid to late oil window maturity levels.

The aromatic fractions in fossil fuels provide many additional clues. A wide variety of polyaromatic compounds have been documented in fossil fuels (Later *et al.*, 1981; White and Lee, 1980; Whitehurst *et al.*, 1982; Radke *et al.*, 1984b; Villar *et al.*, 1988; Rowland *et al.*, 1984). Distributions of these compounds have been shown to be sensitive to maturation (Alexander *et al.*, 1983; 1986, Radke, 1987; 1988; Radke *et al.*, 1980; 1982a; 1982b; 1984a; 1986, Garrigues *et al.*, 1988, Budzinski *et al.*, 1991; 1993a). In particular, the Methylphenanthrene Index (MPI) of Radke and Welte (1983) has been proven useful in maturity assessment, particularly in the middle to late oil window range. Recently, Kruge and others (1989; 1990a; 1990b, 1992) have documented the maturation and organofacies sensitivity of a series of polyaromatic molecular markers in marine shale extracts, coal extracts and coal pyrolyzates.

Aromatic fractions of 39 Scotian Shelf oils, condensates and rock extracts were analyzed by gas chromatography/mass spectrometry (GC/MS) to determine levels of thermal maturation and to correlate the petroleums with one another and with candidate source rocks.

Analytical Methods

Aromatic fractions of 17 petroleums (oils and condensates) and 15 extracts of candidate source rocks (Table 4) were analyzed by a Hewlett Packard 5890A gas chromatograph, coupled to an HP 5970B Mass Selective Detector. The GC was held initially at 100° C for 10 min., then raised to 300° at 3°/min., where it was held for 5 minutes. A 25 m OV-1 column with 0.2 mm inside diameter and 0.33 mm film thickness was employed. The mass spectrometer was run in selective ion monitoring mode, collecting data on the following ions: m/z 91, 156, 162, 168, 169, 170, 176, 182, 184, 192, 196, 198, 202, 206, 212, 216, 220, 228, 230, 231, 242, 252, 253, 256, 268, 282, 296 and 310, which are either the molecular ions of the most common polyaromatic compounds or are principal fragment ions of aromatized biomarkers, such as steroids. Two samples were rerun using full scan mode, to confirm that all principal aromatic peaks were being recorded. Quantitations were done on the molecular ions of the compounds of interest, using the selected ion monitoring data and the Hewlett Packard data system.

GC Peak Recognition

A total of 91 GC/MS peaks representing polyaromatic compounds (Fig. 4) were recognized and quantified in each of the 39 samples. The dimethylnaphthalenes, while dominant in the condensates, were often truncated in the rock extracts, due to the evaporation necessary in the sample preparation. The triaromatic steroids were extremely weak in many of the samples. Therefore these two compound groups were eliminated from further consideration. Table 6 lists the 80 remaining compounds, identified as far as possible by their mass spectral properties and by reference to the literature (Rowland et al., 1984; Wang and Simoneit, 1990; Budzinski et al., 1991; 1993b; 1993c). The chromatographic character of the peaks are more clearly illustrated on the individual mass chromatograms, exemplified by those of lower Missisauga rock extract from Migrant N-20 (Figs. 5-9). As the mass chromatograms show, the polyaromatic peaks are often clusters of isomers, some of which still cannot be specifically identified at the present time, due to the similarity of their mass spectra and the lack of commercially available authentic standards. However, most of the members of each cluster can be separated by GC and the variations in their relative peak distributions from sample to sample can provide much useful empirical information as to maturation levels and organic matter type. The individual GC peaks are labeled according to their mass spectrometric base peaks, which for all compounds under consideration are the molecular ions, and by their order of elution. For example, the eight C₃-alkylnaphthalene peaks, with a molecular ion of m/z 170, are designated peaks 170a through 170h (Fig. 5). Closely eluting peaks, such as the two which comprise peak 212g (Fig. 6), were not resolved in all samples and are summed as a single peak for consistency. There are several exceptions to the labeling system, designating 1) unsubstituted polyaromatics such as phenanthrene (PHN) or 2) heteroatomic aromatics which have molecular ions coincident with hydrocarbons, such as dibenzothiophene (DBT), sharing the m/z 184 trace with C₄-alkylnaphthalenes.

Data Handling

Due to the great number of chromatograms generated from this large data set, visual comparisons were not sufficiently thorough for petroleum-to-petroleum or petroleum-to-rock correlation. A multivariate statistical approach was chosen as the most effective. Once the raw integrated areas of the 80 peaks were transferred to a minicomputer spreadsheet, they could then be normalized. Condensates, such as the South Venture DST 5 (API gravity = 51°), show a much

greater preponderance of lighter aromatic compounds than do relatively low maturity rock extracts, such as the extract from the Abenaki well (Fig. 10). To be able to better compare such samples for possible genetic relationships, the data were normalized within three groups according to retention time (Fig. 4). The first-eluting group includes the compounds with molecular ions at m/z 170, 184 and 182 (C_3 - and C_4 -alkylnaphthalenes, dibenzothiophene, dimethylbiphenyls and methyldibenzofurans). The second group includes compounds with molecular ions at m/z 178, 192, 206, 198, 212 and 202 (phenanthrene, methyl- and dimethylphenanthrenes, methyl- and dimethyldibenzothiophenes, pyrene and fluoranthene). The last-eluting group includes compounds with molecular ions at m/z 216, 230, 228, 242 and 252 (methyl- and dimethylpyrenes and their isomers, chrysene and benzo[*a*]anthracene, methylchrysene isomers, and pentaaromatic hydrocarbons including benzo[*a*]pyrene and its isomers).

Since the variation in peak intensities is so wide, the data must be scaled to avoid distortion during the correlation exercise. Scaling was performed by taking the base 10 logarithm of the normalized peak intensities, in which the tallest peaks (100%) now has a value of 2 and very small or missing peaks (with normalized areas of $< 0.1\%$) are assigned values of -1.

The variations among the Scotian Shelf samples are functions of both maturation and organic matter type. To separate these two effects, methylphenanthrene indices (MPI, Radke and Welte, 1983) and several other maturity parameters using polyaromatic hydrocarbons were calculated for each sample, as presented in detail in the Results and Discussion section. The samples were arranged in order of increasing maturity as so determined and the 5 samples with the lowest and the 8 with the highest maturities were selected. The base 10 logarithms of the areas (normalized by retention time group) of each of the 80 peaks in the low maturity sample set were averaged, as were those with high maturity, forming composite low and high maturity data sets. Averages were employed to cancel differences associated with organic matter type. The low maturity averaged data were subtracted from the high and the differences reported as " Δm " in Table 6. Where Δm has a large positive value, the corresponding peak increases greatly with maturation, while a large negative value indicates that the peak decreases markedly in concentration as maturity increases.

Values of Δm range widely, from -1.61 to +0.80. A subset of data was extracted comprising

those 39 peaks falling within a narrow window where $-0.22 \leq \Delta m \leq 0.17$, based on apparent boundaries in the data set, to dampen the effects of maturation. Next, average peak areas and their standard deviations were computed for all 80 peaks of all 39 samples (\log_{10} , normalized by retention time grouping). The standard deviations range from 0.07 to 0.89 (Table 6). To eliminate the peaks showing little variation (and thus little potential to aid in correlation), the subset of 39 maturation-insensitive peaks were further reduced to those 36 peaks with standard deviations > 0.15 . To confirm that maturation effects were indeed dampened, each of the 36 peaks of the latest subset were averaged for the low and high maturity samples. When these two average sets were co-plotted, they matched closely, indicating that there is little residual maturity effect. These 36 peaks are marked (*) in Table 6 and will be termed the "correlation set".

Even with the original data set scaled logarithmically, normalized by retention time groupings, and reduced from 80 to 36 peaks, visual correlation among 39 samples is inefficient and merely qualitative. The degree of similarity between samples can also be expressed graphically by cross-plotting the correlation sets. For example, if the Chebucto and North Triumph oil samples are compared, a tight linear distribution is apparent. In contrast, if the Chebucto oil is compared with the Panuke J-99 oil, much greater scatter is apparent (Fig. 11). These results can be expressed quantitatively using coefficients of determination (r^2) from the linear regression of any two samples, where r^2 varies from 0 to 1, with a value of 1 indicating a perfect match. With the assistance of a computer, over 700 r^2 values were determined for all possible pairings of the 39 samples (Table 7).

To aid in the interpretation of these results for the oil-oil and oil-source correlation exercise, a cluster analysis was undertaken. The "unweighted average linkage" method (Massart and Kaufman, 1983) was applied to the similarity matrix of Table 7. Briefly, the clustering is begun by choosing the best match from the correlation matrix and linking those two samples on a dendrogram (Fig. 13), with the length of the connecting branches scaled to the r^2 value (the shorter the branch, the better the match). Then, the r^2 values of the two matched samples with all other samples are averaged, and this average replaces the original r^2 values for the two matched samples in the similarity matrix. The next closest match is plotted and matching continues until all samples are included. When a new sample is matched with a pre-existing cluster, a weighted average is computed. When two pre-existing clusters are linked, their r^2 values

are averaged, with each weighted by the number of samples it contains. The term "unweighted average linkage" may thus at first appear to be a misnomer, but refers to the fact that the contributions of each sample are given equal importance.

Maturity Determination

Aromatic maturation parameters, including the methylphenanthrene index (MPI, Radke and Welte, 1983), the ratio of triaromatic to monoaromatic steroids, and the ratio of low to high molecular weight triaromatic steroids (Mackenzie *et al.*, 1983), are commonly applied in evaluating oils and source rock extracts. Since the present study was limited to the analysis of aromatic fractions, in part due to the lack of saturate biomarkers, it was expected at the outset, that maturation determinations could be successfully made. After analysis it became apparent that the monoaromatic steroids were not detectable in most cases and the high molecular weight (C₂₆-C₂₈) triaromatic steroids were weak or not present in the sample set. Thus the two parameters employing steroids were at their maximum values in many cases, suggesting that the overall maturity level for the sample set was rather high. This is not surprising, given the great burial depth of many of the samples (Table 4). Although the MPI was originally calibrated on terrestrial organic matter and the types of organic matter represented in this study were unknown at the outset, the author has observed that the MPI can be a useful relative indicator of high levels of maturity (mid to late oil window) for many types of organic matter. It was thus decided to begin maturation ranking of the samples by using the MPI:

$$\frac{1.5 (3\text{-mPH} + 2\text{-mPH})}{(\text{Phenanthrene} + 9\text{-mPH} + 1\text{-mPH})}$$

converting to an estimated vitrinite reflectance equivalent (R_c) using the empirical relationship of Radke (1987), valid for R_o < 1.35%:

$$R_c = 0.60 \text{ MPI} + 0.40$$

R_c values for the 39 Scotian Shelf samples range from 0.64 to 1.15, with the majority of samples (22) falling between 0.8 and 1.0 (Fig. 12). Not wishing to depend completely on a sole maturity parameter, it was expected that additional maturation indicators could be derived from other polyaromatic hydrocarbon groups. Only maturation variations among groups of isomers, i.e., the dimethylphenanthrenes, were considered for these purposes. Additional maturity

parameters might also be created, by mixing peaks from different compound classes, but the possibility was not investigated here.

Individual isomers were chosen from among the trimethyl- and tetramethyl naphthalenes and dimethylphenanthrenes based on their observed response to natural and artificial maturation and on their computed thermodynamic stabilities (Radke, 1987; Kruge and Landais, 1992; Budzinski et al., 1993a; 1993b; Kruge and Bensley, 1993). To confirm that these peaks were responding to maturation changes in the Scotian Shelf samples, the samples unquestionably exhibiting the lowest and highest maturation levels (as determined by the MPI) were averaged within their maturity groups and subtracted. In all cases, the polyaromatic hydrocarbons used to create the supplemental maturity parameters behaved as predicted — the stable showed a relative increase and the labile, a decrease with increasing maturity. The resulting additional maturity ratios are defined as follows:

Trimethylnaphthalene (TrMN) Ratio

$$\frac{(1,3,6- + 2,3,6\text{-TrMN})}{(1,4,6- + 1,3,5- + 1,2,5- + 1,3,6- + 2,3,6\text{-TrMN})}$$

Tetramethylnaphthalene (TeMN) Ratio

$$\frac{(1,3,6,7- + 2,3,6,7\text{-TrMN})}{(1,2,4,7- + 1,2,5,6- + 1,3,6,7- + 2,3,6,7\text{-TrMN})}$$

Dimethylphenanthrene (DMP) Ratio

$$\frac{(2,6- + 2,7\text{-DMP})}{(1,7- + 2,6- + 2,7\text{-DMP})}$$

They are all expressed in the format:

$$\frac{\text{"stable compounds"}}{\text{"labile + stable"}}$$

so that the ratios increase with maturity and all can only vary between 0 and 1.

To determine a more solid maturity ranking than could be done with R_c alone, linear

regressions were computed for the two phenanthrene ratios ($r^2 = 0.81$) and for the two naphthalene ratios ($r^2 = 0.79$), which indicated good correlations. The resulting best fit lines were rotated to form new "X" axes (X'_{phen} and X'_{naph}), the values for which were in turn also compared by linear regression ($r^2 = 0.64$). The best fit line for this regression was also rotated and the resulting X" values were used as the basis for an objective maturity ranking of the samples. Figure 12 displays the four maturity ratios in their original form, with the samples arrayed according to the X" ranking, along with vitrinite reflectance data for the rock samples. Although noise is present in each case, each curve progressively increases and thus the overall maturation ranking is clear.

Among rock samples, those with shallower burial depths (2-4 km, Cretaceous Logan Canyon and Missisauga Formations) are evidently the least mature according to the aromatic maturity parameters, placing them in the early stages of the oil window. Aromatic data indicate that the Logan Canyon sample from the N. Triumph G-43 well (3.7 km) and the Missisauga sample from the Thebaud well (3.9 km) have higher maturities than suggested by the vitrinite reflectances, which may be due to suppression of the latter. Phenanthrene ratios indicate a lower maturity than the naphthalene for two relatively shallow Jurassic rocks, the Baccaro Member sample from the Demascota well (3.6 km) and the Penobscot Limestone sample from the Penobscot well (4.2 km), probably due to facies effects. High maturity rocks (in the late oil window) include the deep (6.1 km) Mic Mac sample from the S. Venture well and the 5.1 km Missisauga sample from Venture B-52. The deep Missisauga samples from W. Chebucto (5.2 km) and N. Triumph G-43 (4.8 km) also exhibit relatively high maturities, although some facies effects may be operating on these marine samples.

The oil samples show considerable variation in maturity, according to the aromatic parameters, even though most of them have low specific gravities and are classified as condensates. Oils with maturities apparently in the early oil window include those from the Cohasset wells out of the shallow (1.9 km) Logan Canyon reservoir, the oils produced at shallow depths (1.9 to 2.6 km) by the Balmoral and Panuke wells, as well as the moderately deep oil (3.8 km) from the Missisauga in Citnalta. High maturity (late oil window) petroleum include those produced from deep (4.6 to 5.2 km) Mic Mac reservoirs in the Bluenose, Uniacke and Venture H-22 wells, as well as deep (5.0 to 5.2 km) reservoirs in the Missisauga from Arcadia J-16 and

S. Venture O-59. The remaining oil and rock samples are interpreted to have intermediate maturity levels (Fig. 12).

There is an overall trend indicating depth and stratigraphic control of the maturity of the rock extracts and the petroleums, with samples from deeper and/or older strata exhibiting higher levels of maturation.

Oil-Oil and Oil-Source Rock Correlation

A subset of 36 peaks from the aromatic GC/MS data was chosen for correlation purposes by the elimination of peaks strongly susceptible to maturation and of those peaks which showed little variation under any circumstances, as described in the Methods section. The data had been normalized according to retention time and scaled logarithmically prior to the filtering process. There may be some residual maturation effects in the correlation subset, but they should have been minimized.

Upon examination of the matrix of coefficients of determination (Table 7), it can be seen, for example, that in the column corresponding to the Chebucto oil sample, the quality of matches varies considerably. There are several very close matches ($r^2 > 0.95$), namely those with the N. Triumph B-52, S. Venture O-59 (DST 10) and Glenelg oils. It is extremely likely that these oils are genetically related. Further perusal of Table 6 leads to the recognition of close cross-matching among the oils listed above, as well as additional close relations with other oils and several rock extracts. Several additional smaller groupings also become apparent. Many other samples appear quite unrelated. To present this graphically in a concise and quantitative manner, a cluster analysis was performed (as described in the Methods section) and the results drawn as a dendrogram, scaled to r^2 values (Fig. 13).

The dendrogram confirms that the 4 oils just discussed form a tightly knit group, with all linkages above 0.95. Two other tight clusters, one comprised of the Olympia, Thebaud and Penobscot oils, the other of the Bluenose, Arcadia and two Venture H-22 oils, are also apparent. The latter group is characterized by higher maturity (Fig. 12), which may indicate that maturity effects on the correlation, while dampened considerably, were not entirely eliminated. All three groups are linked to each other with r^2 values > 0.9 . The Citnalta and Banquereau oils are closely related to each other, but are linked somewhat more distantly by a branching at $r^2 = 0.85$ to the main group of oils, which is nevertheless still a reasonable match. All these oils, which can be

collectively termed "Group A", are produced from Missisauga or Mic Mac reservoirs. The S. Venture (DST 5) and Uniacke petroleums show more a more limited similarity with the main group, perhaps in part the effect of their elevated maturity.

The Cohasset oils, produced out of the Cree Member of the Logan Canyon Formation, correlate well with one another, but poorly with the remaining samples. The one possible exception is the Sable Island oil from the Mamora Member of the same formation, with which the Cohasset oils have a limited affinity. The Panuke oils bear little resemblance to each other or to any of the other samples.

Among the rock extracts, a number of close correlations are also apparent. The Jurassic Abenaki Formation (Baccaro Member) sample from the Demascota well matches the Jurassic Penobscot Limestone sample (Mic Mac Formation, Penobscot L-30). These samples in turn show similarity in organic facies to the Lower Missisauga extracts of Abenaki J-56 and Migrant N-20 and the Logan Canyon samples of N. Triumph G-43 and Cohasset A-52 (Fig. 13). Lower and Middle Missisauga rocks from Venture B-52, Thebaud, S. Desbarres, and S. Sable exhibit a high degree of correlation to one another ($r^2 \leq 0.9$) and match the first cluster of rocks with a coefficient of determination of 0.86. This group of 10 extracts can be referred to as "Cluster 1". The second large cluster of rock samples has the Missisauga extracts from N. Triumph G-43 and W. Chebucto as its core, with the Missisauga extract from Whycocomagh and Verrill Canyon extracts from the Alma well as associated samples. This "Cluster 2" shows no affinity with the rocks of Cluster 1 (Fig. 13). The two remaining rock samples (Mic Mac from S. Venture and Misaine Member of the Abenaki Formation from Cohasset D-42) show no apparent relation to each other or to the other rock samples. As a whole, the rock extracts demonstrate that major differences in organic facies are represented in the Scotian Shelf samples.

The dendrogram shows a fair correlation ($r^2 = 0.75$) between the Group A petroleums and Cluster 1 rock extracts. The Cohasset oils correlate poorly with all rock samples. Cluster 2 rocks show no apparent similarity with any of the oils. Since the Group A oils show a high degree of genetic similarity to each other and account for the majority of oils analyzed for this study, it may be useful to consider them collectively, by taking the average of their correlation data sets. Then they may be correlated as a group to each rock sample, the results of which can be displayed in stratigraphic context (Fig. 14a). These results are in essential agreement with those

displayed in the dendrogram, but offer a clearer assessment of the quality of the petroleum-to-rock correlation. The Group A oils are still shown in Fig. 14a to match reasonably with the Cluster 1 extracts, with the added refinement that the lower and middle members of the Missisauga appear to provide the best candidate source rocks, along with the Penobscot Limestone and Cree Member samples. Cluster 2 extracts, whether from the Verrill Canyon or Missisauga, continue to show no correlation with Group A oils.

Since the presence of organic sulfur in fossil fuel samples is strongly related to the depositional environment of the original organic matter, it is interesting to examine sulfur content in the samples. A relatively high sulfur content is one of the hallmarks of marine Type II organic matter, whereas lacustrine and terrestrially-derived kerogens (Types I and III) are characteristically low in sulfur. Several of the sulfur compounds (dimethyldibenzothiophenes, Table 6) were important in the cluster analysis, but they merit special consideration. A simple ratio ("Th/Ph") of the sum of the C₂-alkyldibenzothiophenes (m/z 212) to the sum of the C₂-alkylphenanthrenes (m/z 206) is effective in presenting a semi-quantitative view of the sulfur content for the aromatic fraction. If the values of this ratio for the extracts are displayed in stratigraphic order (Fig. 14b), it is apparent that the low sulfur extracts (Th/Ph = 0.06 to 0.25) are those which correlate most closely to Group A oils (Fig. 14a), which have a mean Th/Ph ratio of 0.08. The non-Group A oils in the Scotian Shelf sample set also have low sulfur contents, but they are distinguished by elevated concentrations of another aromatic heterocyclic species, methyldibenzofurans.

It is reasonable to ascribe the highly thiophenic extracts (Cluster 2 and unclassified rocks) to marine origins and the low sulfur Cluster 1 extracts (which provide acceptable matches to the Group A petroleums) to non-marine. By extension, Group A oils are also likely to have non-marine origins. These designations are in broad agreement with the organic petrologic and stratigraphic data.

BASIN MODELING

Similar to modeling of eight other wells (Mukhopadhyay, 1993), a rifting steady state heat flow with $\beta = 3.0$ and present day heat flow between 35 (Abenaki J-56) and 40 (W. Chebucto K-20) mW/m² were chosen as default values for modeling. The model maturity profiles correlate

well with the measured maturity data for those two wells. However, use of rifting heat flow suggests that present day heat flow is the lowest.

Tables 8 and 9 show the following data used in modelling the Abenaki and Chebucto wells: formation name, beginning age, top depth, thickness, approximate lithology, kerogen type and approximate TOC (wt %). The lithologic mixing illustrates the proportions of sandstone/shale/limestone except for the Argo Formation where salt replaces limestone.

Figures 15a through 15j and 16a through 17l for the Abenaki J-56 and W. Chebucto K-20 wells, respectively, demonstrate the burial history, maturity and thermal conductivity through time, cumulative total HC and gas generation and migration through time and maturity or temperature or HC generation of individual source rocks through time.

The lithologic column within the burial history curves in Figures 15a and 16a does not include all formation names; the model names only those formations which are thickest. Figures 15a and 16a show that the maturation zones (example 0.4 to 0.7% R_o) cut across the stratigraphic boundaries between 200 and 87 my whereas, from 75-0 my, the zones run more or less parallel with the stratigraphic boundaries. Depressed maturation zones within the geohistory plots (15a and 16 a) can be seen when sediments were deposited in deeper paleowater depth. This downward shift of maturation zones are an artifact of the BASIN-MOD program.

Figures 15b and 16b show the maturation boundaries limiting the various formations and individual source rocks. The zones are divided into: (a) 0.4/0.5 to 0.7% R_o or early mature oil zone, (b) 0.7 to 1.0% R_o or mid mature oil zone, (c) 1.0 to 1.3% R_o , which is either late mature oil zone or early mature gas zone, (d) 1.3 to 2.0/2.6% R_o or main gas generation zone. Depth for each maturity zone for those two wells are as follows:

TABLE A

Well No.	Zone <u>a</u>		Zone <u>b</u>	Zone <u>c</u>	Zone <u>d</u>
(m)	(m)	(m)	(m)		
Abenaki J-56	2300-3750		3750-4650	4650-5300	5300-6450
W. Chebucto K-20	2000-4100		4100-5200	5200-5900	5900-7100

The cumulative HC generation throughout the sequence (up to 8000 m) varies

with the nature and thickness of the source rocks. In comparison, W. Chebucto K-20 generated much higher HC's compared to Abenaki J-56 well. Comparable cumulative HC generation and expulsion efficiency from those two wells up to 8000m are as follows:

TABLE D

Well No.	Cumm. Oil (bbl HC/Ac.ft rock)	Cumm. Gas (bbl HC/ac.ft rock)	Depth (m)	Expulsion Efficiency (bbl HC/ac.ft rock)
Abenaki J-56 (Fig.15d)	100	250	4000-8000	50 oil, 250 gas
W. Chebucto K-20 (Fig. 16d)	400	1250	4000-8000	250 oil, 1000 gas

Cumm. oil = Cumulative oil; Cumm. Gas = Cumulative Gas

Comparing the history of maturation, temperature and transformation ratio through time and using the measured kinetics, the timing and amount of hydrocarbon generation and migration of selected rock intervals from those two wells were evaluated.

The Abenaki I unit (Kerogen Type IIL) from the Abenaki J-56 well, lies within the maturation zone *a* and is subjected to a maximum of 100°C temperature (Fig. 15f). A very low amount (<10bbl/acre foot) of liquid hydrocarbons were generated. The migration of hydrocarbons from the Abenaki unit is negligible (Fig. 15g). A lower Callovian source rock from the same well which lies within maturation zone *b* and is subjected to a maximum temperature of 110°C (Fig. 15h), has a low transformation ratio (0.4; Fig. 15i). This unit however, expelled about 50 bbl (equivalent) of gas per acre foot of rock (Fig. 15i). The next lower unit, Scatarie Member (Kerogen Type IIB), from Abenaki J-56 well, also generated mainly gas (Fig. 15j).

On the other hand, the Mississauga 2 source rock unit (kerogen Type IIA-IIB) in W. Chebucto K-20, which also lies within maturation zone *b* and is subjected to a maximum temperature of 120°C (Fig. 16f), has a transformation ratio of 0.98 (Fig. 16g). This unit has generated and migrated abundant liquid hydrocarbons (500 bbl/acre foot) starting 125 mybp (Fig. 16h); it has a much lower gas generation. The next lower unit, Mississauga 3, which has Kerogen Type IIB and is subjected to a maximum temperature of 165°C (Fig. 16i), started HC generation 137 mybp. It generated and migrated more gas than liquid hydrocarbons (16j). The next lower unit, Top Jurassic, which has a higher proportion of shale than the other two, has mainly Kerogen

Type III and lies within maturation zone *d*. It is subjected to a maximum temperature of 185°C (Fig. 16k) and has generated and migrated 900 bbl (equiv) of gas/acre foot since 137 mybp.

SUMMARY AND CONCLUSIONS

The report illustrated the results and significance of: (a) hydrocarbon potential and maturation of 50 selected limestone and calcareous shale samples (mainly from the Abenaki Formation) and 50 selected composite shale units from various stratigraphic intervals, (b) kinetics analyses of four selected source rocks, (c) maturation of liquid hydrocarbons, oil-oil and oil-source rock correlation of 39 selected oil, condensates, and source rock extracts by using aromatic biomarkers, and (d) 1-D basin modeling of two selected wells (carbonate-rich Abenaki J-56 and clastic-rich W. Chebucto K-20) using the BASIN-MOD program. The report added several aspects of importance towards the understanding of source rock types, their maturation, kinetic properties, oil-oil and oil-source rock correlation, and hydrocarbon generation and expulsion in the Scotian Basin which point toward the following summary and conclusions:

Source Rock Characterization

Based on organic facies, which uses maceral composition, fluorescence and oxidation criteria of macerals and minerals, hydrogen and oxygen indices (from Rock-Eval pyrolysis) and vitrinite reflectance, the hydrocarbon potential and maturity of selected limestone and composite shale source rocks was evaluated and compared. Volumetrically, the comparable source rock potential between limestone and composite shale samples can be illustrated as follows: Limestone and calcareous shale: 6% are oil- and condensate-prone Kerogen Type IIA-IIB; 42% are condensate- and gas-prone Kerogen Type IIB; 34% are gas-prone Kerogen Type III; and 18% are nonsource Kerogen Type IV. Composite shale: 20% are oil- and condensate-prone Kerogen Type IIA-IIB; 40% condensate- and gas-prone Kerogen Type IIB; 40% are gas-prone Kerogen Type III; and none of them are considered as nonsource for hydrocarbons.

Kinetics of Selected Source Rocks

The distribution of Activation Energies and Arrhenius Constant suggests that Kerogen Type IIA-IIB (peak at 44 to 48 kcal/mole) can generate and expel mainly oil (350 mg HC/g TOC) earlier in its maturation history (<0.8% R_o) than Kerogen Type IIB (peak around 50 and 55 kcal/mole) and III (46-50 kcal/mole). Kinetics of two limestone and calcareous shale source

rocks show sharp peaks between 50-52 kcal/mole and the reaction rate is slower. The reaction rate of Kerogen Type IIA-IIB is considered as medium-fast and compared to Type IIC source rock of Hunt et al. (1991), whereas limestone source rocks can be compared to Kerogen Type I of the Green River Shale.

Maturation of Crude Oil/Condensate and Oil-Oil or Oil-Source Rock Correlation

The majority of petroleums and rock extracts in this study are at advanced levels of thermal maturity (middle to late oil window), as evidenced by the distribution of methylphenanthrene, dimethylphenanthrene, trimethylnaphthalene and tetramethylnaphthalene isomers. This is not surprising, given the great burial depth (4 to 6 km) of many of the samples.

Multiple linear regression and subsequent cluster analysis reveals a closely related family of oils ("Group A") including those from N. Triumph B-52, S. Venture O-59 (DST 10), Glenelg, Olympia, Thebaud, Penobscot, Bluenose, Arcadia, Venture H-22, Citnalta and Banquereau, all produced from Missisauga or Mic Mac reservoirs. The Cohasset oils form a second group and the Panuke oils do not seem to be related to each other or to any other oil.

A group of rock extracts ("Cluster 1") from the lower and middle members of the Missisauga, along with the Penobscot Limestone and Cree Member samples correlate well with each other, indicating similar organic facies. These samples also have reasonable affinities for the Group A oils, with the Lower Missisauga samples from the Abenaki and S. Desbarres wells appearing to be the best candidate source rocks. A second cluster of extracts, including samples from the Verrill Canyon and Missisauga Formations are related to one another, but correlate neither with Group A oils nor Cluster 1 extracts.

To supplement the cluster analysis, a parameter related to the sulfur content of the samples (the ratio of the sum of C₂-alkyldibenzothiophenes to the sum of the C₂-alkylphenanthrenes) was established. It supports the statistical findings, further indicating that the oils and Cluster 1 extracts have a similarly low thiophenic sulfur content, in contrast to the relatively high sulfur Cluster 2 and unclassified extracts. Samples richer in sulfur derive from marine rocks, while the remainder are apparently of non-marine origin, a conclusion supported by the organic petrologic and stratigraphic data.

Basin Modeling

Burial history curves of two selected wells (Abenaki J-56 and W. Chebucto K-20) suggest

that the sedimentation rate between 200 and 75 mybp (syn- and early post-rift sediments) was relatively high and relatively low after that (late post-rift phase).

The best-fit maturation line correlating vitrinite reflectance and depth, is primarily controlled by small changes in present day heat flow (mW/m^2) values within a narrow range of lithospheric stretching factor, β .

Due to the presence of fewer oil-prone source rocks and slower reaction rates of the limestone source rock units, W. Chebucto K-20 source rocks (mainly clastic units) generated and migrated cumulative hydrocarbons five times more than Abenaki J-56 source rocks. Mississauga 2 unit (kerogen Type IIA-IIB) in W. Chebucto K-20 generated and migrated more oil (500 bbl/acre foot) than gas, whereas Mississauga 3 (kerogen Type IIB) migrated major gas (900 bbl equiv/acre foot) between 125-50 mybp. Modelling suggests that major hydrocarbons (about 80%) were generated and expelled between 137 and 50 my before present. Most source rocks also reached their peak temperature during that time; temperature has decreased since then. This suggests that non-expelled oils within the source rocks have been cracked to gas and residue.

ACKNOWLEDGMENTS

The author acknowledges John A. Wade of Basin Analysis Subdivision, Atlantic Geoscience Centre, Bedford Institute of Oceanography, Dartmouth, Nova Scotia for various fruitful discussions during the period of this contract. The author acknowledges M. P. Avery and B. Crilley of Atlantic Geoscience Centre for kerogen isolation, kerogen smear slides, kerogen plugs, vitrinite reflectance data. The author acknowledges Steven Bigelow of CNSOPB for their permission to collect the samples and Karl Makrides of CNSOPB Core Storage Facility at Dartmouth, N. S. for his assistance during sample collection. The author acknowledges Dr. Michael A. Kruege of Southern Illinois University, Carbondale, Illinois and Dan Jarvie, Humble Geochemical Services, Humble, Texas, USA for the Liquid chromatography/aromatic gc-ms and kinetic data and report. The author acknowledges the help of M. G. Fowler of ISPG, Geological Survey of Canada, Calgary, Alberta and Ross Stewart of Calgary, Alberta for their Rock-Eval pyrolysis data.

REFERENCES

- Alexander R., Kagi R. I. and Sheppard P. N. (1983) Relative abundance of dimethylnaphthalene isomers in crude oils. *Jour. Chrom.* **267**:367-372.
- Alexander R., Cumbers K. M. and Kagi R. I. (1986) Alkylbiphenyls in ancient sediments and petroleums. In Leythaeuser D. and Rullkötter J., eds., *Advances in Organic Geochemistry 1985*, p. 841-845.
- Budzinski H., Garrigues P., Connan J., Lee M.L., Andersson J. and Bellocq J. (1991) Determination of maturity indicators in alkylated aromatic series by gas chromatography-mass spectrometry (GC-MS). In Manning D.A.C., ed., *Organic Geochemistry: Advances and Applications in the Natural Environment*, Manchester University Press, Manchester, p. 619-623.
- Budzinski, H., Garrigues, P., Radke, M., Connan, J., Rayez, J. C., Rayez, M. T. (1993a), Use of molecular modeling as a tool to evaluate thermodynamic stability of alkylated polycyclic aromatic hydrocarbons, *Energy & Fuels* **7**:505-511.
- Budzinski, H.; Garrigues, P.; Radke, M.; Connan, J.; Oudin, J. L. (1993b) Thermodynamic calculations on alkylated phenanthrenes, *Org. Geochem.*, **20**:917-926.
- Budzinski H., Garrigues P., Connan J., and Bellocq J. (1993c) Chromatographic fractionation of aromatic compounds from ancient and recent sediments for access to alkylated PAH by GC-FID and GC-MS. *Química Analítica* **12**:69-74.
- Espitalie, J., Deroo, Cr., and Marquis, F. (1985) Rock-Eval pyrolysis and its applications. *Rep. Inst. Fr. Petrol.* **33878**, 72 p.
- Garrigues P., De Sury R., Angelin M.L., Bellocq J., Oudin J.L. and Ewald M. (1988) Relation of methylated aromatic hydrocarbon distribution pattern to the maturity of organic matter in ancient sediments from the Mahakam delta. *Geochim. Cosmochim. Acta*, **52**:375-384.
- Hunt, J. M., Lewan, M. D. and R. J.-C. Hennet, 1991, Modelling oil generation with Time-Temperature Index graphs based on the Arrhenius Equation. *Am. Assoc. Petrol. Geol. Bull.* v. 75. no. 4, p. 795-807.
- Hutton, A. (1987) Petrographic classification of oil shales. *Int. J. Coal Geol.*, **8**:203-231.
- Jones, R. W. (1987) Application of organic facies for hydrocarbon potential. In *Advances in Petroleum Geochemistry* (Edited by Brooks, J.), **2**, 1-90.
- Kruger, M. A. (1993) Oil-source rock correlation and maturity determination using aromatic molecular markers offshore Nova Scotia. Report of the Department of Geology, Southern Illinois University at Carbondale, Illinois. Prepared for Global Geoenergy Research Ltd., Halifax, Nova Scotia.

Kruger, M. A. (1992) Comparison of the flash pyrolysis products of oil asphaltenes, source rocks, and source rock extract asphaltenes, Scotian Shelf, Nova Scotia. Report of the Department of Geology, Southern Illinois University, Carbondale. 9p, 8 Figs. Prepared for Global Geoenergy Research Ltd. Halifax, Nova Scotia.

Kruger, M. A. (1991) Oil-source rock correlation and maturity determination using aromatic molecular markers, offshore Nova Scotia. Report of the Department of Geology, Southern Illinois University. 14p, 12 Figs. Prepared for Global Geoenergy Research Ltd. Halifax, Nova Scotia.

Kruger, M. A. (1990) Aromatic molecular markers in petroleum and rock extracts, offshore Nova Scotia. Report of the Department of Geology, Southern Illinois University, Carbondale. 14p, 14 Figs. Prepared for Global Geoenergy Research Ltd., Halifax, Nova Scotia.

Kruger M. A., Bensley D. F. and Crelling J. C. (1989) An organic geochemical and petrologic study of a suite of sporinite-rich coals from the Lower Kittanning Seam, Upper Carboniferous, USA. International Conference on Coal Formation, Occurrence and Related Properties, Orléans, 12-15 Sept. 1989, Abstract no. 29.

Kruger M. A., Bensley D. F. and Adams J. P. (1990a) Comparison of the behavior of molecular maturity parameters in coal and marine shale extracts. 199th American Chemical Society National Meeting, Boston, 22-27 Apr. 1990, Div. of Geochemistry Abstract no. 10.

Kruger M. A., Crelling J. C., Pascone J. M., Palmer S. R. and Hippo E. J. (1990b) Confined pyrolysis of Illinois Basin coals and maceral concentrates. 200th American Chemical Society National Meeting, Washington, 26-31 Aug. 1990, Div. of Geochemistry Abstract no. 43.

Kruger M. A. and Landais P. (1992) Artificial maturation of coal and maceral concentrates: Saturate and polyaromatic molecular markers. *Preprints of Papers Presented at the 204th ACS National Meeting*; Amer. Chem. Soc. Div. of Fuel Chemistry: Washington, DC, Vol. 37, No. 4, pp. 1595-1600.

Kruger, M. A. and Bensley, D. F. (1993) Flash pyrolysis-gas chromatography/mass spectrometry of Lower Kittanning vitrinites: Changes in the distributions of polyaromatic hydrocarbons as a function of coal rank. Amer. Chem. Soc. Div. of Geochemistry, Abstracts for the Annual Meeting, August 1993, Chicago.

Later D. W., Lee M. L., Bartle K. D., Kong R. C. and Vassilaros D. L. (1981) Chemical class separation and characterization of organic compounds in synthetic fuels. *Anal. Chem.*, **53**:1612-1620.

Mackenzie A.S., Disko U. and Rullkötter J. (1983) Determination of hydrocarbon distributions in oils and sediment extracts by gas chromatography-high resolution mass spectrometry. *Org. Geochem.* **5**:57-63.

Massart D. L. and Kaufman L. (1983) *The Interpretation of Analytical Chemical Data by the Use of Cluster Analysis*. John Wiley and Sons, New York, 237 p.

Mukhopadhyay, P. K. (1993), Analyses and interpretation of geochemical and source rock data from Scotian Shelf wells. SSC Report. File No. OSC92-00217-(004). Contract No. 23420-2-C026/01-OSC. 33p. 10 tables. 167 figures.

Mukhopadhyay, P. K. (1991). Characterization and maturation of selected Cretaceous and Jurassic source rocks and crude oil, Scotian Shelf. SSC Report. File No. OSC91-00026-(004), Contract No. 23420-1-C012/01-OSC, 29p, 8 tables, 16 figures.

Mukhopadhyay, P. K. (1990a). Characterization and maturation of selected oil and condensate samples and correlation with source beds, Scotian Shelf. SSC Report No. OSC90-00245-(014). Contract No. 23420-0-c070/01-OSC.

Mukhopadhyay, P. K. (1990b) Evaluation of organic facies of the Verrill Canyon Formation, SSC Report No. OSC89-00482-(010). Contract No. 23420-9-C673/01-OSC.

Mukhopadhyay, P. K. and Wade, J. A. (1990). Organic facies and maturation of sediments from three Scotian Shelf wells. *Bull. Canadian Petrol. Geol.* v. 38, no. 4. p. 407-425.

Mukhopadhyay, P. K. (1989a) Characterization of amorphous and other organic matter types by microscopy and pyrolysis-gas chromatography. *Org. Geochem.* **14** (3), 269-284.

Mukhopadhyay, P. K. (1989b) Cretaceous organic facies and oil occurrence, Scotian Shelf. SSC Report No. OSC89-00136-(008), Contract No. 23420-9-C559/01-OSC.

Radke M. (1987) Organic geochemistry of aromatic hydrocarbons. In Brooks J. and Welte D. H., eds., *Advances in Petroleum Geochemistry Vol. 2*, p 141-207, Academic Press, London.

Palacas, J. 1984. *Carbonate Source Rocks*. Ame. Assoc. Pet. Geol. Series in Geology. AAPG Publication, Tulsa.

Radke M. (1988) Application of aromatic compounds as maturity indicators in source rocks and crude oils. *Mar. Petrol. Geol.*, **5**:224-236.

Radke M., Schaeffer R. G., Leythaeuser D. and Teichmüller M. (1980) Composition of soluble organic matter in coals: relation to rank and liptinite fluorescence. *Geochim. Cosmochim. Acta*, **44**:1787-1800.

Radke M., Welte D. H. and Willsch H. (1982a) Geochemical study on a well in the Western Canada Basin: relation of the aromatic distribution pattern to maturity of organic matter. *Geochim. Cosmochim. Acta*, **46**:1-10.

Radke M., Willsch H. and Leythaeuser D. (1982b) Aromatic components of coal: relation of

distribution pattern to rank. *Geochim. Cosmochim. Acta*, **46**:1831-1848.

Radke M. and Welte D. H. (1983) The Methylphenanthrene Index (MPI): A maturity parameter based on aromatic hydrocarbons. In Bjørroy M. *et al.*, eds., *Advances in Organic Geochemistry 1981*, p. 504-512.

Radke M., Leythaeuser D. and Teichmüller M. (1984a) Relationship between rank and composition of aromatic hydrocarbons for coals of different origins. *Org. Geochem.* **6**:423-430.

Radke M., Willsch H. and Welte D. H. (1984b) Class separation of aromatic compounds in rock extracts and fossil fuels by liquid chromatography. *Anal. Chem.* **56**:2538-2546.

Radke M., Welte D. H. and Willsch H. (1986) Maturity parameters based on aromatic hydrocarbons: Influence of the organic matter type. In Leythaeuser D. and Rullkötter J., eds., *Advances in Organic Geochemistry 1985*, p. 51-63.

Rowland S. J., Alexander R. and Kagi R. I. (1984) Analysis of trimethylnaphthalenes in petroleum by capillary gas chromatography. *Jour. Chrom.*, **249**:407-412.

Senftle, J. T., Brown, J. H., and Larter, S. R. (1987) Refinement of organic petrographic methods for kerogen characterization. *Int. Jour. Coal Geol.* **7**, 105-118.

Teichmüller, M. (1986) Organic petrology of source rocks, history and state of the art. *Org. Geochem.* **10** (1-3), 581-599.

Tissot, B. and Welte, D. H. (1984) *Petroleum Formation and Occurrence*. Springer-Verlag, Berlin, 699 p.

Villar H. J., Püttmann W. and Wolf M. (1988) Organic geochemistry and petrography of Tertiary coals and carbonaceous shales from Argentina. In Mattavelli L. and Novelli L., eds., *Advances in Organic Geochemistry 1987*, pp. 1011-1021.

Wang T. G. and Simoneit B. R. T. (1990) Organic geochemistry and coal petrology of Tertiary brown coal in the Zhoujing mine, Baise Basin, South China. 3. Characteristics of polycyclic aromatic hydrocarbons. *Fuel* **69**:12-20.

White C. M. and Lee M. L. (1980) Identification and geochemical significance of some aromatic components of coal. *Geochim. Cosmochim. Acta* **44**:1825-1832.

Whitehurst D. D., Buttrill S. E., Derbyshire F. J., Farcasiu M., Odoerfer G. A. and Rudnick L. R. (1982) New characterization techniques for coal-derived liquids. *Fuel* **61**:994-1005.

LIST OF TABLE AND FIGURES

Tables

- 1A. Rock-Eval pyrolysis data of various analyzed limestone and calcareous shale samples along with depth and stratigraphic position of the samples.
- 1B. Rock-Eval pyrolysis data of various composite shale samples along with depth and their stratigraphic position of the samples.
- 2A. Maceral composition in volume percent (organic facies), mean random vitrinite reflectance (% R_o), kerogen type, and oil/gas potential of limestone and calcareous shale samples.
- 2B. Maceral composition in volume percent (organic facies), mean random vitrinite reflectance (% R_o), kerogen type, and oil/gas potential of composite shale samples.
3. Activation energy distributions and Arrhenius factor of four samples (two limestones and two composite shales) based on Rock-Eval 5 derived kinetic parameters. Kinetic data of Green River Shale (Kerogen Type I) was also included for comparison (after Jarvie, 1994).
4. List of source rock and light oil/condensate samples used for aromatic biomarker analysis by GC-MS and for oil-oil or oil-source rock correlation. The asterix denotes the samples which are analyzed in the 1993-1994 contract (after Kruge, 1993).
5. Liquid chromatogray data of three source rock and four light oil/condensate samples analyzed in the 1993-1994 contract (after Kruge, 1993).
6. List of aromatic biomarker peaks and their fragment ions recognized in GC-MS analysis (after Kruge, 1993).
7. Coefficients of determination from multiple linear regression of Scotian Shelf parameters. For details, see text in the oil-oil and oil-source rock correlation chapter (after Kruge, 1993).
8. Lithostratigraphy, kerogen type, and total organic carbon data from Abenaki J-56 well which are used in the Basin-Mod program.
9. Lithostratigraphy and kerogen type data from W. Chebucto K-20 which are used in the Basin-Mod program.

Figures

- 1A. A plot of HI vs. OI from Rock-Eval pyrolysis showing position of various limestone and calcareous shale whole rock samples.
- 1B. A plot of HI vs. OI from Rock-Eval pyrolysis showing position of various limestone and

calcareous shale kerogen samples.

1C. A plot of HI vs. OI from Rock-Eval pyrolysis showing position of various composite shale whole rock samples.

2A. A plot of HI vs. T_{\max} (°C) from Rock-Eval pyrolysis showing position of various limestone and calcareous shale whole rock samples.

2B. A plot of HI vs. T_{\max} (°C) from Rock-Eval pyrolysis showing position of various limestone and calcareous shale kerogen samples.

2C. A plot of HI vs. T_{\max} (°C) from Rock-Eval pyrolysis showing position of various composite shale whole rock samples.

3A. A graph showing activation energy distribution vs. percent of reaction for Green River Shale. The Arrhenous Factor is also given (after Jarvie, 1994).

3B. A graph showing activation energy distribution vs. percent of reaction for calcareous shale sample from Abenaki J-56, 10090-10130'. The Arrhenous Factor is also given (after Jarvie, 1994).

3C. A graph showing activation energy distribution vs. percent of reaction for limestone sample from Demoscota G-43, 11840-11880'. The Arrhenous Factor is also given (after Jarvie, 1994).

3D. A graph showing activation energy distribution vs. percent of reaction for composite shale sample from N. Triumph B-52, 3560- 3715 m. The Arrhenous Factor is also given (after Jarvie, 1994).

3E. A graph showing activation energy distribution vs. percent of reaction for composite shale sample from Venture B-52, 3580-3740 m. The Arrhenous Factor is also given (after Jarvie, 1994).

4. Total ion chromatograms of selected ion data of the aromatic fraction of the Lower Missisauga extract from Migrant N-20. Peaks labeled as in Table 6. Phenanthrene (peak PHN) is twice as tall as shown. Some peak labels are omitted for clarity (after Kruge, 1993).

5. Mass chromatograms of alkylnaphthalene isomers of the Lower Missisauga extract from Migrant N-20. Peaks used in the study are labeled as in Table 6 and molecular structures are shown (after Kruge, 1993).

6. Partial m/z 198, 212 and 182 mass chromatograms of the Lower Missisauga extract from Migrant N-20. Peaks are labeled as in Table 6 and molecular structures are shown (Kruge, 1993).

7. Partial m/z 178, 192 and 206 mass chromatograms of the Lower Missisauga extract from

Migrant N-20, with 178 and 192 combined. Peaks are labeled as in Table 6 and molecular structures are shown (after Kruge, 1993).

8. Partial m/z 202, 216 and 230 mass chromatograms of the Lower Missisauga extract from Migrant N-20, with 202 and 216 combined. Peaks are labeled as in Table 6 and examples of molecular structures are shown (after Kruge, 1993).

9. Partial m/z 228, 242 and 252 mass chromatograms of the Lower Missisauga extract from Migrant N-20. Masses 228 and 242 are combined. Peaks are labeled as in Table 6 and examples of molecular structures are shown (Kruge, 1993).

10. Comparison of normalized quantitation results of a condensate and a rock extract, illustrating variations in distribution of classes of compounds as a function of sample type. The extract is distinguished by the predominance of the larger aromatic hydrocarbons (after, Kruge, 1993).

11. Cross-plots of the aromatic correlation sets of two petroleums with high ($r^2 = 0.99$) and two with low ($r^2 = 0.65$) degrees of similarity (after Kruge, 1993).

12. Maturity parameters based on calculated equivalent vitrinite reflectance (R_c), dimethylphenanthrene ratio (DMP), tetramethylnaphthalene ratio (TeMN), and trimethylnaphthalene ratio (TrMN). See text for definition of parameters and for explanation of the technique used to determine the sequence of samples presented. A semilog plot of measured vitrinite reflectance is also shown. The arrow marks an obviously suppressed sample. The full names for the samples are given in Table 5 and examples of the sample code are shown in Table 7 (after Kruge, 1993).

13. Dendrogram showing clustering based on a similarity matrix derived from multiple linear regression of aromatic data from Scotian Shelf samples. See text for further explanation. The full names for the samples are given in Table 4 and examples of the sample code are shown in Table 7 (after Kruge, 1993).

14. Summary diagram of the extent of correlation between candidate source rock extracts and Group A petroleums, based on a) multiple linear regression and b) the ratio of C_2 -alkyldibenzothiophenes to C_2 -alkylphenanthrenes (Th/Ph), which serves as a relative indicator of organic sulfur content. Group A oils have a mean Th/Ph ratio of 0.08. The full names for the samples are given in Table 4 and examples of the sample code are shown in Table 7. See text for further discussion (after Kruge, 1993).

15. Various diagrams which are derived from 1-D modeling of Abenaki J-56 well.

- a. Burial history plot of various formations with maturation boundaries.
- b. Depth versus reflectance plot showing best-fit line related to measured R_c .
- c. Depth versus thermal conductivity.
- d. Depth versus cumulative in situ hydrocarbons in bbls per acre foot of rock.
- e. Depth versus expelled gas in bbls equivalent per acre foot of rock.

- f. Change in temperature and maturity through time for Abenaki 1 unit.
- g. The figure shows two plots. The upper shows the heat flow through time and the lower shows cumulative hydrocarbons from Abenaki 1 unit in bbls per acre foot.
- h. Change in temperature and maturity through time for L. Callovian unit.
- i. Change in transformation ratio through time and the cumulative hydrocarbons through time for L. Callovian unit through time.
- j. Change in maturity through time and cumulative hydrocarbons through time in bbls per acre foot for Scatarie Member.

16. Various diagrams which are derived from 1-D modeling of W. Chebucto K-20 well.

- a. Burial history plot of various formations with maturation boundaries.
- b. Depth versus reflectance plot showing best-fit line related to measured R_o .
- c. Depth versus thermal conductivity.
- d. Depth versus cumulative in situ hydrocarbons in bbls per acre foot of rock.
- e. Depth versus expelled gas in bbls equivalent per acre foot of rock.
- f. Change in temperature and maturity through time for Mississauga 2 unit.
- g. Change in transformation ratio through time for Mississauga 2 and Naskapi units.
- h. The figure shows two plots. The upper shows the heat flow through time and the lower shows cumulative hydrocarbons from Mississauga 2 unit in bbls per acre foot.
- i. Change in temperature and maturity through time for Mississauga 3 unit.
- j. The figure shows two plots. The upper shows the heat flow through time and the lower shows cumulative hydrocarbons from Mississauga 3 unit in bbls per acre foot.
- k. Change in temperature and maturity through time for Top Jurassic unit.
- l. The figure shows two plots. The upper shows the heat flow through time and the lower shows cumulative hydrocarbons from Mississauga 3 unit in bbls per acre foot.

TABLE IA : Rock-Eval pyrolysis data of limestone and calcareous shale samples

Borehole No.	Depth (ft/m)	Formation/ Member	R _o	T _{max} (°C)	TOC (wt%)	S ₁	S ₂	S ₁ +S ₂	S ₃	PI	HI	OI
Abenaki J-56	10090-10130'(w) (k)	Baccaro	0.51	432	2.39	0.91	4.91	5.82	1.04	0.16	205	43
				429	74.15	23.7	206		0.10	282	27	
	10800-10810'(w) (k)	same	0.53	435	1.22	0.53	1.58	2.11	0.72	0.25	129	59
				431	78.9	45.3	174		25.1	0.21	220	31
	12040-12050'(w) (k)	same	0.57	435	0.90	0.32	0.89	1.21	0.62	0.26	98	68
				431	65.78	28.1	146		8.70	0.16	222	13
	12360-12370'(w) (k)	same	0.58	432	0.53	0.18	0.25	0.43	0.53	0.42	47	100
				432	90.62	40.7	262		16.0	0.13	288	17
	12890-12900'(w) (k)	same	0.57	435	1.05	0.38	0.79	1.17	0.62	0.32	75	59
	13040-13050 (w) (k)	same	0.60	433	0.63	0.47	0.45	0.92	0.57	0.51	71	90
				433	86.2	36.3	213		21.6	0.15	247	25
	13270-13280'(w) (k)	Misaine	0.78	438	0.91	0.10	0.40	0.50	0.44	0.20	43	48
				432	67.1	21.1	116		17.9	0.15	172	26
	13790-13800'(w) (k)	Scatarie	0.76	433	0.36	0.14	0.16	0.30	0.40	0.47	44	111
				441	83.04	32.0	179		34.0	0.15	215	40
	14080-14090'(w) (k)	Mohican	0.82	437	0.60	0.32	0.40	0.72	0.32	0.44	66	53
				444	13.35	4.68	15.7		6.59	0.23	117	49
	14140-14150'(w) (k)	same	0.89	440	0.65	0.48	0.43	0.91	0.43	0.53	66	66
	14220-14230'(w) (k)	same	0.94	439	0.85	0.52	0.64	1.16	0.50	0.45	75	58
				443	45.1	14.2	66.8		11.1	0.18	148	24
	14350-14360'(w) (k)	same	0.92	437	0.95	2.22	0.93	3.15	0.46	0.70	97	48
	14410-14435'(w) (k)	Iroquois	0.91	439	0.80	1.84	0.62	2.46	0.69	0.75	77	86

R_o = mean random vitrinite reflectance; T_{max} = maturation parameter from Rock-Eval pyrolysis in °C. TOC = total organic carbon in weight %.
 S₁ and S₂ = in mg HC/g TOC; S₃ = mg CO₂/g TOC; PI = production index = S₁/S₁+S₂; HI = hydrogen index in mg HC/g TOC; OI = oxygen index =
 mg CO₂/g TOC. (w) = whole rock; (k) kerogen concentrate.

TABLE 1A : Rock-Eval pyrolysis data of limestone and calcareous shale samples

Borehole No.	Depth (ft/m)	Formation/Member	R ₀	T _{max} (°C)	TOC (wt%)	S ₁	S ₂	S ₁ +S ₂	S ₃	PI	HI	OI
Cohasset D-42	10000-10030' (w)	L. Mississauga	0.64	437	0.58	0.15	0.47	0.62	0.63	0.24	81	108
	10830-10850' (w)	Baccaro	0.71	431	0.23	0.13	0.10	0.23	0.34	0.57	43	147
	13430-13450' (w)	same	0.63		0.06							
	13740-13760' (w)	same	1.23		0.07							
	14160-14180' (w)	Misaine	1.23		0.17	0.03	0.07	0.10	0.15	0.30	41	88
Cohasset L-97	14370-14390' (w)	same	1.33		0.16	0.03	0.05	0.08	0.21	0.38	31	131
	14490-14520' (w)	same	1.34	457	0.25	0.02	0.04	0.06	0.49	0.33	16	196
	3100-3110 m (w) (k)	L. Mississauga	0.63	440	0.83	0.12	0.50	0.62	0.44	0.19	60	53
	3195-3205 m (w)	Baccaro	0.62	428	29.82	79.6	66.4		28.0	0.55	222	93
	3990-4000 m (w) (k)	same		435	1.22	1.28	1.63	2.91	3.38	0.44	133	277
	4155-4165 m (w)	same			0.34	0.08	0.16	0.24	2.22	0.33	47	652
	4470-4480 m (w)	Misaine	1.41	455	9.37	11.7	20.0		23.2	0.37	213	247
	4545-4555 m (w)	same	1.50		0.13							
					0.18	0.01	0.03	0.04	0.32	0.25	16	177
					0.15							

R₀ = mean random vitrinite reflectance; T_{max} = maturation parameter from Rock-Eval pyrolysis in °C. TOC = total organic carbon in weight %.
 S₁ and S₂ = in mg HC/g TOC; S₃ = mg CO₂/g TOC; PI = production index = S₁/S₁+S₂; HI = hydrogen index in mg HC/g TOC; OI = oxygen index = mg CO₂/g TOC.

TABLE 1A : Rock-Eval pyrolysis data of limestone and calcareous shale samples

Borehole No.	Depth (ft/m)	Formation/Member	R _o	T _{max} (°C)	TOC (wt%)	S ₁	S ₂	S ₁ +S ₂	S ₃	PI	HI	OI
Cohasset L-97	4620-4630 m (w)	Scatarie	1.6		0.09							
	4855-4865 m (w)	Mohican	1.92		0.26							
	4870-4880 m (w)	same	1.82		0.18							
Demoscota G-43	11480-11510' (w) (k)	Verrill Canyon	0.73	438 440	1.00 74.02	0.39 33.8	0.77 145	1.17	0.73 7.02	0.34 0.19	77 195	73 9
	11840-11880' (w) (k)	Baccaro	0.72	440 423	0.94 74.1	0.38 18.2	0.76 85.3	1.14	0.59 49.1	0.33 0.18	80 115	62 66
	14550-14570' (w)	same		429	0.19	0.23	0.16	0.39	0.71	0.59	84	373
	14820-14840' (w)	same	0.75		0.10							
	15060-15080' (w)	Misaine	1.17		0.12							
	15250-15270' (w)	Scatarie	1.32		0.30							
Penobscot L-30	12280-12290' (w) (k)	Mic Mac	0.61	439 439	0.80 76.1	0.19 20.6	0.44 106	0.63	0.70 8.75	0.30 0.16	55 139	87 11
	12440-12450' (w) (k)	same	0.67	440 439	1.11 76.9	0.22 20.4	0.73 116	0.95	0.85 5.5	0.23 0.15	65 151	76 7
	12580-12600' (w) (k)	same	0.73	442 440	1.04 77.3	0.21 24.9	0.60 117	0.81	0.77 7.0	0.26 0.18	57 151	74 9
	12730-12740' (w) (k)	same	0.73	443 441	1.12 79.3	0.25 25.7	0.60 140	0.85	1.31 5.5	0.29 0.16	53 176	116 6

(w) = whole rock; (k) kerogen concentrate; R_o = mean random vitrinite reflectance; T_{max} = maturation parameter from Rock-Eval pyrolysis in °C. TOC = total organic carbon in weight %; S₁ and S₂ = in mg HC/g TOC; S₃ = mg CO₂/g TOC; PI = production index = S₁/S₁+S₂; HI = hydrogen index in mg HC/g TOC; OI = oxygen index = mg CO₂/g TOC.

TABLE IA : Rock-Eval pyrolysis data of limestone and calcareous shale samples

Borehole No.	Depth (ft/m)	Formation/Member	R ₀	T _{max} (°C)	TOC (wt%)	S ₁	S ₂	S ₁ +S ₂	S ₃	PI	HI	OI
Penobscot L-30	12980-12990' (w)	Mic Mac	0.77	441	1.13	0.23	0.83	1.06	0.74	0.22	73	65
				439	77.9	20.0	129		5.9	0.13	165	7
	13320-13340' (w)	same	0.85	444	1.00	0.22	0.64	0.86	0.71	0.26	64	71
				443	76.2	21.9	106		4.9	0.17	139	7
	13690-13700' (w)	same	0.84	444	0.71	0.15	0.37	0.52	0.66	0.29	52	92
				443	76.5	20.6	111		6.5	0.16	145	8
	13810-13840' (w)	same	0.95	451	0.89	0.17	0.44	0.61	0.98	0.28	40	110
				446	75.6	20.6	101		5.4	0.17	133	7
Uniacke G-72	13970-14000' (w)	same	1.02	452	0.56	0.17	0.30	0.47	0.76	0.36	53	135
				436	1.74	0.40	0.94	1.34	1.74	0.30	54	100
	4270-4280 m (w)	Cimalta Limestone	0.71	431	72.5	13.6	81.9		34.1	0.		
				430	8.64	0.38	3.58	3.96	13.6	0.10	41	157
	4985-4990 m (w)	Penobscot Limestone	1.03	418	60.9	6.3	62.0		41.7	0.09	101	68
				439	1.43	0.29	0.46	0.75	2.17	0.39	32	151
	5020-5030 m (w)	Mic Mac	1.18	423#	67.2	13.3	68.0		35.9	0.16	101	53
				427#	10.8	1.00	7.39		20.5	0.12	68	190
	5140-5150 m (w)	same	1.28	425#	60.3	12.1	53.8		35.6	0.15	89	65
				428#	1.18	0.86	0.82	1.68	1.28	0.51	69	108
	5665-5675 m (w)	same	1.52	430#	1.73	0.59	0.84	1.43	1.80	0.41	48	104
				423#	71.7	20.4	79.3		36.2	0.20	110	50

(w) = whole rock; (k) kerogen concentrate; R₀ = mean random vitrinite reflectance; T_{max} = maturation parameter from Rock-Eval pyrolysis in °C. TOC = total organic carbon in weight %; S₁ and S₂ = in mg HC/g TOC; S₃ = mg CO₂/g TOC; PI = production index = S₁/S₁+S₂; HI = hydrogen index in mg HC/g TOC; OI = oxygen index = mg CO₂/g TOC. # = possible contamination

TABLE 1B : Rock-Eval pyrolysis data of composite shale samples

Borehole No.	Depth (ft/m)	Formation/Member	R _o	T _{max} (°C)	TOC (wt%)	S ₁	S ₂	S ₁ +S ₂	S ₃	PI	HI	OI
Alma F-67	3945 - 4130 m	Verrill Canyon	0.90	441	1.97	3.00	3.22	6.22	1.41	0.48	166	73
	4400 - 4595 m	same	1.15	433	1.47	2.62	5.81	8.43	1.70	0.31	399	117
	4700 - 4890 m	same	1.40	435	2.99	6.39	9.43	15.82	1.81	0.40	315	60
Chebucto K-90	2450 - 2650 m	Marmora	0.50	429	1.42	2.19	0.90	3.09	3.11	0.70	62	217
	3545 - 3620 m	Cree	0.68	435	2.76	2.22	1.36	3.58	3.18	0.62	49	115
	3730 - 3800 m	same	0.76	439	3.07	1.85	1.54	3.39	4.24	0.55	53	68
	3960 - 4115 m	Naskapi	0.80	443	2.50	2.05	1.34	3.39	1.70	0.63	49	69
	4335 - 4450 m	Mississauga	0.90	444	2.34	1.24	1.54	2.78	1.74	0.45	65	74
	4655 - 4885 m	same	1.25	432	2.39	1.12	1.55	2.67	2.42	0.42	64	101
Glencol J-48	4975 - 5090 m	same	1.55	440	1.60	0.92	0.54	1.46	1.39	0.63	33	86
	2850 - 2970 m	Cree	0.60	433	2.93	1.38	1.83	3.21	3.12	0.43	62	106
	3540 - 3750 m	U. Mississauga	0.68	440	1.55	0.57	0.92	1.49	1.07	0.39	59	69
	3975 - 4150 m	same	0.90	444	1.74	0.39	1.06	1.45	0.89	0.27	61	51

R_o = mean random vitrinite reflectance; T_{max} = maturation parameter from Rock-Eval pyrolysis in °C. TOC = total organic carbon in weight %.
 S₁ and S₂ = in mg HC/g TOC; S₃ = mg CO₂/g TOC; PI = production index = S₁/S₁+S₂; HI = hydrogen index in mg HC/g TOC; OI = oxygen index = mg CO₂/g TOC.

TABLE 1B : Rock-Eval pyrolysis data of composite shale samples

Borehole No.	Depth (ft/m)	Formation/Member	R _o	T _{max} (°C)	TOC (wt%)	S ₁	S ₂	S ₁ +S ₂	S ₃	PI	HI	OI
Glenelg J-48	4255 - 4400 m	U. Mississauga	1.02	431	4.59	1.26	1.46	2.72	5.03	0.46	31	109
	4570 - 4750 m	M. Mississauga	1.10	439	1.46	0.86	0.57	1.43	1.83	0.57	39	125
	4900 - 5025 m	Verrill Canyon	1.30	430	4.78	1.36	3.28	4.64	4.60	0.29	68	96
N.Truimph B-52	2660 - 2850 m	Cree		431	2.27	7.87	2.36	10.23	2.25	0.77	103	99
	2860 - 3115 m	same	0.44	433	3.46	10.41	4.06	14.47	2.15	0.72	117	62
	3210 - 3400 m	same	0.54	434	4.69	10.60	6.20	16.80	2.05	0.63	132	43
	3410 - 3560 m	Naskapi	0.57	435	3.52	10.04	5.75	15.79	2.25	0.64	163	63
	3560 - 3715 m	Naskapi	0.63	436	3.55	10.77	4.96	15.73	1.93	0.68	142	54
N.Truipmh G-43	2660 - 2750 m	Cree	0.58	426	2.08	5.06	3.86	8.92	1.41	0.57	185	67
	2840 - 2975 m	same	0.59	433	3.35	7.08	5.58	12.66	2.07	0.56	166	61
	3360 - 3550 m	same	0.65	427	3.01	6.30	5.39	11.69	1.53	0.54	179	50
	3600 - 3710 m	Naskapi	0.69	426	3.75	10.59	9.37	19.96	1.63	0.53	249	43
	4530 - 4600 m	Mississauga	0.83	426	2.83	6.74	11.6	18.34	1.97	0.37	410	69
	4680 - 4755 m	same	1.12	418	3.64	11.43	11.9	23.3	1.83	0.49	331	51
S.Desbarres O-76	2345 - 2425 m	Cree	0.52	428	2.86	5.08	3.52	8.60	1.63	0.59	123	56
	3150 - 3195 m	M. Mississauga	0.70	426	1.74	3.60	3.24	6.84	1.34	0.53	187	76
	3910 - 4075 m	Mic Mac	0.97	430	1.30	1.98	3.88	5.86	1.01	0.34	298	77
	4355 - 4515 m	same	1.14	422	2.24	4.45	6.06	10.51	1.59	0.42	270	70
	5025 - 5250 m	same	1.50	426	1.89	5.24	5.49	10.73	1.60	0.49	290	84
	5590 - 5705 m	same	1.78	431	2.55	5.73	10.4	16.13	2.23	0.36	408	87

R_o = mean random vitrinite reflectance; T_{max} = maturation parameter from Rock-Eval pyrolysis in °C. TOC = total organic carbon in weight %.
 S₁ and S₂ = in mg HC/g TOC; S₃ = mg CO₂/g TOC; PI = production index = S₁/S₁+S₂; HI = hydrogen index in mg HC/g TOC; OI = oxygen index = mg CO₂/g TOC.

TABLE 1B : Rock-Eval pyrolysis data of composite shale samples

Borehole No.	Depth (ft/m)	Formation/Member	R _o	T _{max} (°C)	TOC (wt%)	S ₁	S ₂	S ₁ +S ₂	S ₃	PI	HI	OI
Venture B-43	5310 - 5475 m	Mic Mac	1.44	434#	4.14	1.66	1.11	2.77	4.69	0.60	27	116
Venture B-52	2310 - 2485 m	Cree	0.45	430	1.35	0.82	0.76	1.58	2.86	0.52	56	211
	2875 - 3045 m	Naskapi	0.57	436	1.64	1.31	0.71	2.02	1.84	0.65	43	112
	3680 - 3740 m	M. Mississauga	0.72	438	2.87	0.57	1.47	2.04	1.70	0.28	51	59
	5810 - 5960 m	Mic Mac	1.66	429#	5.31	1.00	1.40	2.40	5.72	0.42	27	110
Venture H-22	3860 - 3875 m	M. Mississauga	0.74	445	1.44	0.45	0.90	1.35	0.83	0.34	62	57
	4175 - 4385 m	same	0.77	445	1.10	0.41	0.68	1.09	0.75	0.38	61	68
	4410 - 4525 m	same		447	1.10	0.35	0.82	1.17	0.57	0.30	74	51
	4550 - 4710 m	same	0.86	448	1.19	0.48	0.61	1.09	0.75	0.44	51	63
W.ChebuctoK-20	2770 - 2930 m	Cree	0.43	432	2.22	6.62	4.19	10.8	1.49	0.61	188	67
	2950 - 3035 m	same	0.49	433	2.21	5.08	4.23	9.31	1.40	0.55	191	63
	3070 - 3170 m	same	0.55	434	2.87	8.80	5.67	14.47	1.42	0.61	197	49
	3340 - 3440 m	same	0.63	435	2.63	5.92	5.77	11.69	1.67	0.51	219	63
	3890 - 4010 m	Naskapi	0.65	438	3.15	6.48	6.30	12.78	1.50	0.51	200	47
	4165 - 4350 m	Mississauga	0.72	440	2.91	5.93	5.12	11.05	1.67	0.54	175	57
	4400 - 4570 m	same	0.75	439	3.06	8.97	6.57	15.54	1.49	0.58	214	48
	4825 - 5000 m	same	1.09	435#	2.99	7.61	5.92	13.53	1.85	0.56	197	61

R_o = mean random vitrinite reflectance; T_{max} = maturation parameter from Rock-Eval pyrolysis in °C. TOC = total organic carbon in weight %.
 S₁ and S₂ = in mg HC/g TOC; S₃ = mg CO₂/g TOC; PI = production index = S₁/S₁+S₂; HI = hydrogen index in mg HC/g TOC; OI = oxygen index =
 mg CO₂/g TOC. # = possible contamination.

TABLE 2A: Maceral Composition in volume percent (organic facies), R_o , kerogen type, and oil/gas potential of limestone and calcareous shale samples

Borehole No.	Depth (ft/m)	Fm/Mbr	R_o	Vit	Int	Exe	Res	Lam	Tel	Ltd	ao m 1	aom 2	aom 3	Bit	Ra. Int	Kerog. Type	Oil/Gas Potent.
Abenaki J-56	10090'	Baccaro	0.51	34.5	7.5	27.0		8.0	0.5	2.5		19.0	0.5	0.5		IIA-IIIBL	Oil/Cond
	10800'	same	0.53	38.5	15.0	25.5	0.5	2.0	2.5			8.0	7.0	1.0		IIB	Cond/Gas
	12040'	same	0.57	50.0	16.5	23.5		1.0	1.0	1.0		6.0		1.0		IIB	Cond/Gas
	12360'	same	0.58	33.5	20.5	20.5		3.5	2.0	3.0		15.0				IIA-IIIBL	Oil/Cond
	12890'	same	0.57	50.5	21.5	17.0		1.0	1.0	1.0		8.0				IIB	Cond/Gas
	13040'	same	0.60	43.0	18.0	26.5	2.5	1.0		2.5		4.5		2.0		IIB	Cond/Gas
	13270'	Misaine	0.78	57.0	22.0	19.5		0.5				0.5		0.5		III	Gas
	13790'	Scatar.	0.76	61.5	15.5	22.0						1.0				IIB	Cond/Gas
	14080'	Mohican	0.82	64.0	13.5	13.0	2.0	1.0	1.0			1.5		1.0	3.0	IIB/III	Gas
	14140'	same	0.89	68.0	17.5	9.0	1.0	2.5	0.5			1.0				III	Gas
	14220'	same	0.94	63.5	18.0	7.0	0.5	0.5	0.5			1.0		0.5	8.5	IIB	Cond/Gas
	14350'	same	0.92	63.5	25.0	8.0		2.0				1.5				III	Gas
	14410'	Iroquo.	0.91	59.0	25.5	10.0	0.5		0.5			1.5		3.0		III	Gas
CohassetD-42	10000'	L.Miss.	0.64	58.0	19.5	17.0		1.5	1.0			2.5		0.5		IIB	Cond/Gas
	10830'	Baccaro	0.71	44.5	33.0	15.5	0.5	1.0		0.5		1.5		0.5	3.0	III	Gas
	13430'	same	0.63	55.5	14.5	12.5	1.0	2.0*	3.0*	0.5		7.5*		3.5		III/IV	Gas
	13740'	same	1.23	10.0	73.0	5.0*						5.5*	0.5	4.0	2.0	IV	Nonsour.
	14160'	Misaine	1.23	19.0	68.0	6.5*		0.5*	0.5*	0.5*		1.5*		1.0	2.0	IV	Nonsour.
	14370'	same	1.33	28.0	54.5	6.0*			0.5*	0.5*		0.5		4.0	5.5	IV	Nonsour.

LEGEND

Fm = Formation/Member; R_o = mean random vitrinite reflectance; Vit = vitrinite; Int = Inertinite; Exe = Exenite; Res = Resinite; Lam = Lamalinite; Tel = telalinite
 Ltd = liptodetrinite; aom 1 = amorphous organic matter 1 (oil); aom 2 = amorphous organic matter 2 (oil and condensate); aom 3 = amorphous organic matter 3 (gas);
 Bit = solid bitumen; Ra. Int = rank inertinite or oxidized liptinite; Kerog. Type = Kerogen Type; Oil/gas potent. = oil and gas potential.
 * = oxidized phytoclasts; Nonsour. = nonsource rock. Baccaro = Baccaro Member; Misaine = Misaine Member; Scatar. = Scatarie Member; Mohican = Mohican Formation; Iroquo. = Iroquois Formation; L.Miss. = Lower Mississauga Formation; VerrCan = Verrill Canyon Formation.

TABLE 2A: Maceral Composition in volume percent (organic facies), R_o , kerogen type, and oil/gas potential of organic-rich limestone and calcareous shale samples

Borehole No.	Depth (ft/m)	Fm	R_o	Vit	Int	Exe	Res	Lam	Tel	Ltd	ao m 1	aom 2	aom 3	Bit	Ra. Int	Kerog. Type	Oil/Gas Potent.
CohassetD-42	1450'	Misaine	1.34	32.0	50.5	5.5		1.0*	1.0*			0.5		1.0	8.5	III	Gas
CohassetL-97	3100 m	L.Miss.	0.63	47.0	23.0	26.0	0.5	0.5	0.5*			1.0*	1.0	0.5		IIB	Cond/Gas
	3195 m	Baccaro	0.62	34.5	41.0	16.5		0.5	1.5			1.5	4.5			III	Gas
	3990 m	same	-	50.0	25.0	10.0	15.0									IIB	Cond/Gas
	4155 m	Baccaro	-	55.0	35.0	5.0								5.0		III	Gas
	4470 m	Misaine	1.41	32.5	49.0				2.0					0.5	16.0	III-IV	Gas-nonsour.
	4545 m	same	1.50	30.5	51.0	0.5			0.5					0.5	17.0	IV	Nonsour.
	4620 m	Scatar.	1.6	29.0	55.5										15.5	IV	Nonsour.
	4855 m	Mohican	1.92	20.5	42.5				3.0			9.0		0.5	24.5	IIB	Cond/Gas
	4870 m	Same	1.82	20.0	34.0				0.5					0.5	45.0	IIA-IIB overmat.	Oil/Cond
Demoscota G-43	11480'	VerrCan	0.73	52.5	11.5	28.0	1.5	3.0				3.5				IIB	Cond/Gas
	11840'	Baccaro	0.72	54.5	23.0	17.0	1.0	2.0*	1.0	0.5		0.5		0.5		III	Gas
	14550'	Baccaro	-	53.0	25.0	17.0			4.0*			1.0				III	Gas
	14820'	Baccaro	0.75	46.0	51.0	9.0*				1.0		2.0*	1.0			III-IV	Nonsour.
	15060'	Misaine	1.17	38.5	48.5	12.0	*		0.5*			0.5*				IV	Nonsour.
	15250'	Scatar.	1.32	39.0	42.0	16.0		2.0	1.0							IIB	Cond/Gas

LEGEND

Fm = Formation/Member; R_o = mean random vitrinite reflectance; Vit = vitrinite; Int = Inertinite; Exe = Exenite; Res = Resinite; Lam = Lamalgnite; Tel = telalgnite
 Ltd = liptodetrinite; aom 1 = amorphous organic matter 1 (oil); aom 2 = amorphous organic matter 2 (oil and condensate); aom 3 = amorphous organic matter 3 (gas);
 Bit = solid bitumen; Ra. Int = rank inertinite or oxidized liptinite; Kerog. Type = Kerogen Type; Oil/gas potent. = oil and gas potential.
 * = oxidized phytoclasts; Nonsour. = nonsource rock. Baccaro = Baccaro Member; Misaine = Misaine Member; Scatar. = Scatarie Member; Mohican = Mohican Member; Iroquois Formation = Lower Mississauga Formation; VerrCan = Verrill Canyon Formation; MicMac = Mic Mac Formation; CitLs = Citnalta Limestone Member; PenLs = Penobscot Limestone Member.

TABLE 2A: Maceral Composition in volume percent (organic facies), R_o , kerogen type, and oil/gas potential of organic-rich limestone and calcareous shale samples

Borehole No.	Depth (ft/m)	Fm	R_o	Vit	Int	Exe	Res	Lam	Tel	Ltd	ao m 1	aom 2	aom 3	Bit	Ra. Int	Kerog. Type	Oil/Gas Potent.
Penobscot L-30	12280'	MicMac	0.61	48.5	30.5	19.5		0.5	0.5			0.5				IIB	Cond/Gas
	12440'	same	0.67	49.0	28.0	21.5						1.5				IIB	Cond/Gas
	12580'	same	0.73	48.0	29.5	19.0		0.5	0.5			2.5				IIB	Cond/Gas
	12730'	same	0.73	51.5	27.0	19.5		1.0						1.0		IIB	Cond/Gas
	12980'	same	0.77	50.5	27.0	16.5		2.0	1.0	2.5		0.5				IIB	Cond/Gas
	13320'	same	0.85	29.0	48.0	19.0	2.0	0.5	0.5			1.0				III	Gas
	13690'	same	0.84	33.0	34.5	28.0	1.5	1.0	1.0			1.0				IIB	Cond/Gas
	13810'	same	0.95	30.0	42.5	19.0		1.0	1.0					1.5	5.0	III	Gas
	13970'	same	1.02	35.5	40.0	17.0	0.5	0.5	0.5					2.0	4.0	III	Gas
	4270 m	CitLs	0.71	43.5	37.0	15.5	2.5	0.5	0.5	0.5		0.5					IIB
4985 m	PenLs	1.03	26.5	63.0	8.0	1.0		1.0*						0.5		III	Gas
5020 m	MicMac	1.18	40.5	40.5	6.0			2.0						2.0	8.5	IIB	Cond/Gas
5140 m	MicMac	1.28	30.0	46.5	8.0						1.0			2.0	12.5	IIB	Cond/Gas
5665 m	MicMac	1.52	20.5	63.5	11.5										4.0	III	Gas
5735 m	MicMac	1.57	29.0	55.0	11.0									3.0	III	Gas	

LEGEND

Fm = Formation/Member; R_o = mean random vitrinite reflectance; Vit = vitrinite; Int = inertinite; Exe = exenite; Res = resinite; Lam = lamalinite; Tel = telalinite
 Ltd = liptodetrinite; aom 1 = amorphous organic matter 1 (oil); aom 2 = amorphous organic matter 2 (oil and condensate); aom 3 = amorphous organic matter 3 (gas);
 Bit = solid bitumen; Ra. Int = rank inertinite or oxidized liptinite; Kerog. Type = Kerogen Type; Oil/gas potent. = oil and gas potential.
 * = oxidized phytoclasts; Nonsour. = non-source rock. Baccaro = Baccaro Member; Misaine = Misaine Member; Scatar. = Scatarie Member; Mohican = Mohican Formation; Iroquois Formation; L.Miss. = Lower Mississauga Formation; VerrCan = Verrill Canyon Formation; MicMac = Mic Mac Formation; CitLs = Citnalla Limestone Member; PenLs = Penobscot Limestone Member; MicMac = Mic Mac Formation.

TABLE 2B: Maceral Composition in volume percent (organic facies), R_o , kerogen type, and oil/gas potential of composite shale samples

Borehole No.	Depth (ft/m)	Fm/Member	R_o	Vit	Int	Exe	Res	Lam	Tel	Ltd	ao m 1	aom 2	aom 3	Bit	Ra. Int	Kerog. Type	Oil/Gas Potent.
Alma F-67	3945 m	VerrCan	0.90	53.0	12.0	27.0	2.0	3.0*	1.0			2.0				IIB	Cond/Gas
	4400	same	1.15	42.0	8.0	5.0		3.0				40.0		2.0		IIA-IIB	Oil/Cond
	4700	same	1.40	28.0	5.0	5.0		2.0	1.0			45.0	5.0	3.0	6.0	IIA-IIB	Oil/Cond
ChebuctoK-90	2450 m	Marmor.	0.50	60.0	25.5	9.0			0.5			3.0	2.0			III	Gas
	3545 m	Cree	0.68	64.0	24.5	8.0	0.5		0.5			2.0	0.5			III	Gas
	3730 m	Cree	0.76	62.0	28.0	5.0	2.0					2.0	1.0			III	Gas
	3960 m	Naskapi	0.80	58.0	28.0	8.0	1.0						5.0			III	Gas
	4335 m	Miss.	0.90	72.0	7.0	19.6							0.5			IIB	Gas
	4655 m	Miss.	1.25	68.0	20.0	10.0			1.0*				1.0			III	Gas
	4975 m	Miss.	1.55	71.0	16.0	10.0								2.0	1.0	III	Gas
Glengel J-48	2850 m	Cree	0.60	54.0	29.0	15.0*			1.5					0.5		III	Gas
	3540 m	U.Miss.	0.68	48.0	31.5	19.5*			1.0*							III	Gas
	3975 m	same	0.90	63.5	19.0	16.0*	0.5		1.0*							III	Gas
	4255 m	same	1.02	38.0	50.0	9.5			1.0*				0.5	1.0		III	Gas
	4570 m	M.Miss.	1.10	46.0	35.0	15.5*		1.5*	2.0							III	Gas
	4900 m	VerrCan	1.30	36.0	41.0	11.0						10.0		2.0		IIB	Cond/Gas

LEGEND

Fm = Formation/Member; R_o = mean random vitrinite reflectance; Vit = vitrinite; Int = Inertinite; Exe = Exenite; Res = Resinite; Lam = Lamalinite; Tel = telalinite
 Ltd = liptodetrinite; aom 1 = amorphous organic matter 1 (oil); aom 2 = amorphous organic matter 2 (oil and condensate); aom 3 = amorphous organic matter 3 (gas);
 Bit = solid bitumen; Ra. Int = rank inertinite or oxidized liptinite; Kerog. Type = Kerogen Type; Oil/gas potent. = oil and gas potential.

* = oxidized phytoclasts; Nonsour. = nonsource rock. Marmor. = Marmora Member; Cree = Cree Member; U.Miss. = Upper Mississauga Formation; M.Miss. Middle Mississauga Formation; VerrCan = Verrill Canyon Formation.

TABLE 2B: Maceral Composition in volume percent (organic facies), R_o , kerogen type, and oil/gas potential of composite shale samples

Borehole No.	Depth (ft/m)	Fm	R_o	Vit	Int	Exe	Res	Lam	Tel	Ltd	ao m 1	aom 2	aom 3	Bit	Ra. Int	Kerog. Type	Oil/Gas Potent.
N.Truiumph B-52	2660 m	Cree		27.0	46.0	22.0			2.0			2.0		1.0		IIB	Con/Gas
	2860 m	same	0.44	16.0	45.0	24.0	1.0					9.0		5.0		IIB	Cond/Gas
	3210 m	same	0.54	21.0	35.0	3.0						20.0	21.0			IIB	Cond/Gas
	3410 m	Naskapi	0.57	15.0	23.0	3.0						15.0	42.0			IIB	Cond/Gas
N.Truiumph G-43	3560 m	Naskapi	0.63	17.0	13.0	1.0		3.0				43.0	21.0	2.0		IIA-IIB	Oil/Cond
	2660 m	Cree	0.58	31.0	31.0	5.0		5.0				25.0	2.0	1.0		IIB	Cond/Gas
	2840 m	same	0.59	27.0	35.0			4.0				24.0	6.0	4.0		IIB	Cond/Gas
	3360 m	same	0.65	21.0	32.0	26.0						5.0	13.0			IIB	Cond/Gas
S.Desbarres O-76	3600 m	Naskapi	0.69	12.0	20.0	1.0		1.0	1.0			43.0	17.0	5.0		IIA-IIB	Oil/Cond
	4530 m	Miss.	0.83	18.0	16.0			2.0				64.0				IIA-IIB	Oil/Cond
	4680 m	same	1.12	29.0	19.0	1.0						44.0	4.0	3.0		IIA-IIB	Oil/Cond
	2345 m	Cree	0.52	40.0	28.0	29.0			1.0			2.0				IIB	Cond/Gas
S.Desbarres O-76	3150 m	M.Miss.	0.70	42.0	17.0	40.0						1.0				IIB	Cond/Gas
	3910 m	MicMac	0.97	37.0	11.0	39.0		1.0	1.0			11.0				IIA-IIB	Oil/Cond
	4355 m	same	1.14	37.0	18.0	18.0			3.0			13.0		2.0	9.0	IIA-IIB	Oil/Cond
	5025 m	same	1.50	35.0	16.0	24.0			2.0			18.0			5.0	IIA-IIB	Oil/Cond

LEGEND

Fm = Formation/Member; R_o = mean random vitrinite reflectance; Vit = vitrinite; Int = Inertinite; Exe = Exenite; Res = Resinite; Lam = Lamalinite; Tel = telalinite
 Ltd = liptodetrinite; aom 1 = amorphous organic matter 1 (oil); aom 2 = amorphous organic matter 2 (oil and condensate); aom 3 = amorphous organic matter 3 (gas);
 Bit = solid bitumen; Ra. Int = rank inertinite or oxidized liptinite; Kerog. Type = Kerogen Type; Oil/gas potent. = oil and gas potential.
 * = oxidized phytoclasts; Nonsour. = nonsource rock. Marmor. = Marmora Member; Cree = Cree Member; U.Miss. = Upper Mississauga Formation; Miss. = Mississauga Formation; M.Miss = Middle Mississauga Formation; MicMac = Mic Mac Formation.

TABLE 2B: Maceral Composition in volume percent (organic facies), R_o, kerogen type, and oil/gas potential of composite shale samples

Borehole No.	Depth (ft/m)	Fm	R _o	Vit	Int	Exe	Res	Lam	Tel	Ltd	ao m I	aom 2	aom 3	Bit	Ra. Int	Kerog. Type	Oil/Gas Potent.
S. Desbarres O-76	5590 m	MicMac	1.78	16.0	16.0							68.0				IIA-IIIB	Oil/Cond
Venture B-43	5310 m	same	1.44	45.0	42.0	9.0							4.0			III	Gas
Venture B-52	2310 m	Cree	0.45	45.0	35.0	20.0*										III	Gas
	2875 m	Naskapi	0.57	33.0	47.0	19.0*							1.0			III	Gas
	3680 m	M.Miss.	0.72	32.0	49.0	19.0										III	Gas
	5810 m	Mic Mac	1.66	35.0	55.0	10.0										III	Gas
Venture H-22	3860 m	M.Miss.	0.74	21.0	60.0	18.0										III	Gas
	4175 m	same	0.77	14.0	65.0	19.0		1.0	1.0							III	Gas
	4410 m	same	-	18.0	54.0	27.0*	1.0									III	Gas
	4550 m	same	0.86	17.0	53.0	29.0*						1.0				III	Gas
W. Chebucto K-20	2770 m	Cree	0.43	59.5	12.5	16.5	0.5		2.5			8.0	0.5			IIB	Cond/Gas
	2950 m	same	0.49	55.5	20.5	14.5	1.5	1.5	0.5	0.5		5.0	0.5			IIB	Cond/Gas
	3070 m	same	0.55	53.5	19.0	14.0	0.5	1.5	1.5			9.0	0.5	0.5		IIB	Cond/Gas
	3340 m	same	0.63	58.5	13.0	21.0	0.5	2.5				4.5				IIB	Cond/Gas
	3890 m	Naskapi	0.65	53.0	15.0	20.0	1.5	0.5	2.0			8.0				IIB	Cond/Gas
	4165 m	Miss.	0.72	43.5	15.0	18.0*	1.0	0.5		0.5		21.0*	0.5			IIB	Cond/Gas
	4400 m	same	0.75	39.0	24.5	13.5		0.5	0.5			22.0*				IIB	Cond/Gas
	4825 m	same	1.09	41.5	15.0	6.5			1.0			25.5				IIB	Cond/Gas

LEGEND: Fm = Formation/Member; R_o = mean random vitrinite reflectance; Vit = vitrinite; Int = inertinite; Exe = Exenite; Res = Resinite; Lam = Lamalinite; Tel = telalinite; Ltd = liptodetrinite; aom 1 = amorphous organic matter 1 (oil); aom 2 = amorphous organic matter 2 (oil and condensate); aom 3 = amorphous organic matter 3 (gas); Bit = solid bitumen; Ra. Int = rank inertinite or oxidized liptinite; Kerog. Type = Kerogen Type; Oil/gas potent. = oil and gas potential.
 * = oxidized phytoclasts; Nonsour. = non-source rock. Naskapi = Naskapi Member; Cree = Cree Member; U.Miss. = Upper Mississauga Formation; Miss. = Mississauga Formation; M.Miss = Middle Mississauga Formation; MicMac = Mic Mac Formation.

TABLE 3. ROCK-EVAL 5 DERIVED KINETIC PARAMETERS
 Five heating rates: 1, 5, 15, 25, and 50 C/min

Sample id.	Green River Shale (AP22)	Venture B-52 3680 m	N. Triumph B-52 3560 m	Demoscola G-43 11840-11880 ft	Abenaki J-56 10090-1013- ft
Kerogen Type	I	III	IIA-IIIB	III	IIA-IIIBL
.. Arrhenius factor (/my)	1.3844244E+27	3.8702914E+25	4.4367602E+24	9.6091408E+25	1.8258914E+26
Percent of Reaction					
Activation Energy (kcal/mole)					
39			1.86		
40			0.94		
41			4.48		
42		0.29			0.06
43		0.78			0.51
44	0.23	1.06	22.33		
45		1.83	35.36		
46	1.70		19.20		
47	0.07	17.34	5.84		
48		38.20	3.87		
49	7.32	8.78			
50	2.31	13.57		34.03	49.38
				23.14	13.45
51		6.34	6.12	21.20	20.72
52		0.59		5.00	
53	84.50			6.82	9.97
54					2.68
55		11.22		5.77	
56	2.42			4.04	0.90
57	0.31				1.09
58					
59	1.14				
60					1.24
61					
62					
63					
64					
65					
66					
67					
68					
69					
70					
Standard Error (kcal/mole)	437	895	300	2805	581

Table 4--Scotian Shelf oil/condensates and source rock samples.

Well	SIU Sample No.	DST/ RFT	Depth (m)	Formation	Member
Petroleum Samples					
Arcadia J-16	NS9116	5	5156-5175	Missisauga	Lower
Balmoral M-32	NS9302	3	1954-1958	L.Canyon	Cree
Banquereau C-21	NS9113	1	3585-3596	Missisauga	Lower
Bluenose 2G-47	NS9115	8	4577-4590	Mic Mac	
Chebucto K-90	NS9001	4	4227-4238	Missisauga	Upper
CitnaltaI59	NS9112	3	3777-3781	Missisauga	Lower
Cohasset A-52	NS9004	5	1888-1891	Logan Canyon	Cree
Cohasset D-42	NS9002	7	1861-1865	Logan Canyon	Cree
Glenelg J-48	NS9010	8	3491-3495	Missisauga	Upper
N. Triumph B-52	NS9006	4	3771-3777	Missisauga	Upper
Olympia A-12	NS9007	5	4664-4678	Missisauga	Lower
Panuke B-90	NS9003	1	2293-2299	Missisauga	Upper
Panuke PP3 J-99	NS9301	1	2572-2580	L. Canyon	Cree
Penobscot L-30	NS9114	5	2642	Missisauga	Upper
S. Venture O-S9	NS9008	5	5035-5050	Missisauga	Lower
S. Venture O-S9	NS9009	10	4255-4267	Missisauga	Lower
Sable Island 3H-58	NS9005	5	1435-1436	Logan Canyon	Mamora
Thebaud C-74	NS9118	9	3865-3888	Missisauga	Lower
Uniacke G-72	NS9303	6	5191-5199	Mic Mac	
Venture H-22	NS9117	5	5021-5025	Mic Mac	
Venture H-22	NS9304	7	4957-4962	Mic Mac	
Extract Samples					
Abenaki J-56	NS9305		3080	Missisauga	Lower
Alma F-67	NS9013		4500	Verrill Canyon	
Alma F-67	NS9108		5045	Verrill Canyon	
Cohasset A-52	NS9011		2036	Logan Canyon	Cree
Cohasset D-42	NS9012		4426	Abenaki	Misaine
Demascota G-43	NS9306		3615	Abenaki	Baccaro
Migrant N-20	NS9101		3587	Missisauga	Lower
N. Triumph G-43	NS9103		4845	Missisauga	undif.
N. Triumph G-43	NS9109		3695	Logan Canyon	Naskapi
Penobscot L-30	NS9105		2118	Logan Canyon	undif.
Penobscot L-30	NS9307		4214	Mic Mac	Penobscot
S. Desberres 0-76	NS9107		3861	Missisauga	Lower
S. Sable B-44	NS9111		3938	Missisauga	Middle
S. Venture O-S9	NS9014		6115	Mic Mac	undif.
Thebaud C-74	NS9110		3911	Missisauga	Lower
Venture B-52	NS9102		5121	Missisauga	Lower
W. Chebucto K-20	NS9104		5210	Missisauga	undif.
Whycocomagh N-30	NS9106		3360	Missisauga	undif.

Liquid chromatographic data for 1993 Scotian Shelf samples (normalized %).

	Saturates	Aromatics	Polar 1	Polar 2
Petroleums				
Panuke PP3 J-99 (DST1)	68	26	5	1
Balmoral M-32 (DST3)	70	24	5	1
Uniacke G-72 (DST6)	46	47	6	1
Venture H-22 (DST7)	57	36	4	3
Rock Extracts				
Abenaki J-56 (3080m)	43	16	10	30
Demascota G-43 (3615m)	38	26	6	30
Penobscot L-30 (4214m)	51	32	2	15

Table 5

Table 6: GC/MS peak identification.

Peak Compound (Isomers specified if known)	Δm	Standard *Used in Deviation correlation
m/z 170 (C3-Alkyl-naphthalenes)		
170a methylethyl-naphthalene	0.06	0.23 *
170b methylethyl-naphthalene	0.02	0.22 *
170c 1,3,7-trimethyl-naphthalene (TrMN)	0.25	0.22
170d 1,3,6-TrMN	0.11	0.18 *
170e 1,4,6- & 1,3,5-TrMN	-0.04	0.17 *
170f 2,3,6-TrMN	0.12	0.17 *
170g 1,6,7- & 1,2,6-TrMN	-0.19	0.16 *
170h 1,2,5-TrMN	- 1.02	0.36
m/z 184 (C4-alkyl-naphthalenes and dibenzothiophene)		
184a 1,3,5,7-tetramethyl-naphthalene (TeMN)	-0.02	0.18 *
184b 1,3,6,7-TeMN	-0.02	0.16 *
184c 1,2,4,7-TeMN	-0.55	0.29
184d 1,2,5,7-TeMN	-0.78	0.33
184e 2,3,6,7-TeMN	-0.03	0.17 *
184f 1,2,6,7-TeMN	-0.55	0.25
184g 1,2,3,7-TeMN	-0.50	0.31
184h 1,2,3,6-TeMN	-0.56	0.33
184i 1,2,5,6-TeMN	-1.61	0.67
DBT dibenzothiophene	-0.50	0.89
m/z 182 (Dimethylbiphenyls and methyl-dibenzofurans)		
182a	0.64	0.44
182b 3,5-dimethylbiphenyl (DMBP)	0.52	0.28
182c 3,3'-DMBP	0.39	0.24
182d 3,4'-DMBP	0.40	0.23
182e	0.80	0.40
182f 3,4-DMBP	0.31	0.20
mDBFa methyl-dibenzofuran	-0.59	0.26
mDBFb methyl-dibenzofuran	-0.84	0.33
mDBFc methyl-dibenzofuran	-0.46	0.33
m/z 198 (Methyl-dibenzothiophenes)		
198a 4-methyl-dibenzothiophene (MDBT)	0.32	0.30
198b 3- & 2-MDBT	0.08	0.29 *
198c 1-MDBT	-0.85	0.76

Table 6

Table 6: Continued

Peak Compound asomers specified if known)	Δm	Standard *Used in Deviation correlation
m/z 212 (C2-Alkyldibenzothiophenes)		
212a ethyldibenzothiophene	0.17	0.38 *
212b 4,6-dimethyldibenzothiophene (DMDBT ⁻)	0.35	0.38
212c 2,4-DMDBT	0.43	0.35
212d 2,6- & 3,6-DMDBT	0.26	0.37
212e 3,7-DMDBT	-0.03	0.38 *
212f 1,4- & 1,6- & 1,8-DMDBT	-0.07	0.46 *
212g 1,3- & 1,2- & 1,9-DMDBT	0.05	0.45 *
m/z 178 (Phenanthrene)		
PHN phenanthrene	0.00	0.07
m/z 192 (Methylphenanthrenes)		
192a 3-methylphenanthrene (MP)	0.38	0.15
192b 2-MP	0.27	0.14
192c 9-MP	0.10	0.12
192d 1-MP	-0.04	0.09
m/z 206 (C2-Alkylphenanthrenes)		
206a 3-ethylphenanthrene (ET)	0.16	0.28 *
206b 2-EP & 9-EP & 3,6-DMP	-0.09	0.25 *
206c 1-EP	0.62	0.41
206d 2,6-dimethylphenanthrene (DMP)	0.60	0.26
206e 2,7-DMP	0.56	0.25
206f 1,3- & 2,10- & 3,9- & 3,10-DMP	0.41	0.23
206g 1,6- & 2,9-DMP	0.35	0.19
206h 1,7-DMP	0.09	0.18 *
206i 2,3-DMP	0.40	0.20
206j 1,9-DMP	-0.02	0.20 *
206k 1,8-DMP	0.05	0.19 *
206l	-0.80	0.41
m/z 202 (Fluoranthene and pyrene)		
FLAN fluoranthene	-0.64	0.28
PYR pyrene	-0.54	0.31
m/z 216 (Methylpyrenes and isomers)		
216a	-0.22	0.16 *
216b benzo[a]fluorene	-0.05	0.19 *
216c	-0.01	0.79 *
216d 2-methylpyrene	0.04	0.21 *
216e 4-methylpyrene	-0.17	0.17 *
216f 1-methylpyrene	-0.22	0.19 *

Table 6: Continued

Peak Compound (Isomers specified if known)	Δm	Standard Deviation	*Used in correlation
m/z 230 (Dimethylpyrenes and isomers)			
230a	0.02	0.46	*
230b	0.32	0.32	
230c	-0.39	0.72	
230d	0.13	0.23	*
230e	-0.21	0.70	*
230f	0.04	0.36	*
230g	-0.12	0.16	*
230h	-0.17	0.18	*
230i	-0.41	0.32	
m/z 228 (Chrysene and benzo[a]anthracene)			
BAN benzo[a]anthracene	-0.78	0.36	
CHR chrysene	-0.04	0.30	*
m/z 242 (Methylchrysenes and isomers)			
242a methylchrysene	0.17	0.37	*
242b methylchrysene	0.04	0.40	*
242c methylbenzo[a]anthracene (?)	-0.22	0.29	*
242d methylchrysene	-0.13	0.37	*
m/z 252 (Benzo[a]pyrene and isomers)			
252a benzo[<i>h</i>]fluoranthene	-0.53	0.42	
252b benzo[a]pyrene	-0.49	0.43	
252c benzo[<i>e</i>]pyrene	-1.54	0.85	

Table — Continued																	
ALMAF-4R	ALMAF-5R	COHASA-2R	COHASD-4R	DEMASG-3R	MIGRAN-3R	NTRIUG-3R	NTRIUG-4R	PENOB-2R	PENOB-4R	SDESBO-3R	SSALB-3R	SVENTO-6R	THEBAC-3R	VENTUB-5R	WCHEBK-5R	WHYCON-3R	Sample
0.32	0.20	0.82	0.16	0.83	0.78	0.70	0.22	0.64	0.85	0.89	0.86	0.07	0.88	0.87	0.11	0.52	ARCADI-50
0.15	0.06	0.63	0.02	0.46	0.50	0.38	0.08	0.53	0.54	0.66	0.60	0.01	0.57	0.52	0.02	0.26	BALMOM-30
0.21	0.11	0.74	0.03	0.59	0.63	0.52	0.15	0.68	0.66	0.72	0.64	0.05	0.62	0.55	0.05	0.37	BANQUC-10
0.33	0.16	0.80	0.09	0.80	0.78	0.67	0.22	0.64	0.83	0.92	0.88	0.05	0.91	0.87	0.11	0.54	BLUENG-80
0.29	0.19	0.86	0.09	0.75	0.76	0.67	0.23	0.76	0.81	0.84	0.78	0.06	0.77	0.73	0.10	0.50	CHEBUK-40
0.21	0.09	0.71	0.02	0.55	0.63	0.49	0.14	0.66	0.62	0.75	0.65	0.04	0.62	0.56	0.05	0.37	CITNAI-30
0.13	0.06	0.61	0.04	0.37	0.44	0.34	0.07	0.57	0.44	0.54	0.46	0.02	0.45	0.40	0.01	0.20	COHASA-50
0.16	0.11	0.63	0.06	0.42	0.48	0.36	0.11	0.68	0.47	0.52	0.46	0.02	0.48	0.43	0.03	0.28	COHASD-70
0.22	0.12	0.77	0.05	0.61	0.65	0.54	0.16	0.71	0.68	0.73	0.64	0.04	0.63	0.59	0.04	0.37	GLENEI-80
0.29	0.16	0.84	0.08	0.72	0.74	0.64	0.21	0.75	0.78	0.83	0.76	0.06	0.75	0.71	0.08	0.47	NTRIUB-40
0.30	0.17	0.80	0.09	0.74	0.74	0.62	0.21	0.72	0.78	0.85	0.79	0.05	0.83	0.78	0.09	0.50	OLYMFA-50
0.17	0.10	0.46	0.11	0.34	0.39	0.28	0.08	0.54	0.35	0.40	0.36	0.13	0.36	0.32	0.03	0.27	PANUKB-10
0.17	0.29	0.60	0.05	0.41	0.44	0.49	0.27	0.54	0.54	0.55	0.51	0.02	0.40	0.39	0.16	0.40	PANUKJ-10
0.29	0.14	0.79	0.06	0.67	0.72	0.58	0.21	0.70	0.72	0.87	0.78	0.04	0.79	0.75	0.09	0.48	PENOB-50
0.14	0.08	0.61	0.04	0.38	0.44	0.36	0.10	0.59	0.45	0.50	0.44	0.02	0.44	0.38	0.01	0.22	SABLEH-50
0.33	0.21	0.83	0.10	0.74	0.75	0.64	0.25	0.75	0.79	0.83	0.79	0.06	0.83	0.77	0.11	0.51	SVENTO-10
0.37	0.47	0.77	0.24	0.78	0.75	0.73	0.43	0.67	0.81	0.80	0.77	0.06	0.79	0.85	0.31	0.68	THEBAC-90
0.36	0.20	0.83	0.10	0.76	0.78	0.65	0.24	0.76	0.80	0.85	0.82	0.07	0.86	0.81	0.11	0.54	UNIACG-60
0.32	0.45	0.78	0.17	0.69	0.67	0.76	0.44	0.62	0.80	0.73	0.71	0.06	0.65	0.66	0.29	0.59	VENTUH-50
0.33	0.21	0.83	0.13	0.82	0.80	0.69	0.24	0.70	0.84	0.91	0.85	0.05	0.88	0.88	0.13	0.57	VENTUH-70
0.34	0.21	0.85	0.12	0.84	0.80	0.75	0.27	0.65	0.90	0.94	0.90	0.06	0.89	0.88	0.15	0.57	ABENAJ-3R
0.39	0.28	0.93	0.19	0.92	0.87	0.86	0.32	0.73	0.95	0.93	0.92	0.13	0.87	0.84	0.19	0.66	ALMAF-4R
0.57	0.57	0.44	0.21	0.54	0.62	0.64	0.81	0.54	0.53	0.40	0.45	0.56	0.51	0.56	0.68	0.79	ALMAF-5R
0.44	0.33	0.33	0.38	0.40	0.41	0.54	0.82	0.45	0.40	0.23	0.28	0.30	0.30	0.31	0.38	0.76	COHASA-2R
0.44	0.33	0.27	0.27	0.86	0.88	0.85	0.38	0.77	0.90	0.86	0.87	0.18	0.82	0.79	0.25	0.67	COHASD-4R
0.21	0.38	0.27	0.27	0.26	0.26	0.31	0.25	0.14	0.22	0.12	0.20	0.26	0.20	0.21	0.29	0.26	DEMASG-3R
0.54	0.40	0.86	0.26	0.93	0.93	0.92	0.44	0.73	0.97	0.85	0.88	0.25	0.89	0.87	0.32	0.78	MIGRAN-3R
0.62	0.41	0.88	0.26	0.93	0.90	0.90	0.53	0.80	0.91	0.87	0.86	0.29	0.86	0.85	0.40	0.81	NTRIUG-3R
0.64	0.54	0.85	0.31	0.92	0.90	0.90	0.61	0.70	0.93	0.79	0.83	0.34	0.79	0.79	0.47	0.84	NTRIUG-4R
0.81	0.82	0.38	0.25	0.44	0.53	0.61	0.46	0.46	0.45	0.32	0.36	0.40	0.39	0.46	0.90	0.79	PENOB-2R
0.54	0.45	0.77	0.14	0.73	0.80	0.70	0.46	0.73	0.73	0.65	0.59	0.20	0.64	0.65	0.28	0.73	PENOB-4R
0.53	0.40	0.90	0.22	0.97	0.91	0.93	0.45	0.65	0.89	0.89	0.91	0.21	0.89	0.88	0.31	0.77	SDESBO-3R
0.40	0.23	0.86	0.12	0.85	0.87	0.79	0.32	0.65	0.89	0.95	0.95	0.12	0.90	0.88	0.22	0.65	SSALB-3R
0.45	0.28	0.87	0.20	0.88	0.86	0.83	0.36	0.59	0.91	0.95	0.95	0.16	0.94	0.89	0.27	0.68	SVENTO-6R
0.56	0.30	0.18	0.26	0.25	0.29	0.34	0.40	0.20	0.21	0.12	0.16	0.15	0.15	0.14	0.43	0.35	THEBAC-3R
0.51	0.31	0.82	0.20	0.89	0.86	0.79	0.39	0.64	0.89	0.90	0.94	0.15	0.95	0.95	0.26	0.69	VENTUB-5R
0.56	0.38	0.79	0.21	0.87	0.85	0.79	0.46	0.65	0.88	0.88	0.89	0.14	0.95	0.95	0.26	0.69	WCHEBK-5R
0.68	0.76	0.25	0.29	0.32	0.40	0.47	0.90	0.28	0.31	0.22	0.27	0.43	0.26	0.33	0.69	0.69	WHYCON-3R
0.79	0.71	0.67	0.26	0.78	0.81	0.84	0.79	0.73	0.77	0.65	0.68	0.35	0.69	0.76	0.69	0.69	

Table 7

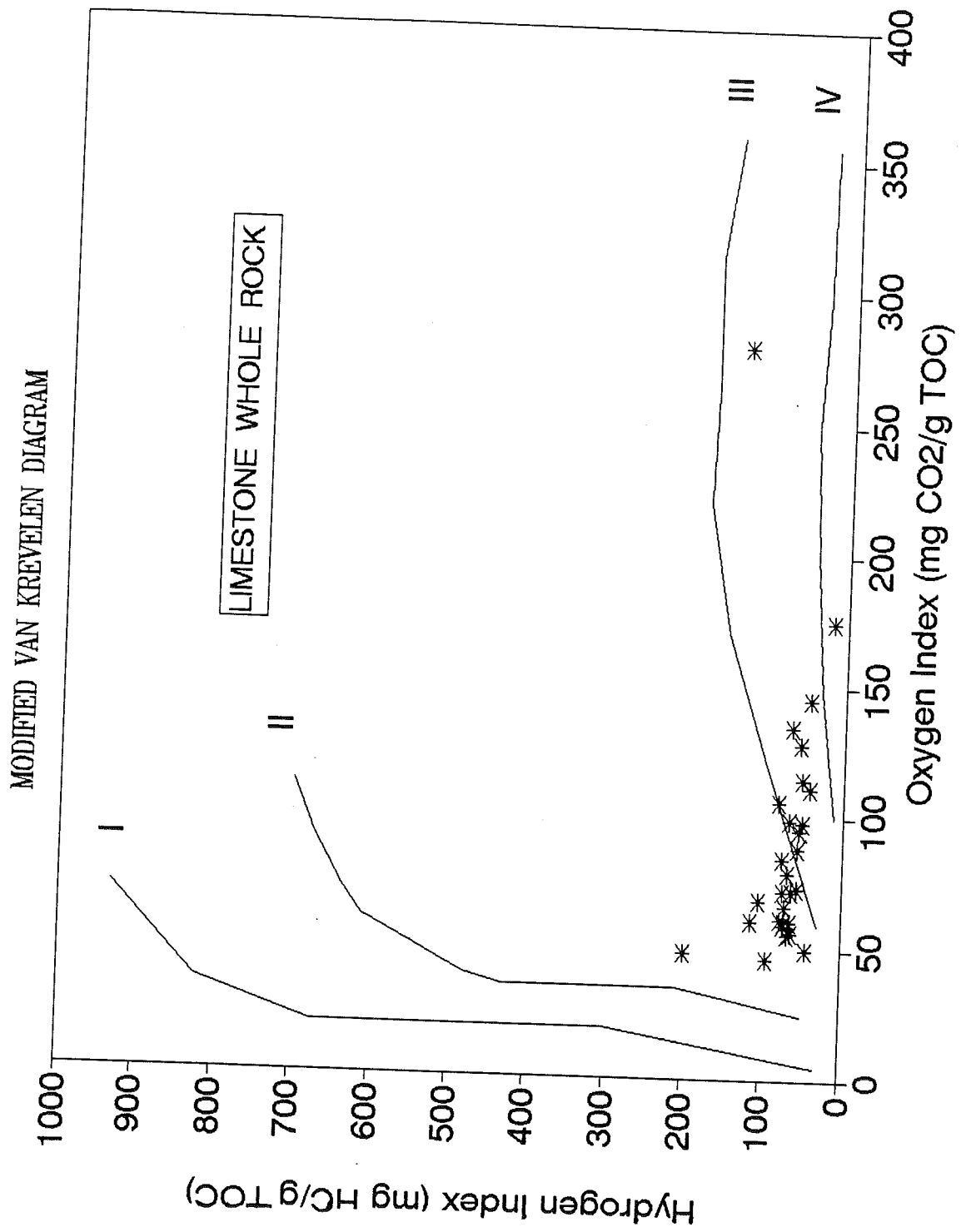


Figure 1A

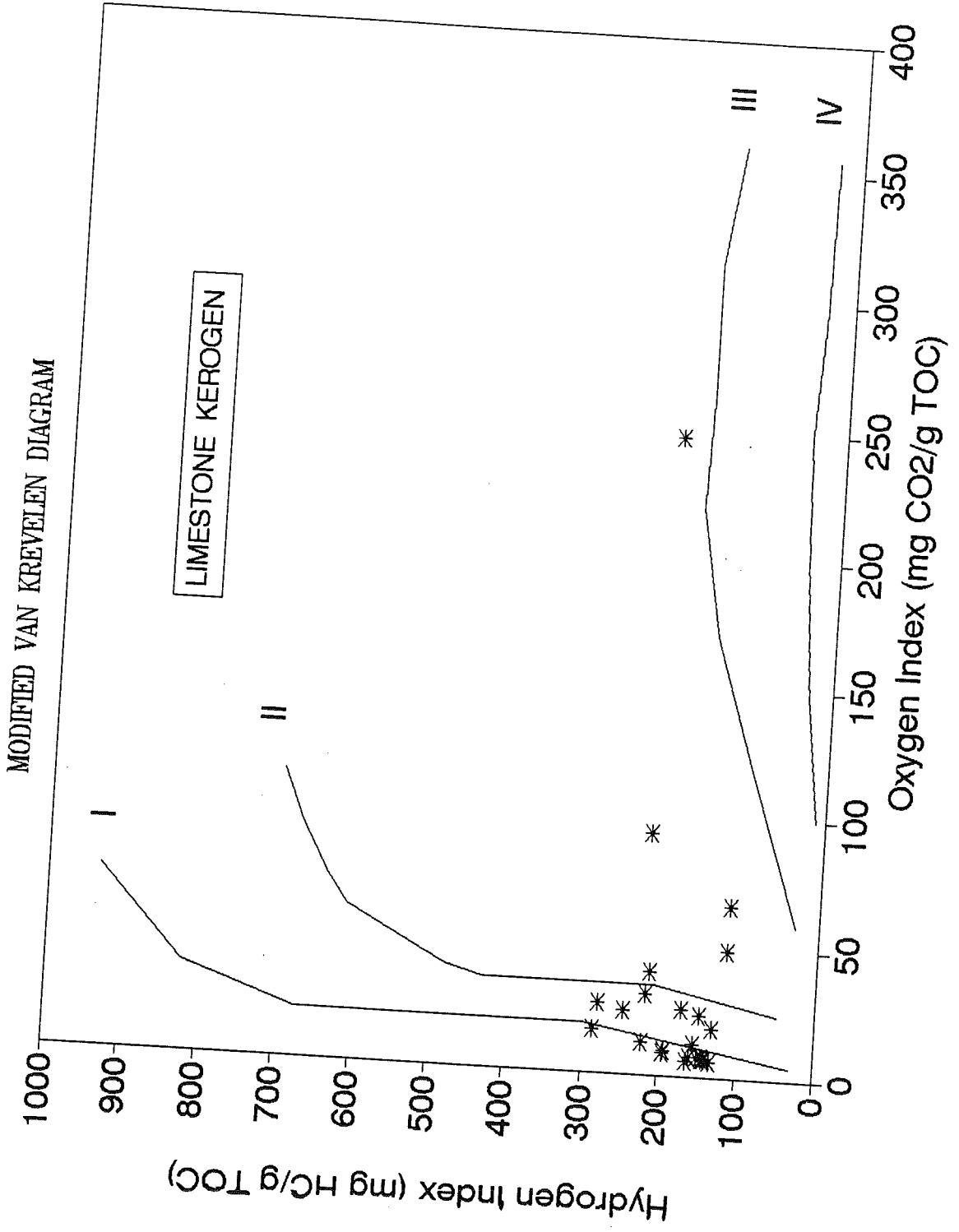


Figure 1B

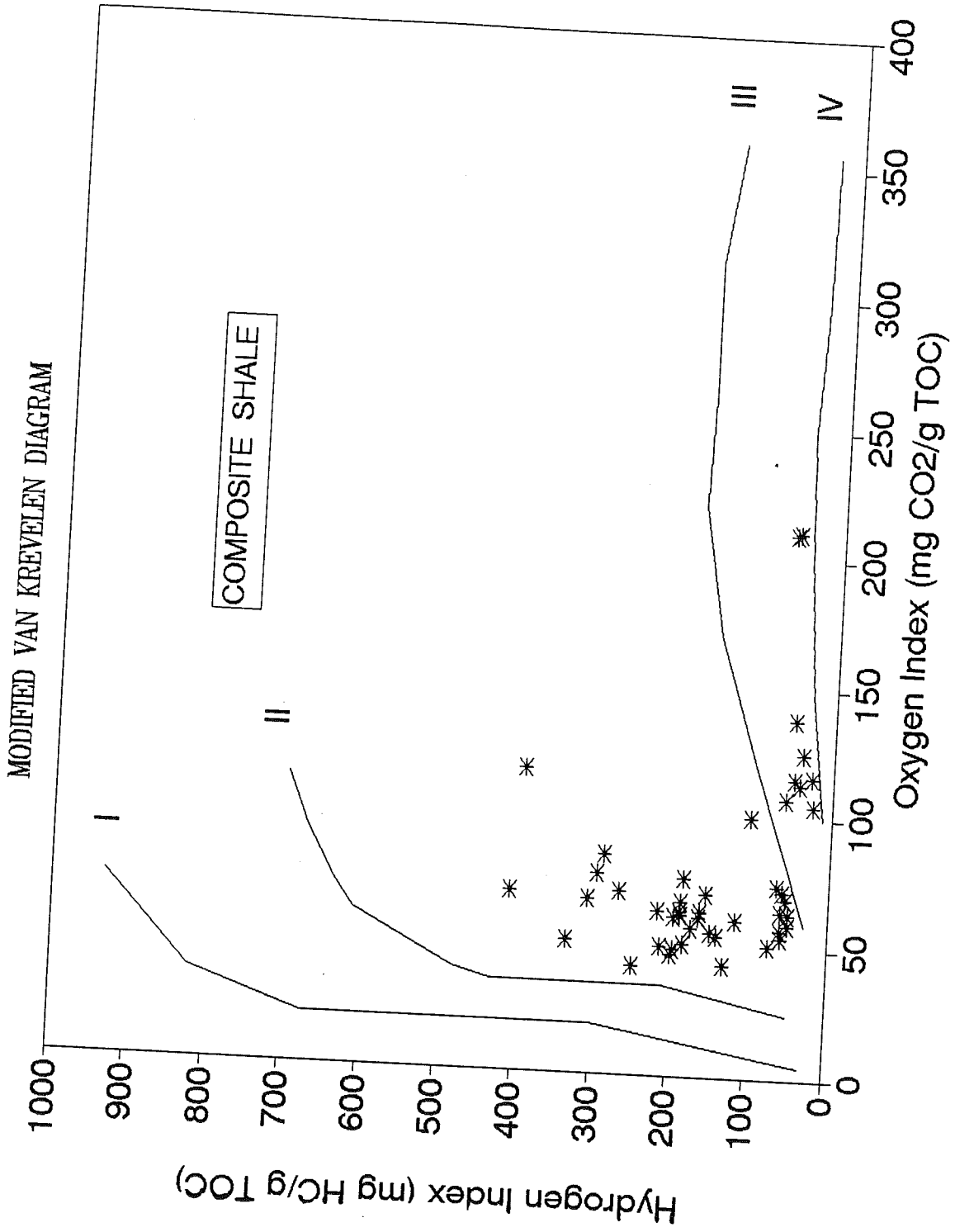


Figure 1C

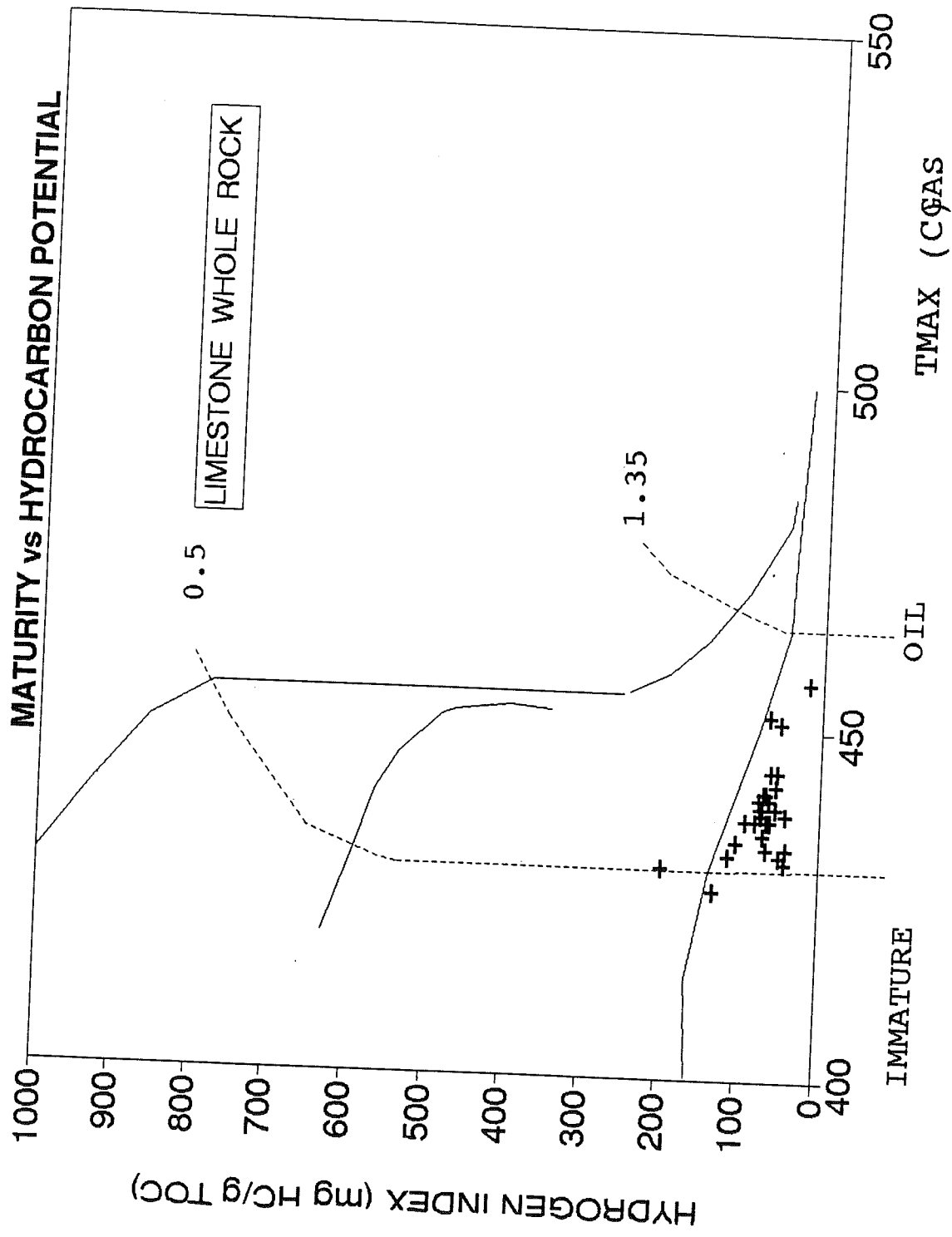


Figure 2A

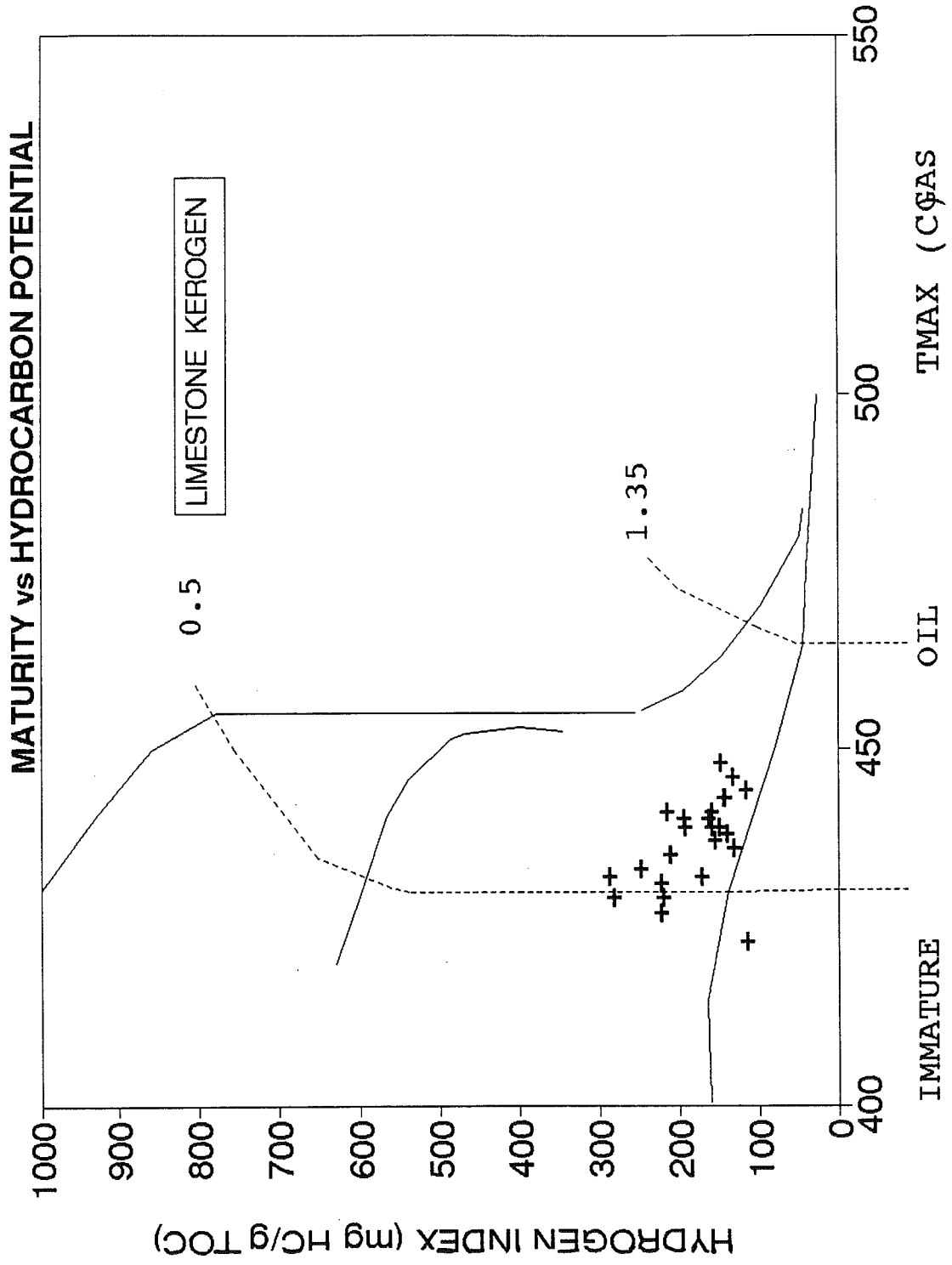


Figure 2B

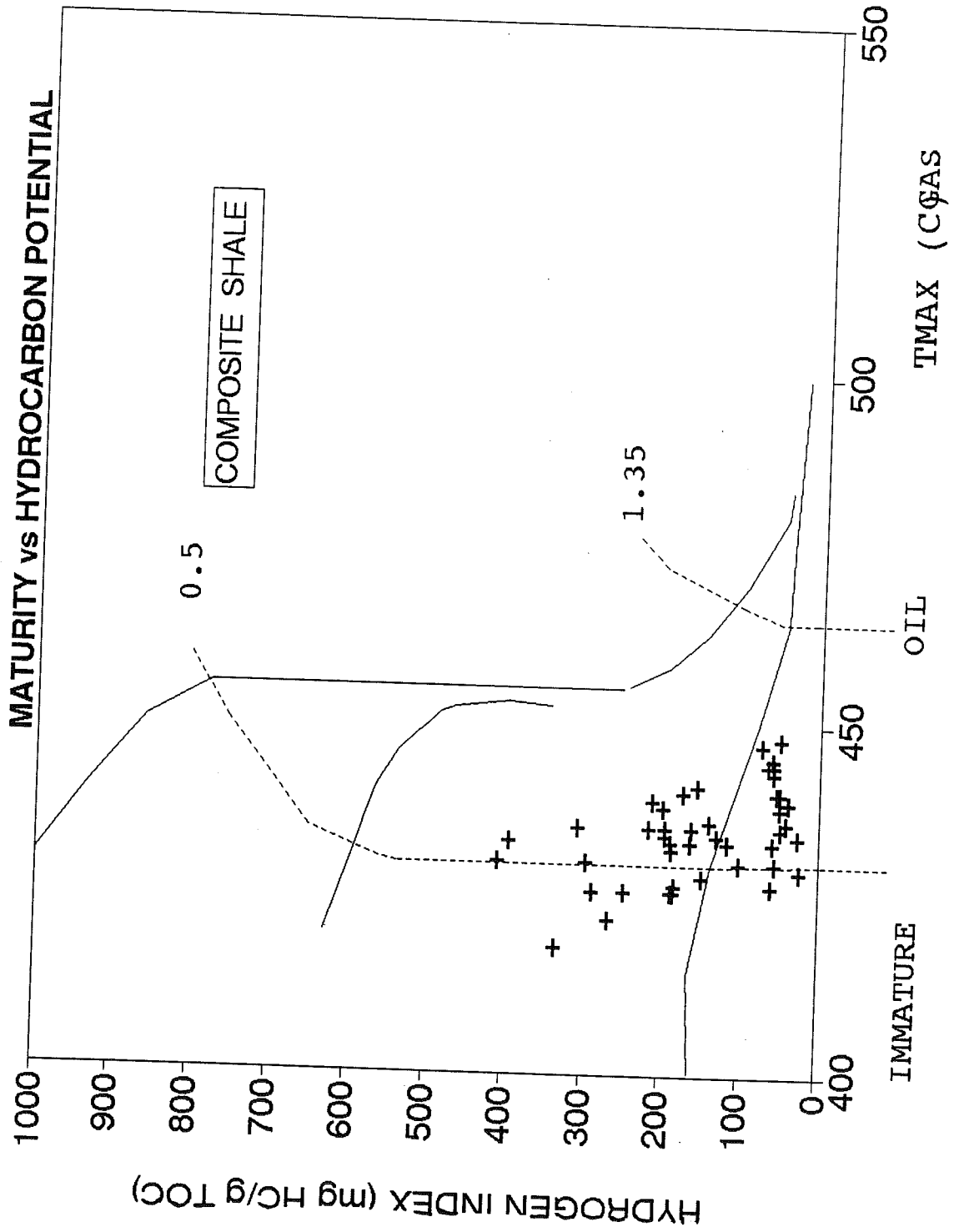


Figure 2C

ROCK-EVAL DERIVED KINETIC PARAMETERS GREEN RIVER SHALE (AP22)

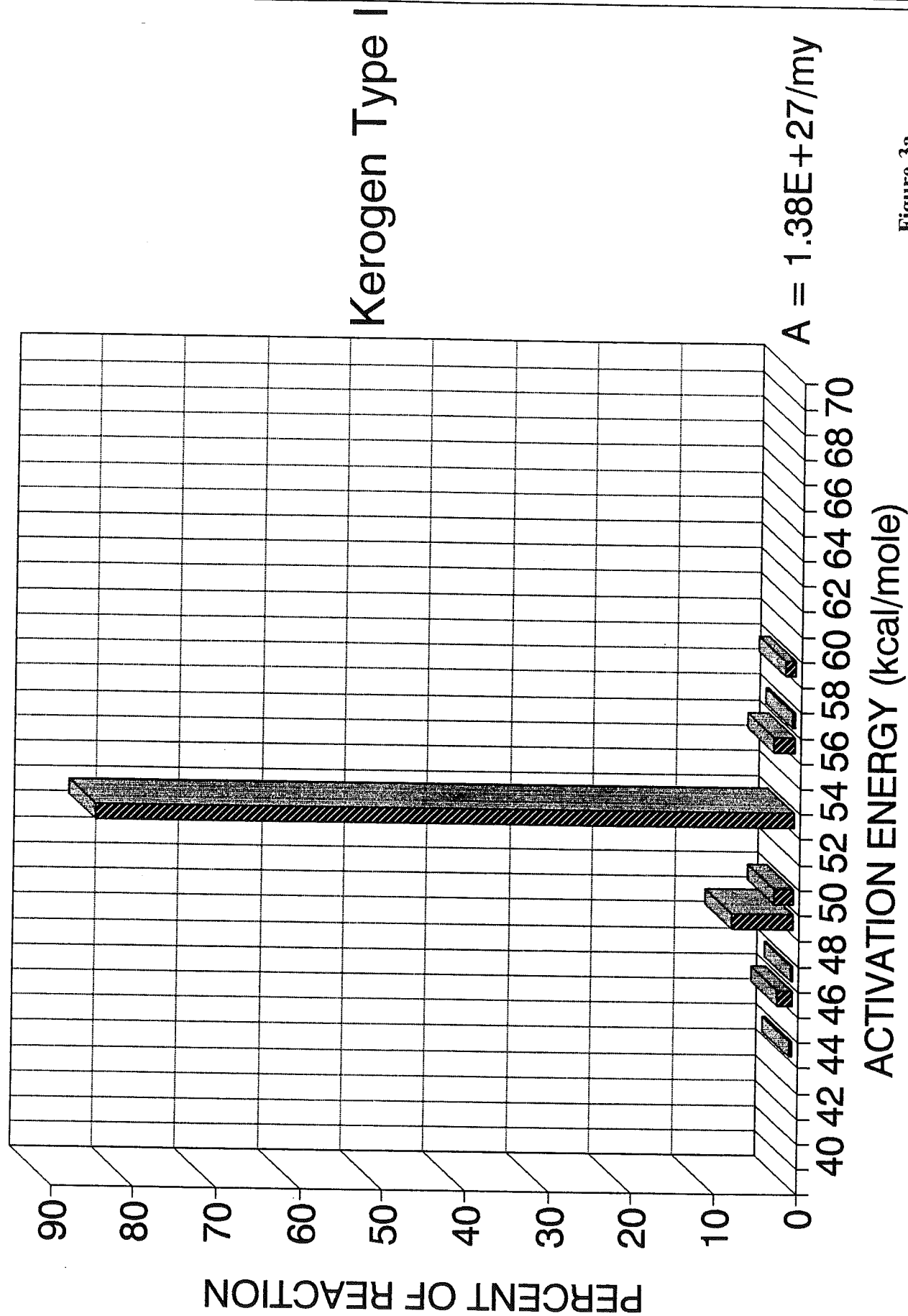


Figure 3a

ROCK-EVAL DERIVED KINETIC PARAMETERS

ABENAKI J-56 10090 - 10130 ft

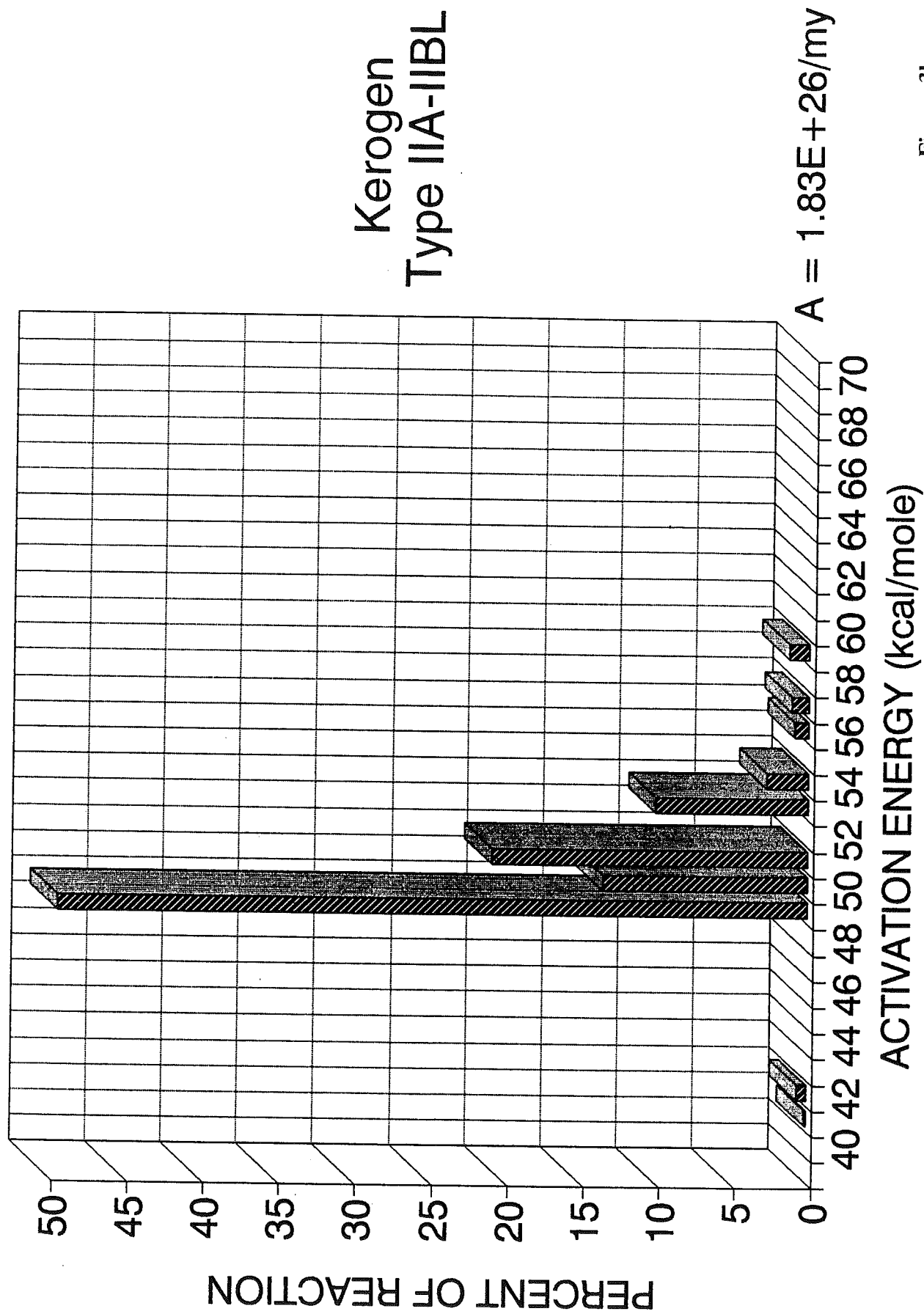


Figure 3b

ROCK-EVAL DERIVED KINETIC PARAMETERS

DEMOSCOTA G-43 11840 - 11880 ft

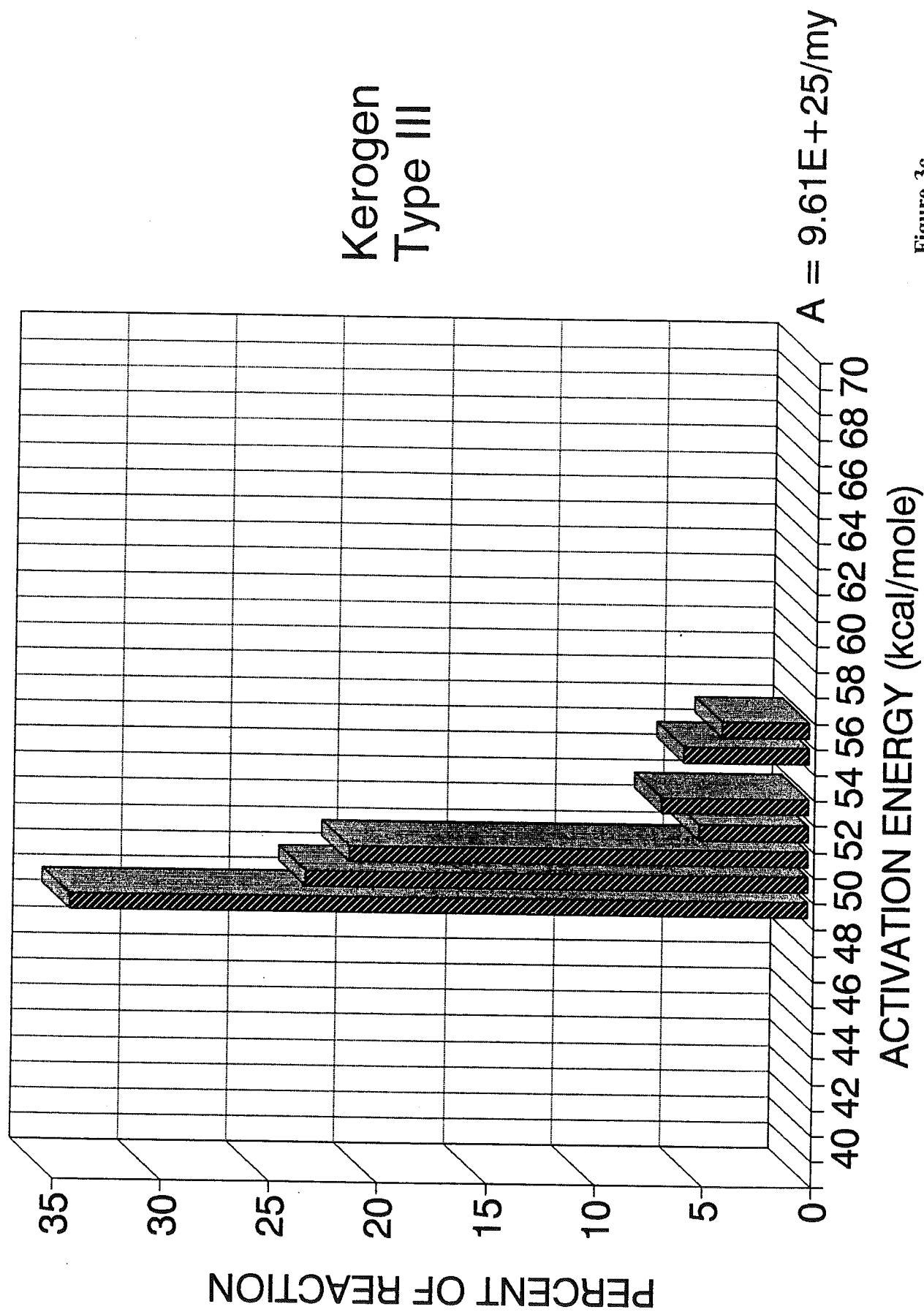


Figure 3c

ROCK-EVAL DERIVED KINETIC PARAMETERS

NORTH TRIUMPH B-52 3560 m

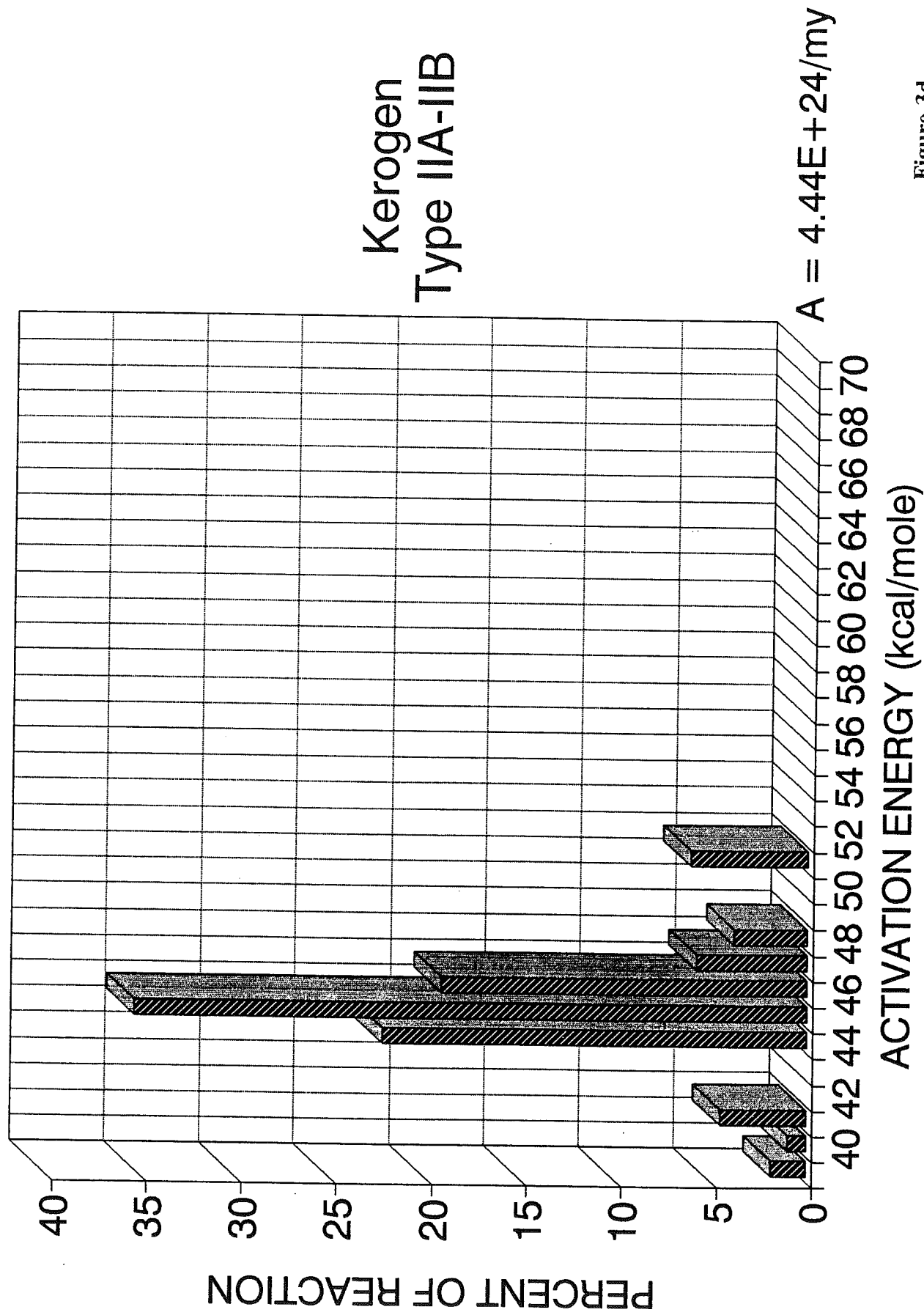


Figure 3d

ROCK-EVAL DERIVED KINETIC PARAMETERS

VENTURE B-52 3680 m

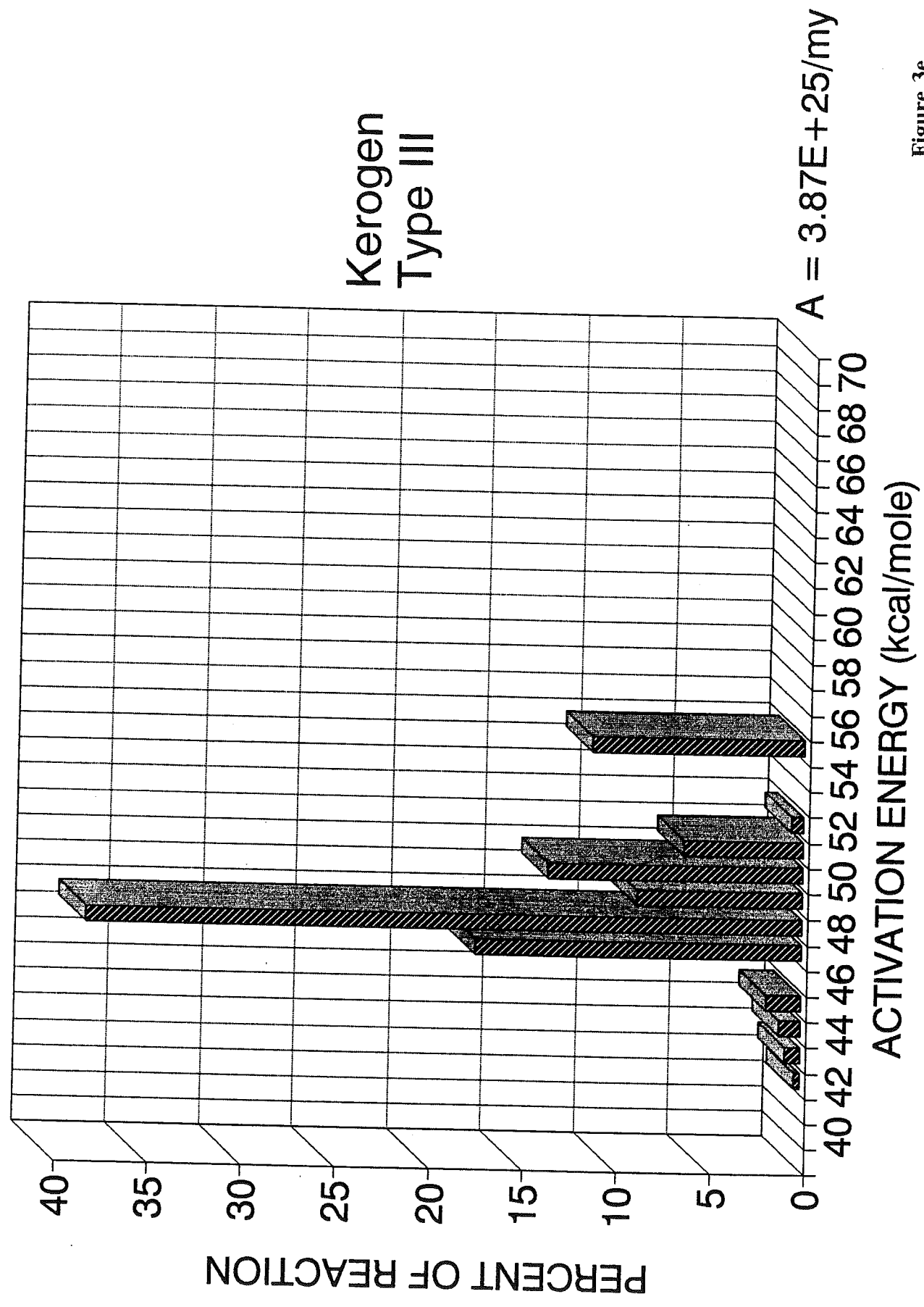
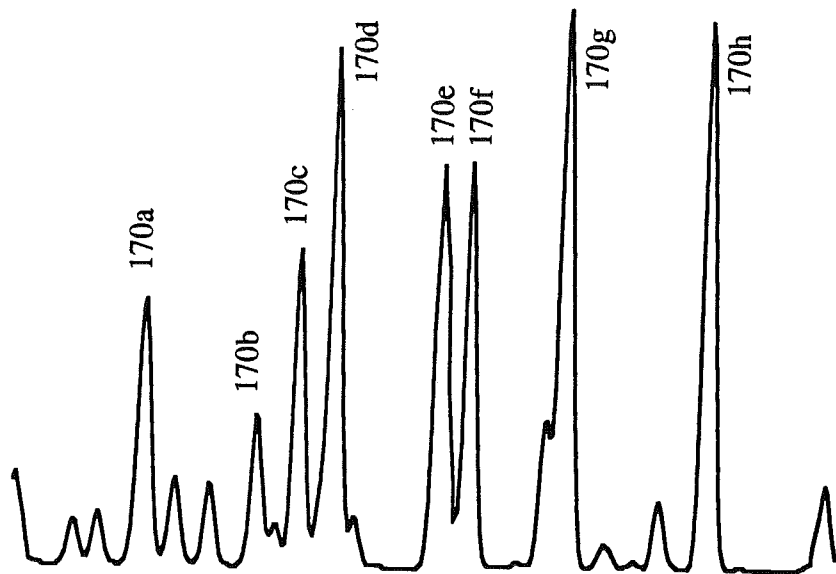
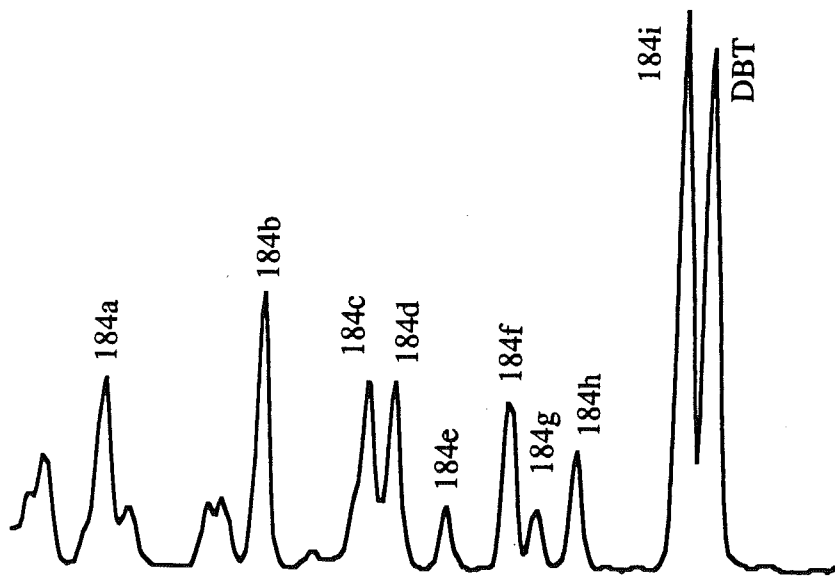


Figure 3e



M/Z 170
C3-ALKYLNAPHTHALENES



M/Z 184
C4-ALKYLNAPHTHALENES
and DIBENZOTHIOPHENE

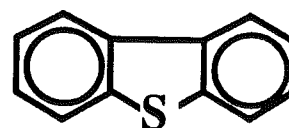
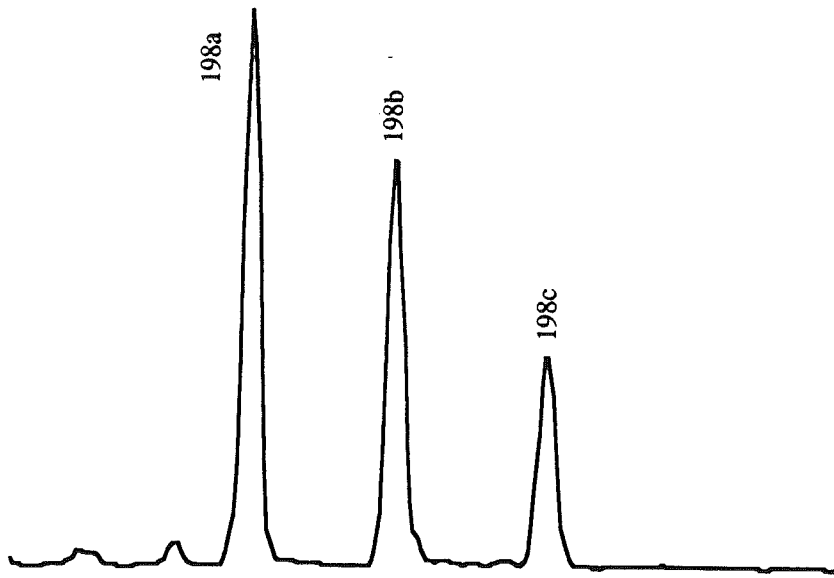
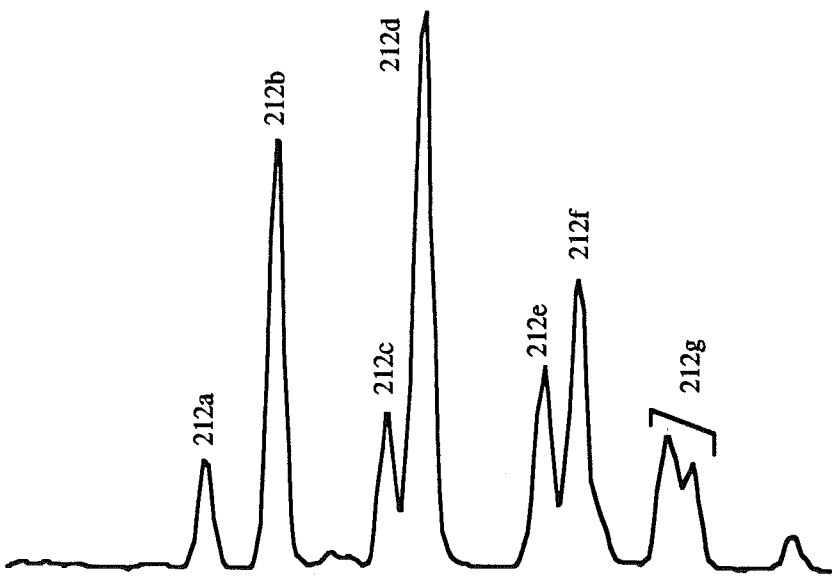
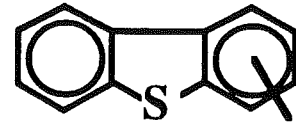


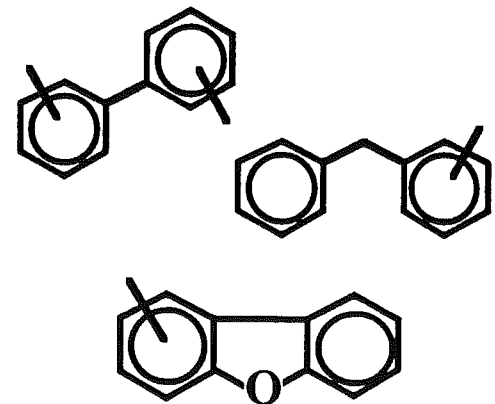
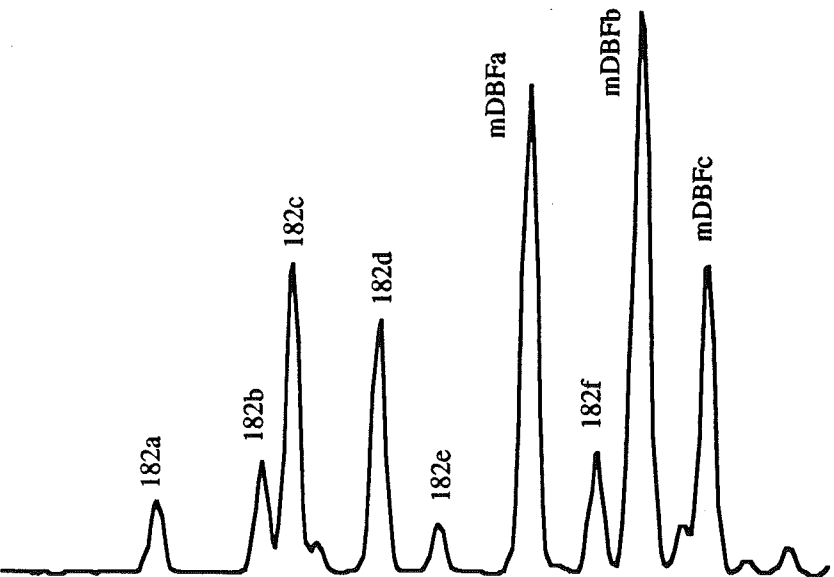
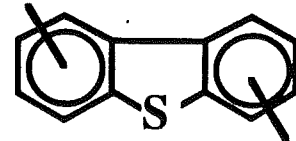
Figure 5



**M/Z 198
METHYLDIBENZOTHIOPHENES**



**M/Z 212
DIMETHYLDIBENZOTHIOPHENES**

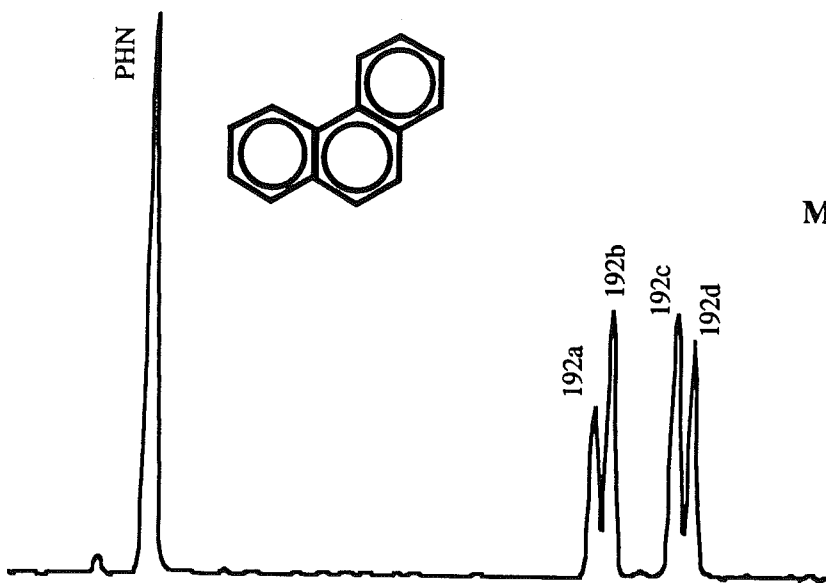
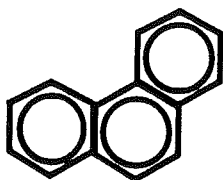


**M/Z 182
DIMETHYLBIPHENYLS
METHYLDIPHENYLMETHANE
METHYLDIBENZOFURANS**

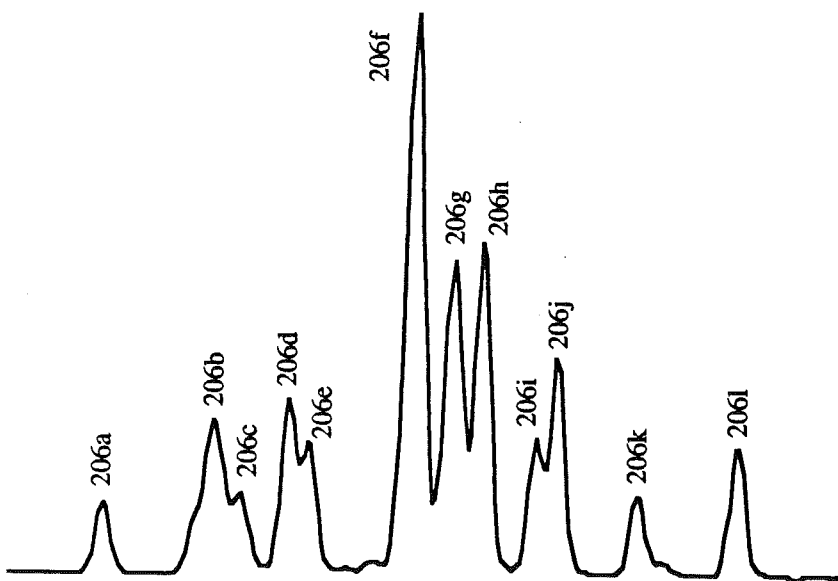
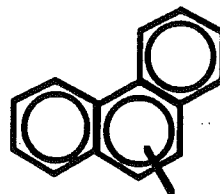
Figure 6

**M/Z 178
PHENANTHRENE**

PHN



**M/Z 192
METHYLPHENANTHRENES**



**M/Z 206
DIMETHYLPHENANTHRENES**

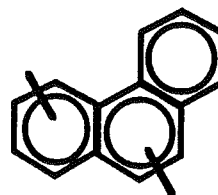
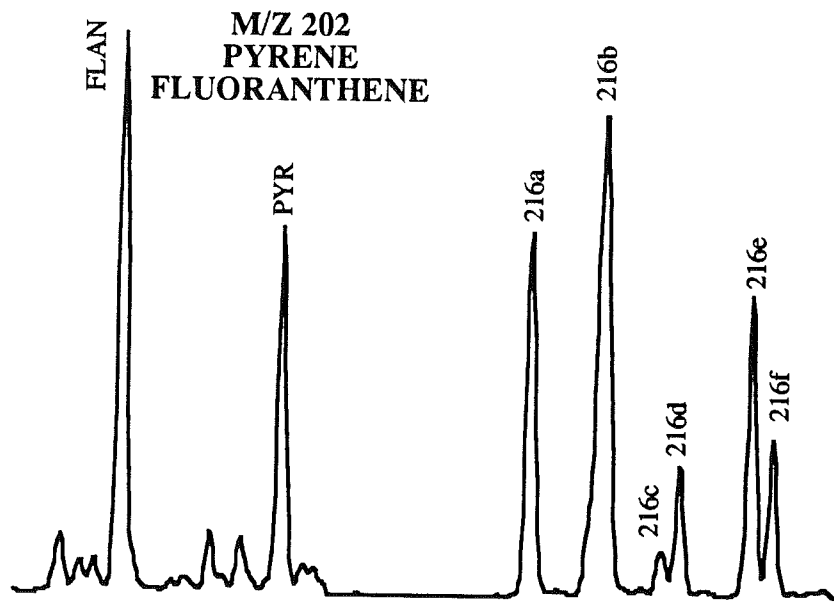
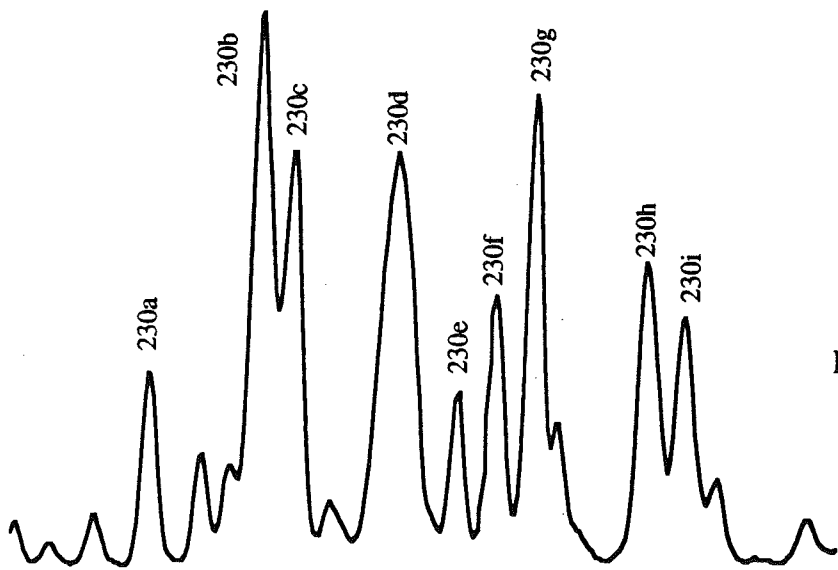
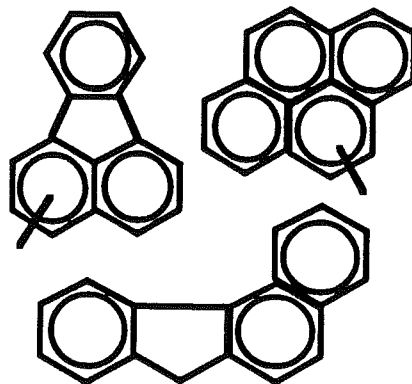


Figure 7

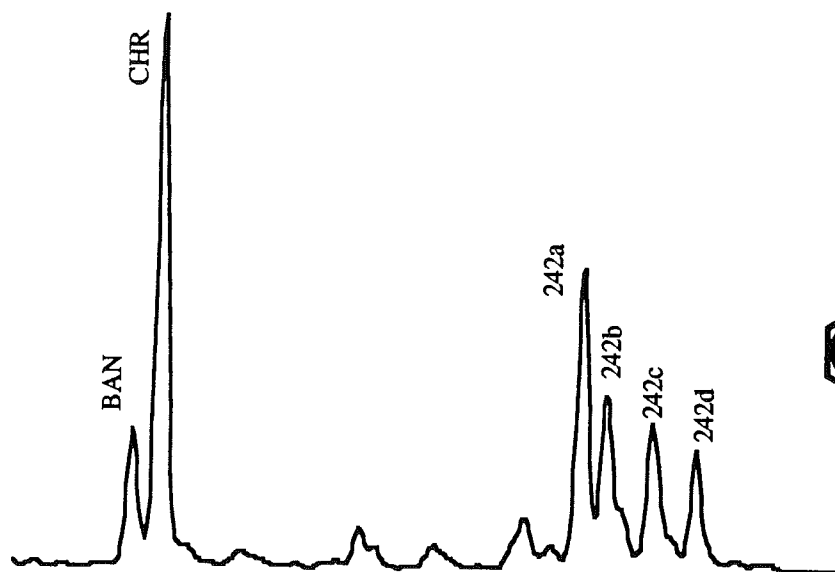


M/Z 216
METHYLPYRENES
METHYLFLUORANTHENES
BENZOFLUORENE

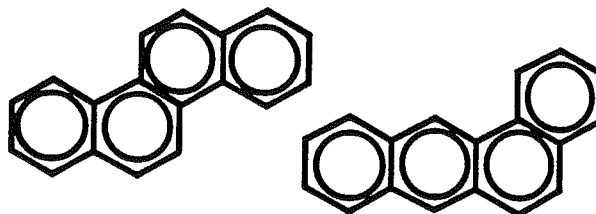


M/Z 230
DIMETHYLPYRENES
DIMETHYLFLUORANTHENES
METHYLBENZOFLUORENES

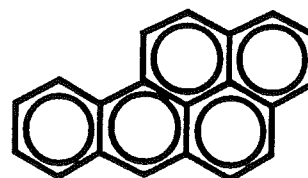
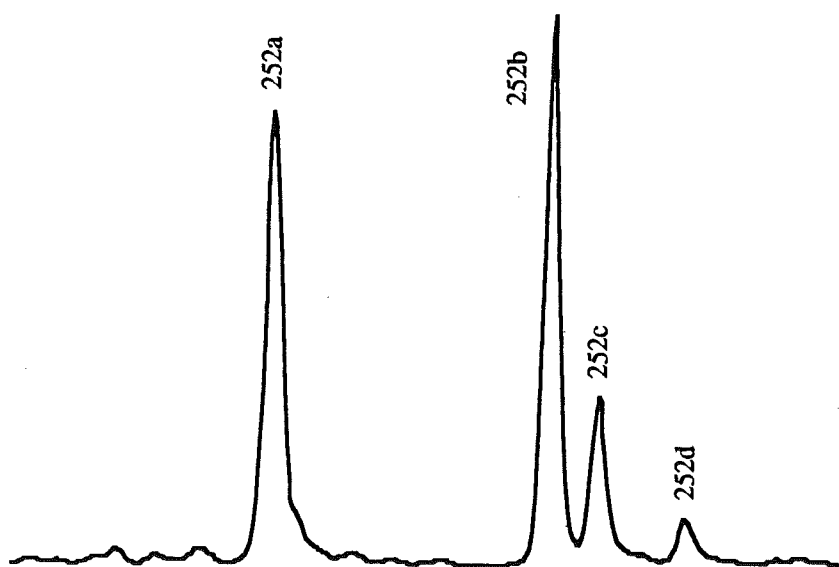
Figure 8



**M/Z 228
CHRYSENE
BENZO[a]ANTHRACENE**



**M/Z 242
METHYLCHRYSENES & ISOMERS**



**M/Z 252
BENZO[a]PYRENE & ISOMERS**

Figure 9

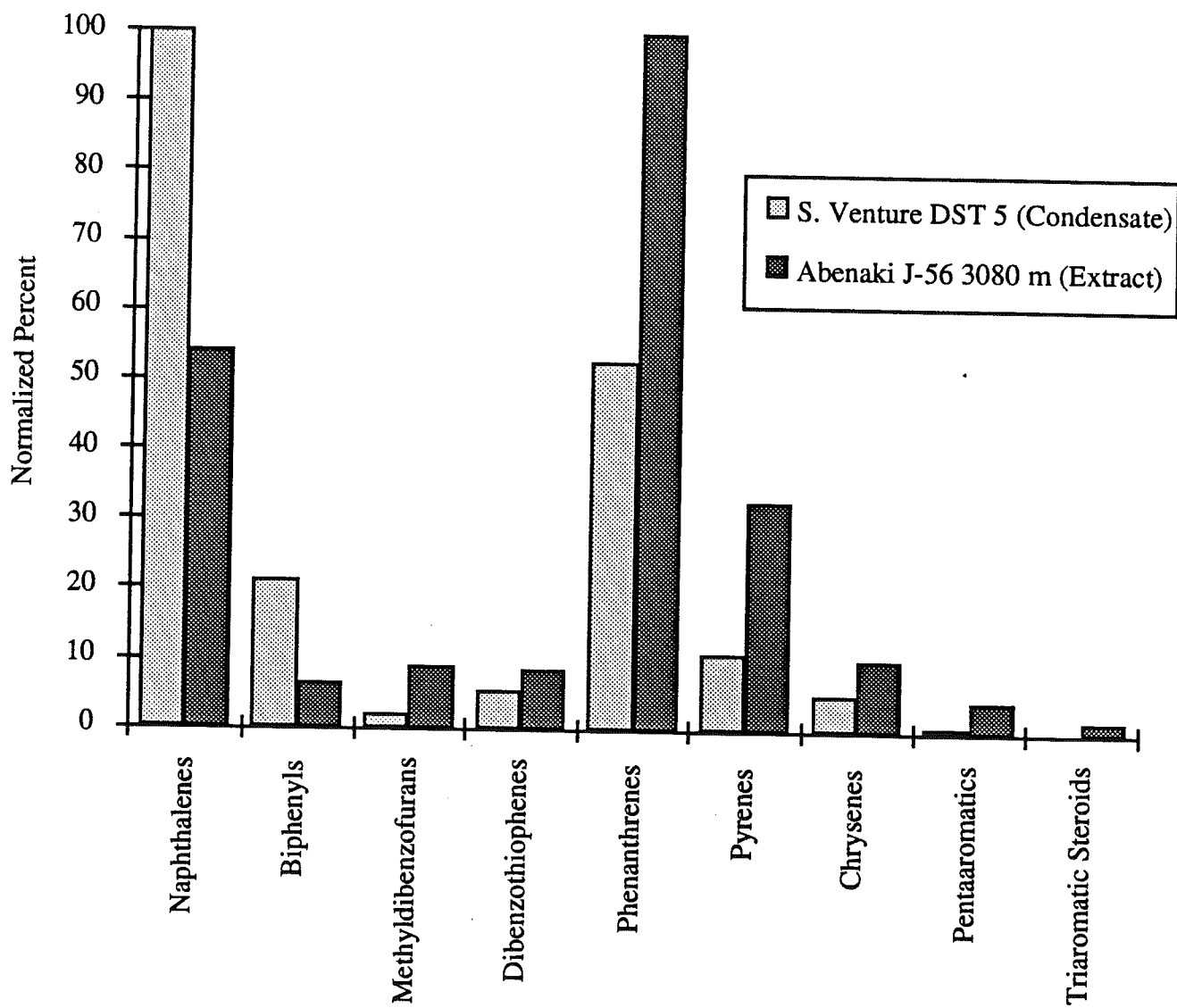


Figure 10

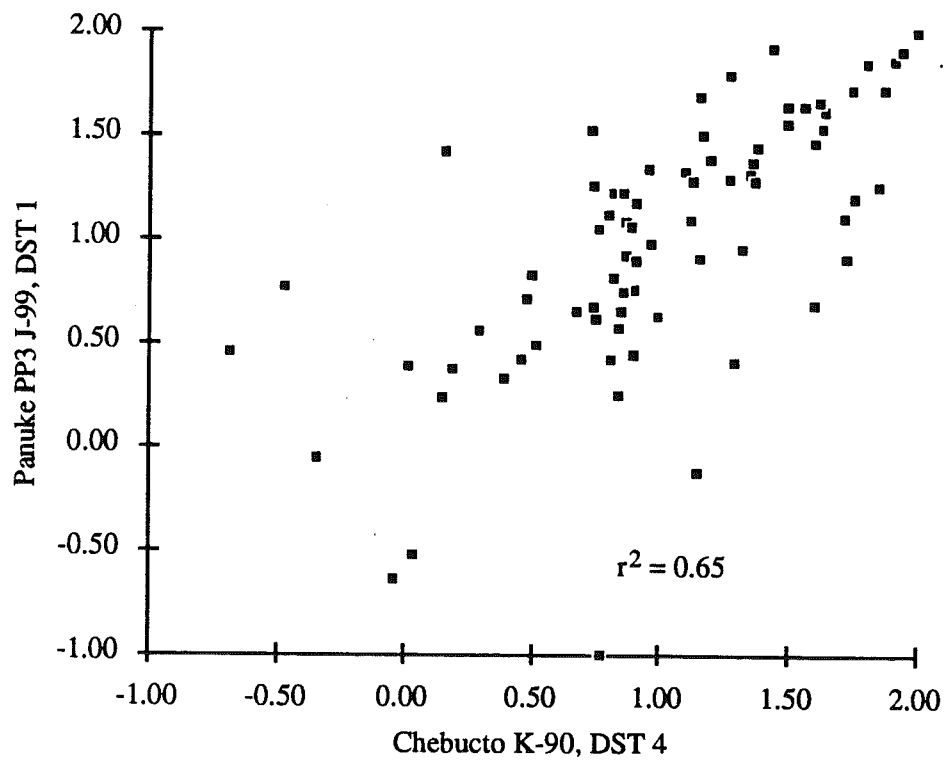
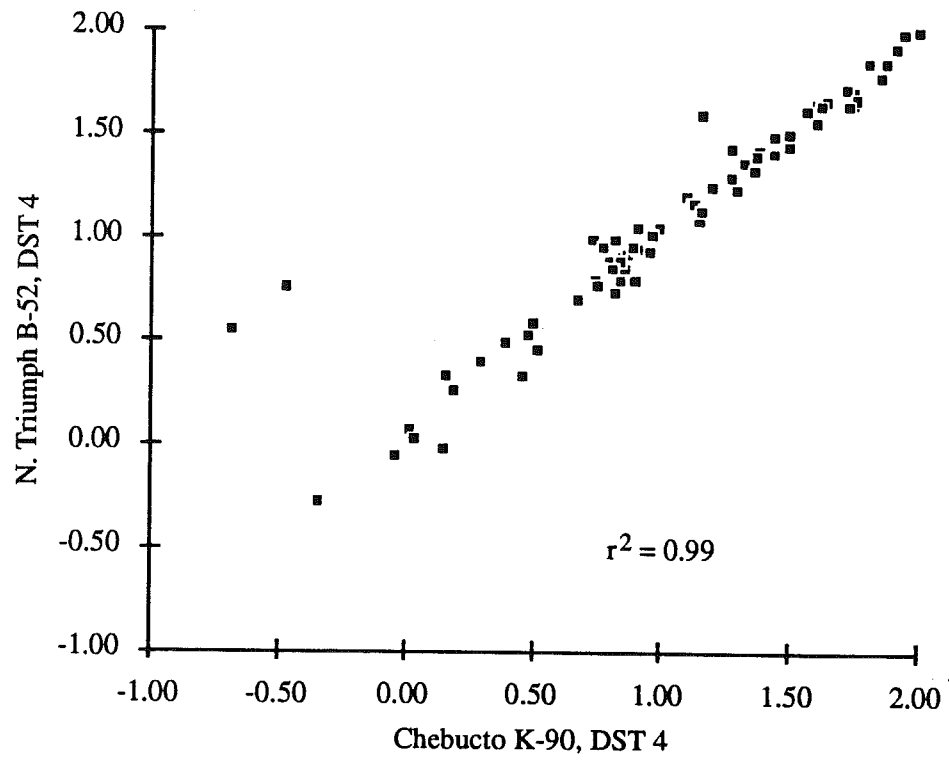


Figure 11

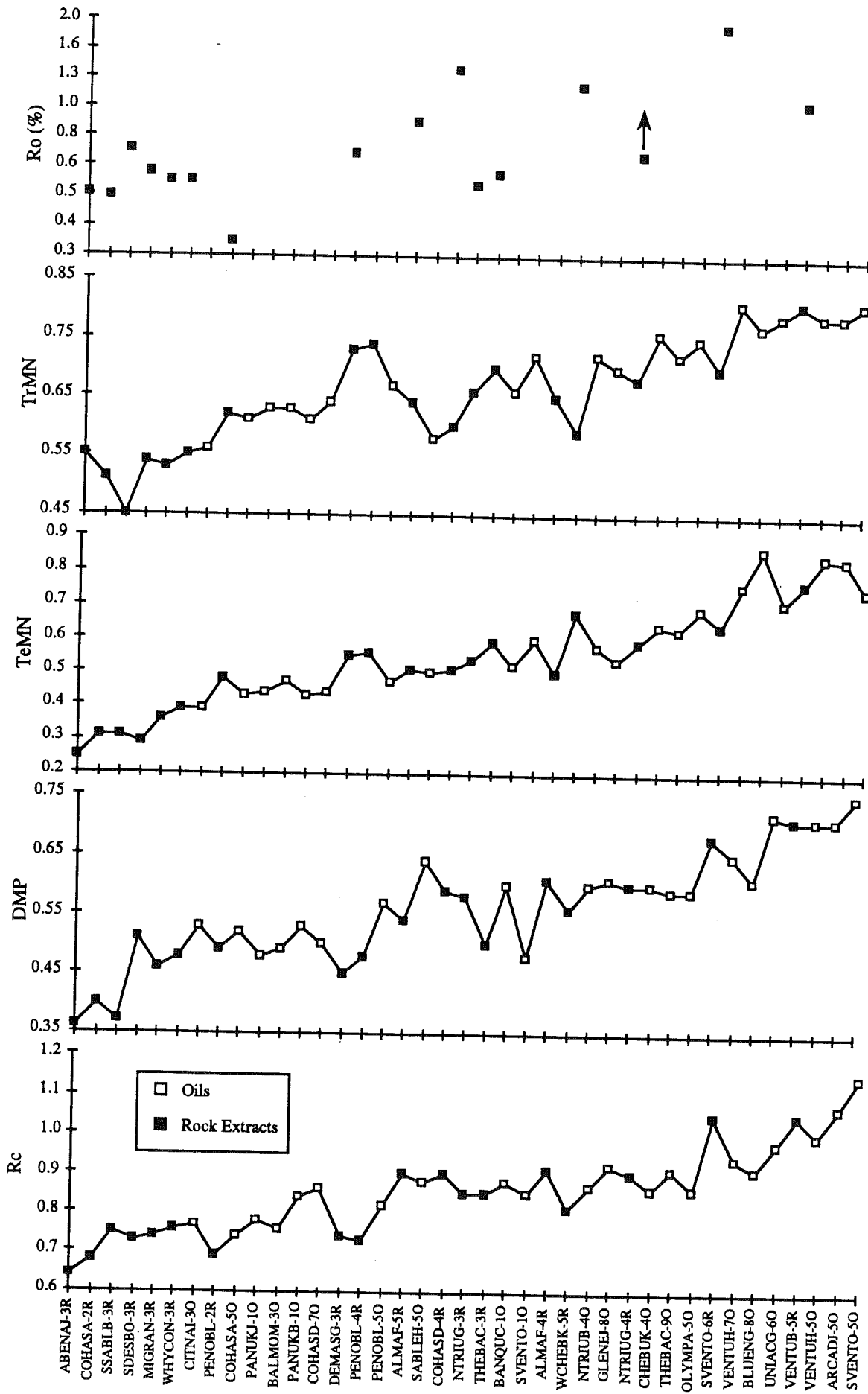


Figure 12

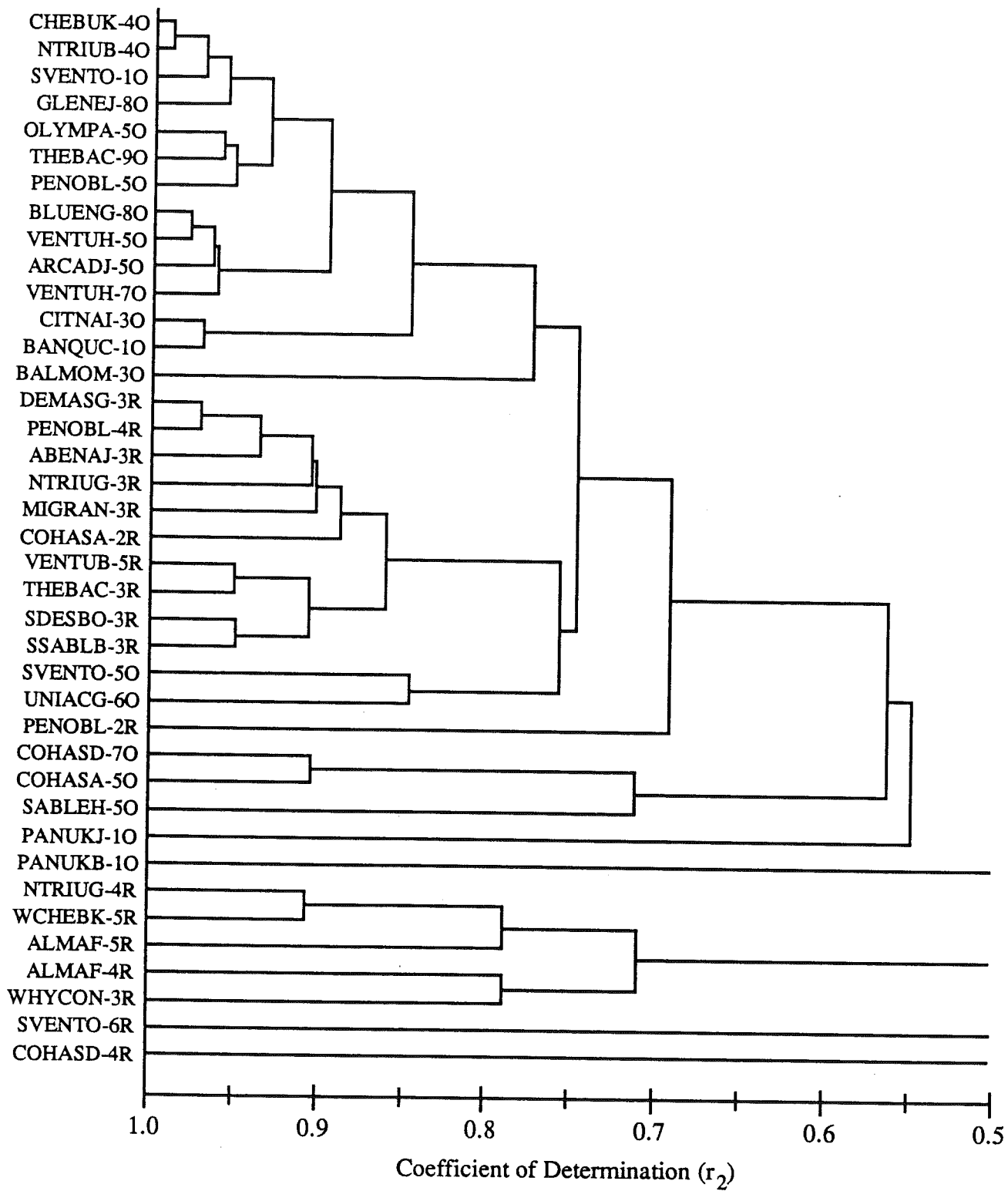
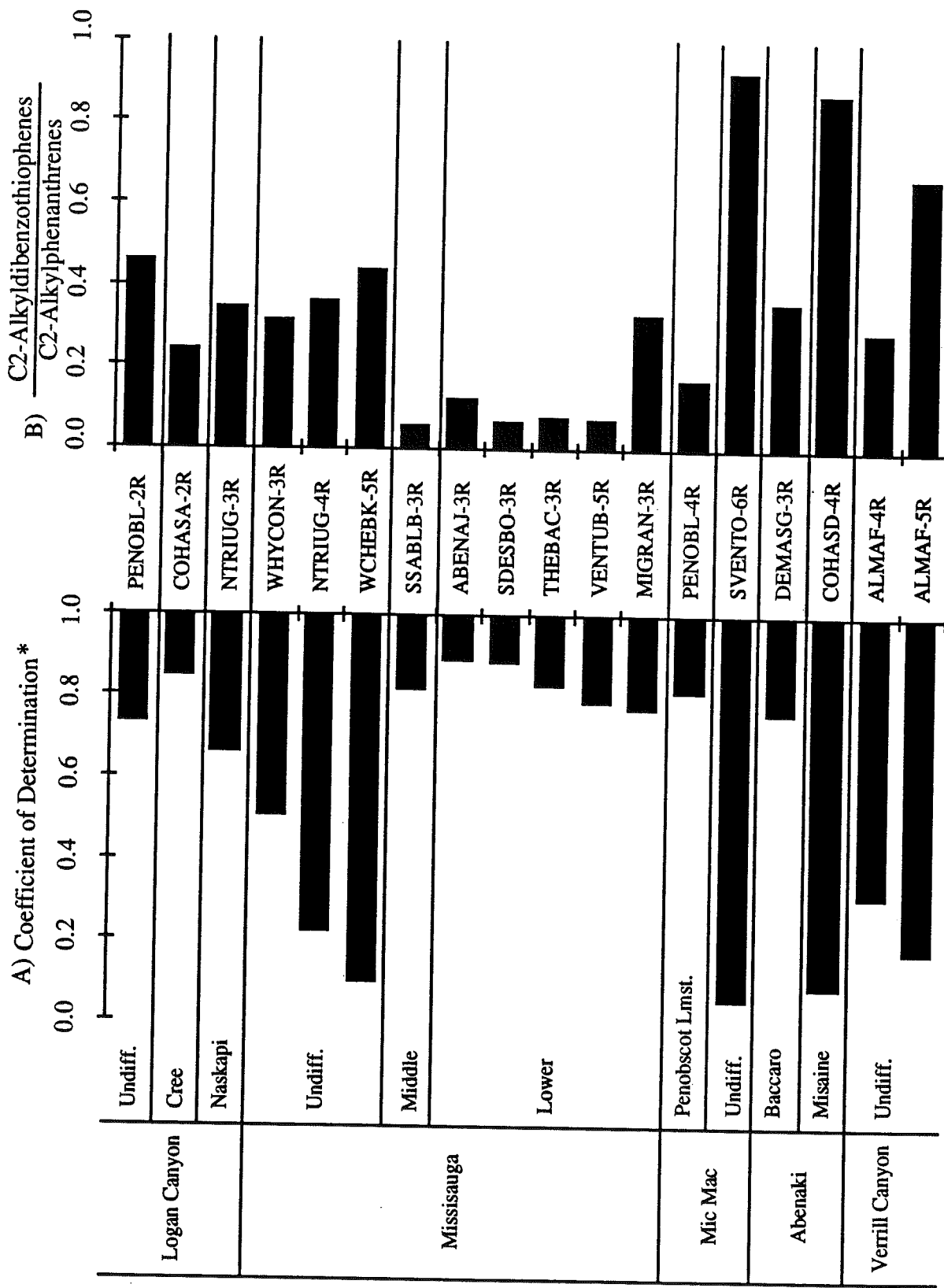


Figure 13



* From comparison of rock extracts with Group A oils

Figure 14

Well: Abenaki J-56

FM_NAME	BEGIN_AGE	TOP_DEPTH	THICKNESS	LITHOLOGY	KEROGEN_NAME	TOC
Quarternary	1.6	~107	~93	90/10/0		
Unconformity1	15					
Miocene	25	~200	~200	30/70/0		
Unconformity2	52					
E..Eocene	58	~400	~42	90/10/0		
Unconformity3	64					
E. Paleocene	66.4	~442	~143	40/60/0		
Maastrichtian	70	~585	~287	Shale	Type III	1.5
Unconformity4	80					
Campanian	88.5	~872	~78	10/90/0	Type III	1.4
Turonian	90	~950	~36	Shale	IIB	1.75
Petrel	91	~986	~74	40/45/15	Type III	0.9
Cenoman	116	~1060	~915	55/45/0	Type III	1.5
Naskapi	119	~1975	~99	40/60/0	IIB	2.5
Barrem/Miss	124	~2074	~239	85/15/0	Type III	1.0
Haut/"O"	131	~2313	~242	50/40/10	Type III	0.5
Val/Berr	144	~2555	~666	65/35/0	Type III	1.0
Top Jura	144.5	~3221	~41	20/10/70	Type III	1.0
Abenaki 1	152	~3262	~298	20/10/70	Type III	1.0
Abenaki 2	163	~3560	~372	20/10/70	IIL	1.75
L. Callovian	164	~3932	~57	0/10/90	IIIL	0.75
Misaine	166	~3989	~124	10/90/0	Type III	0.3
Scatarie	168	~4113	~174	0/10/90	IIB	0.5
Mohican	200	~4287	~3713	30/30/40	IIB	0.5
BU	201				Type III	0.5
Argo	204	~8000	~4000	0/5/95	Type III	0.2
Eurydice	220	~12000	~1500	40/50/10	Type III	10.0

Table 8

Abenaki J-56

Scotian Shelf

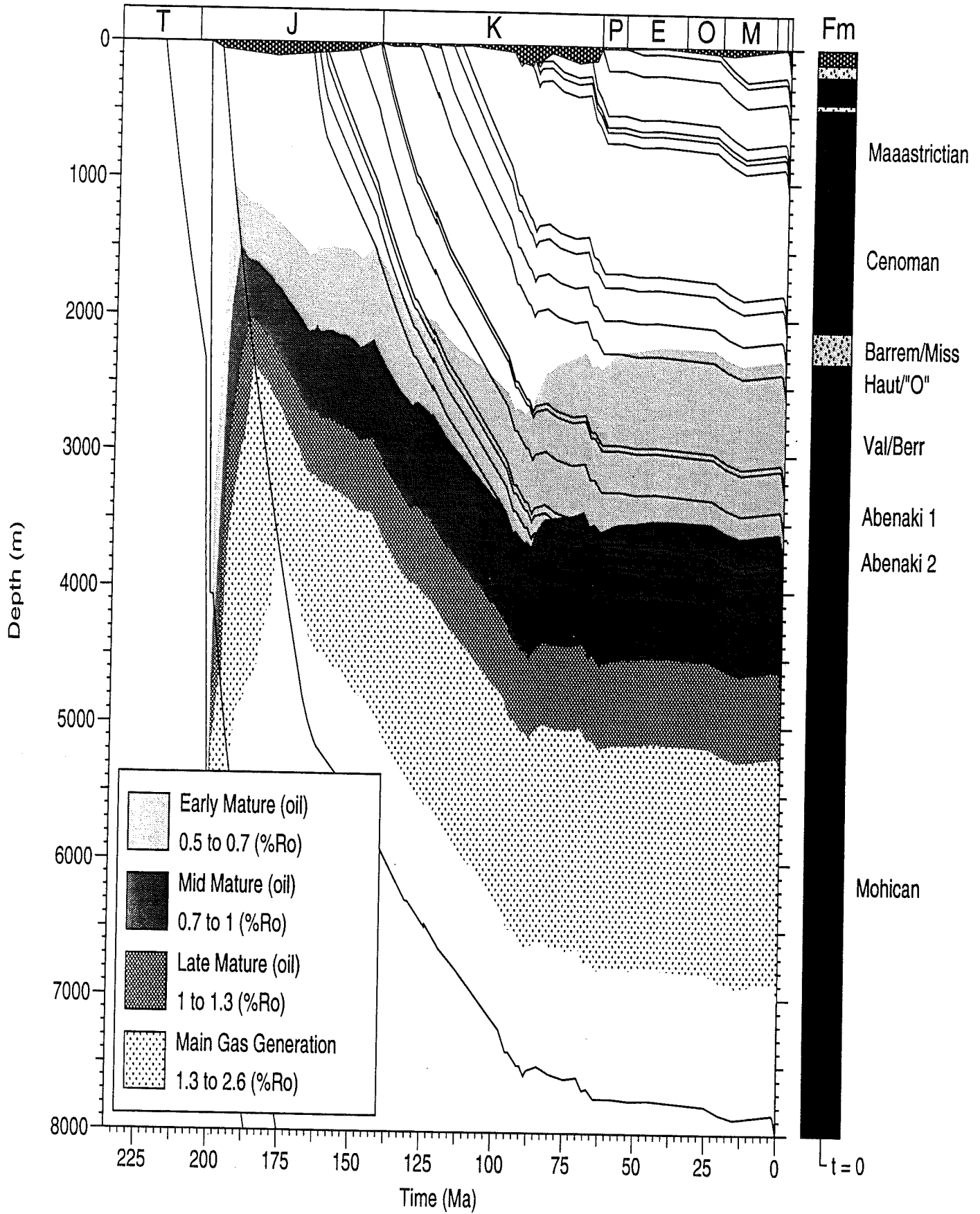


Figure 15a

Abenaki J-56

Scotian Shelf

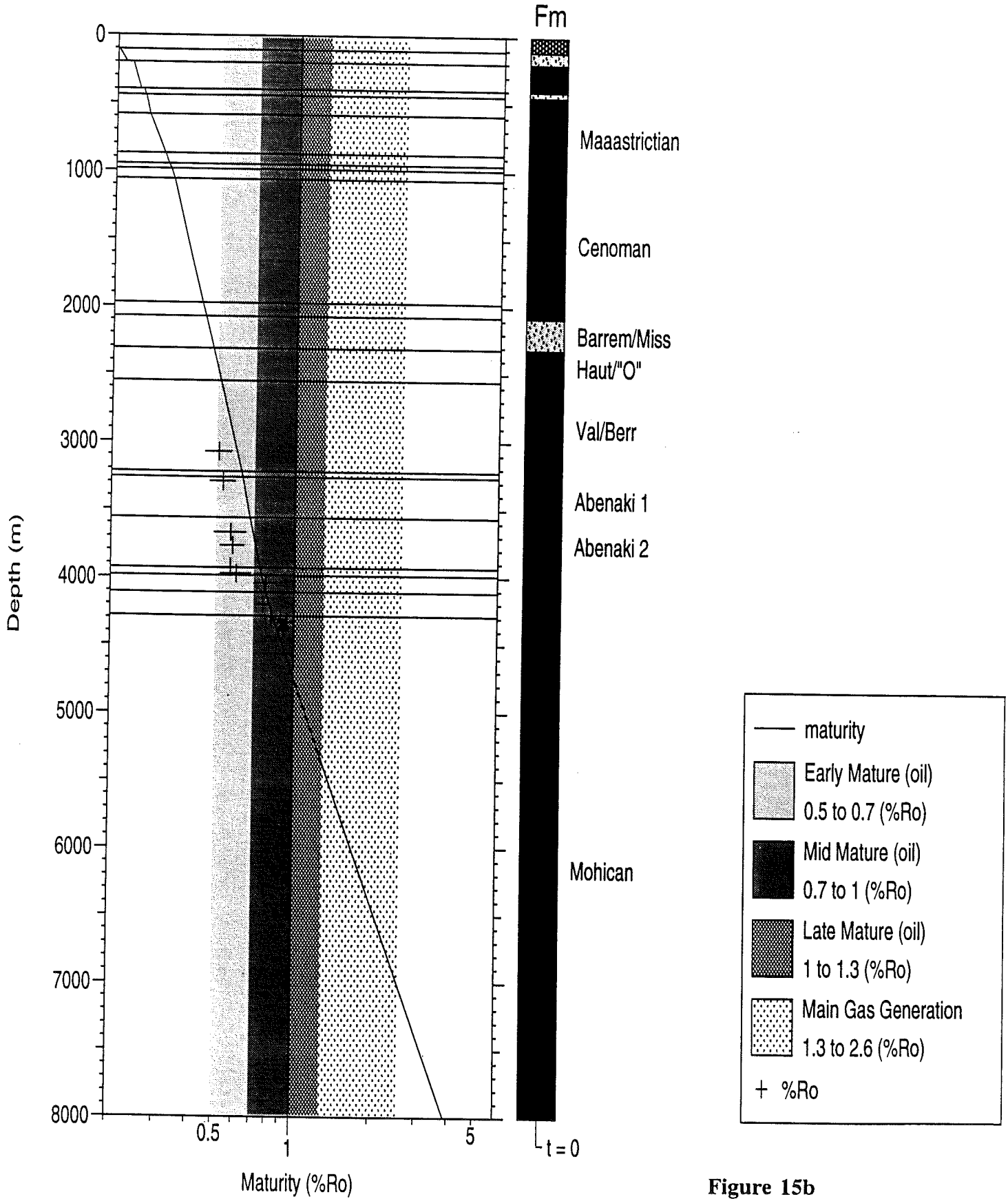


Figure 15b

Abenaki J-56

Scotian Shelf

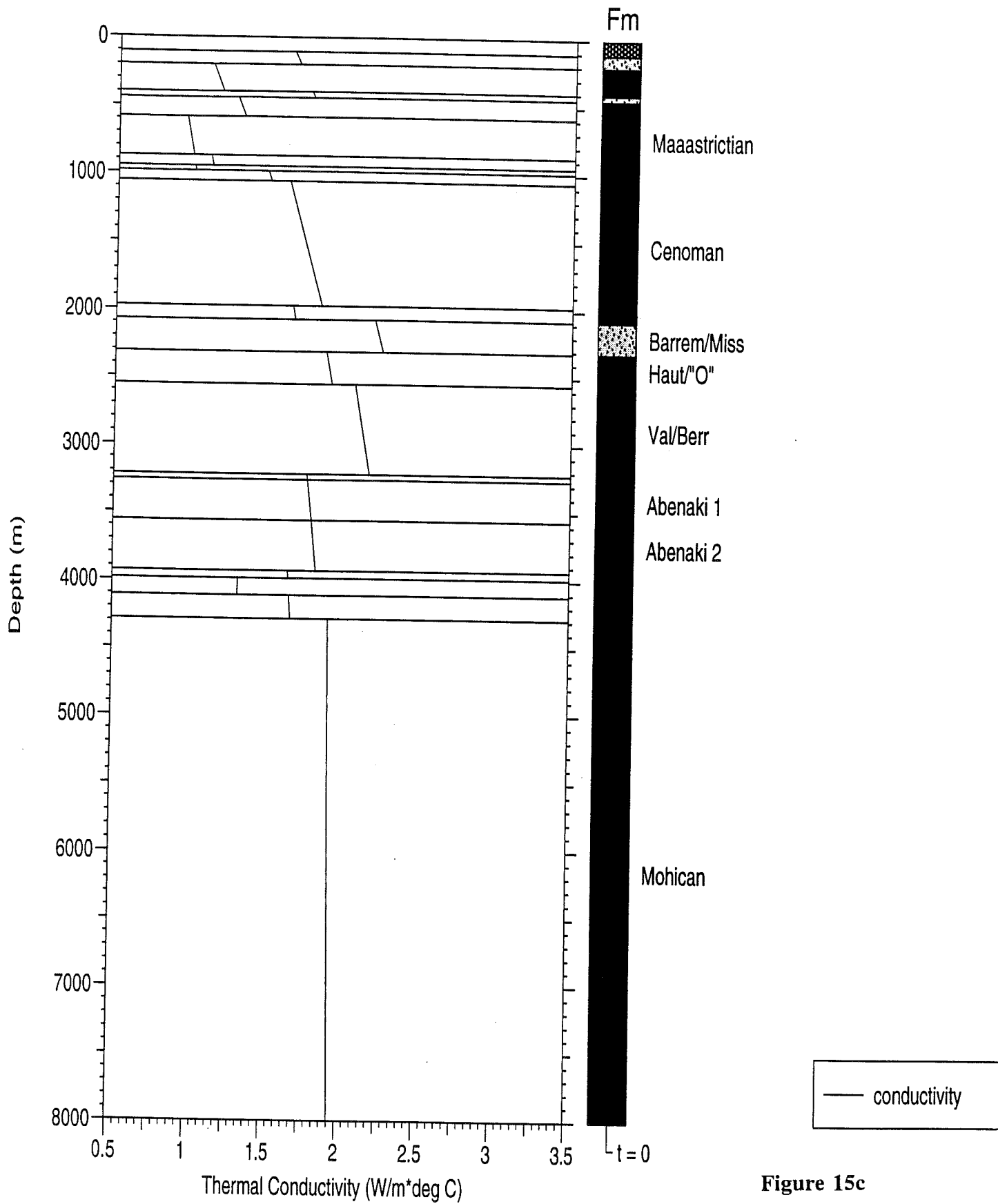


Figure 15c

Abenaki J-56

Scotian Shelf

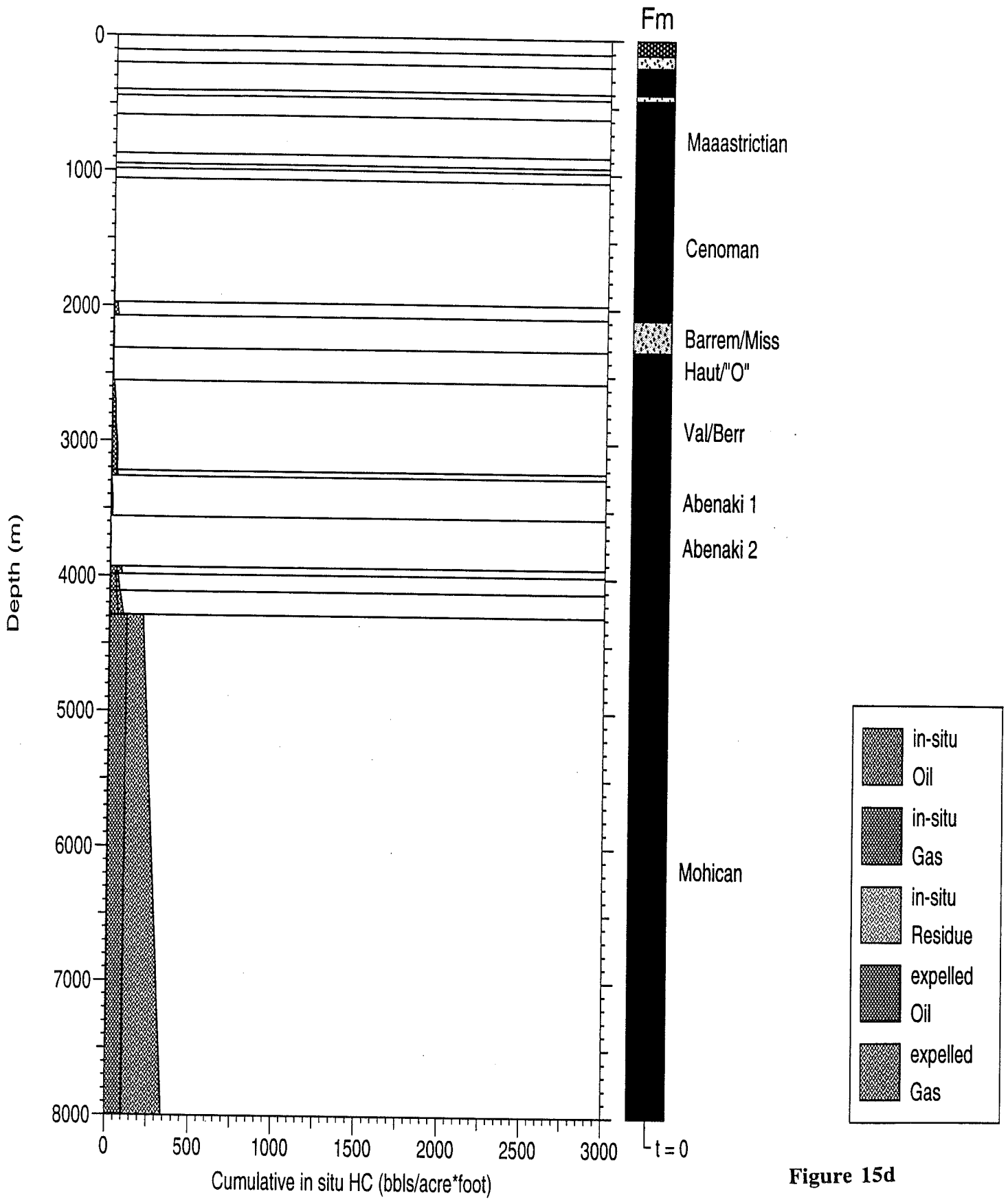


Figure 15d

Abenaki J-56

Scotian Shelf

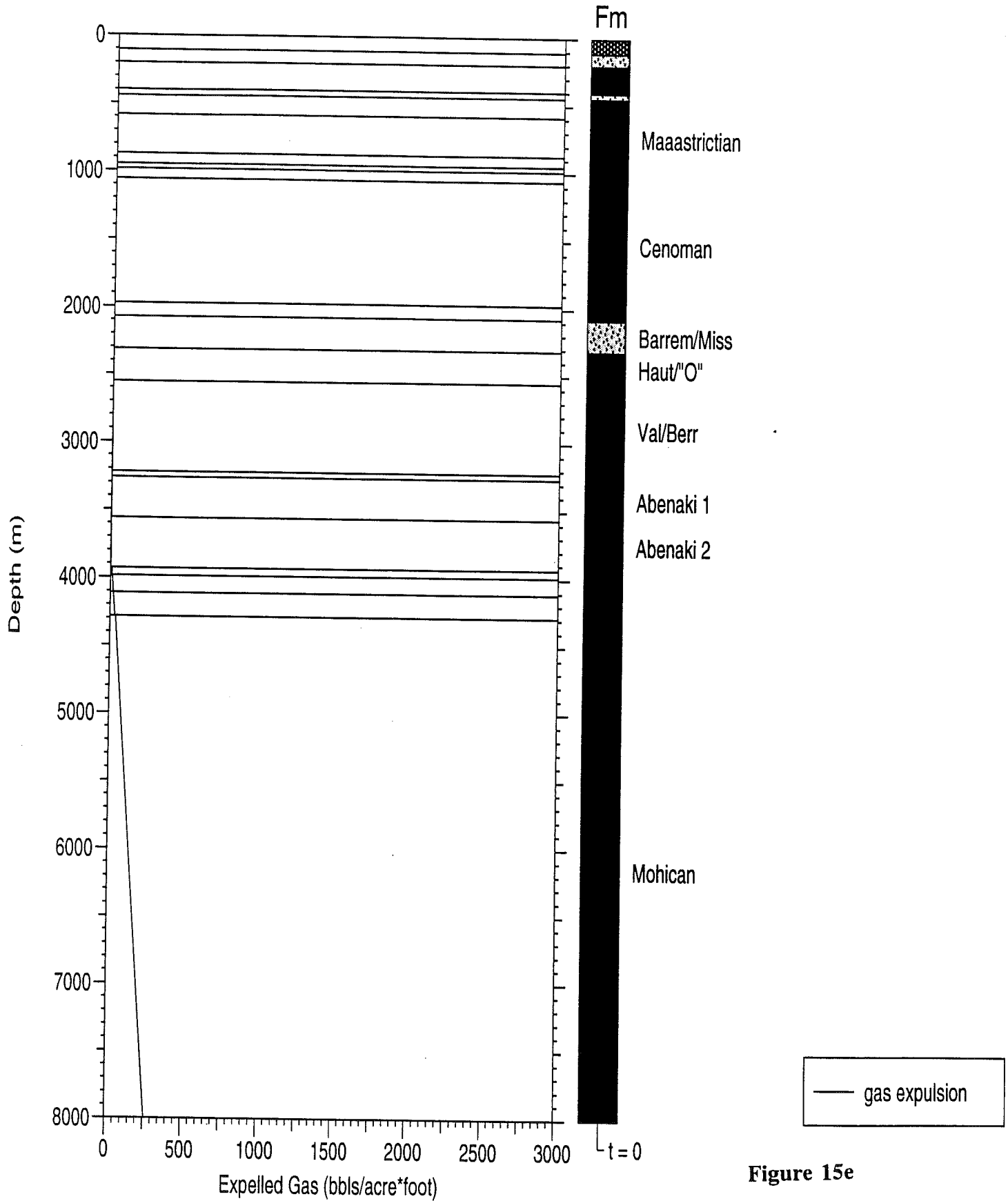


Figure 15e

Abenaki J-56

Scotian Shelf

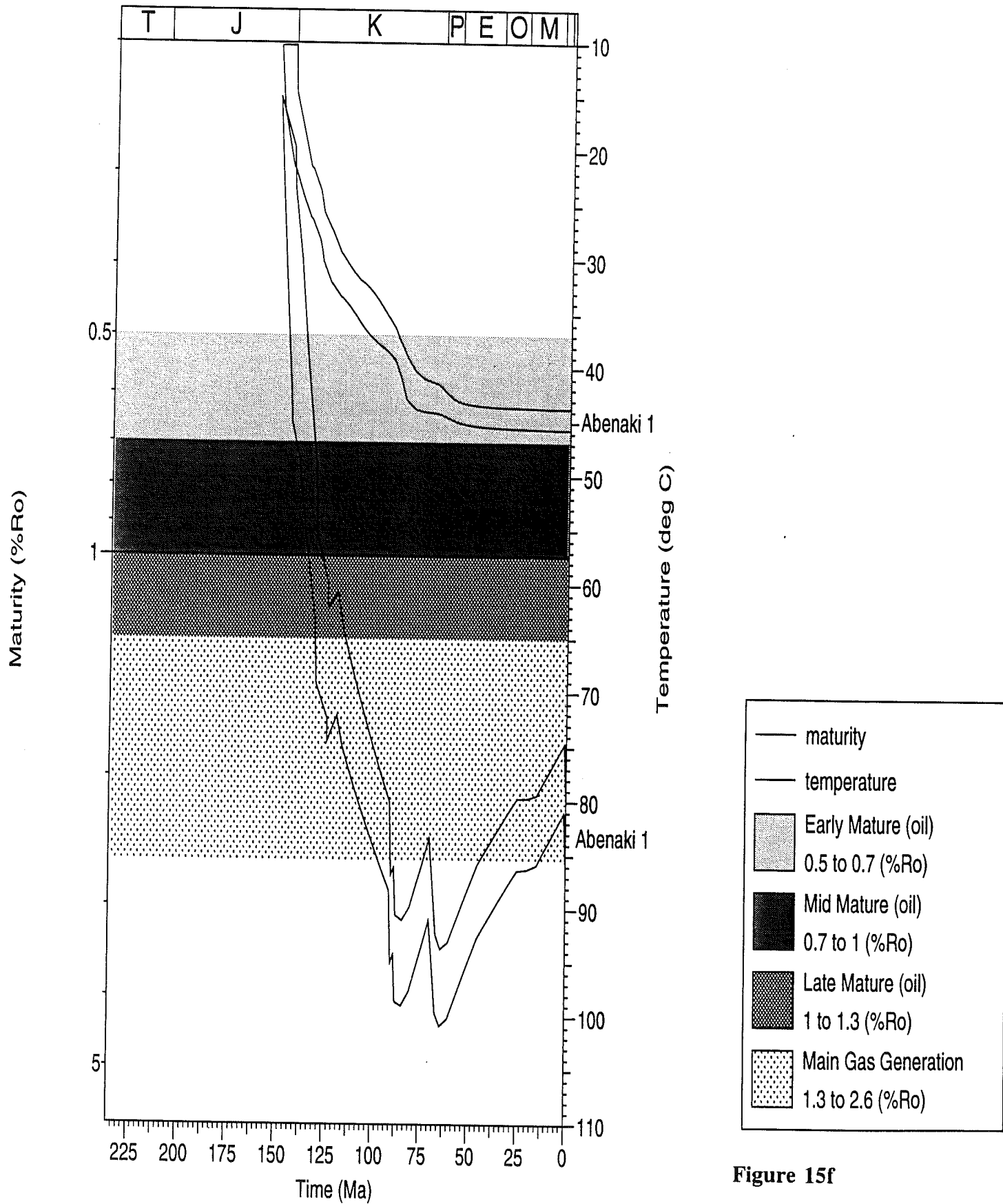


Figure 15f

Abenaki J-56

Scotian Shelf

Abenaki 1

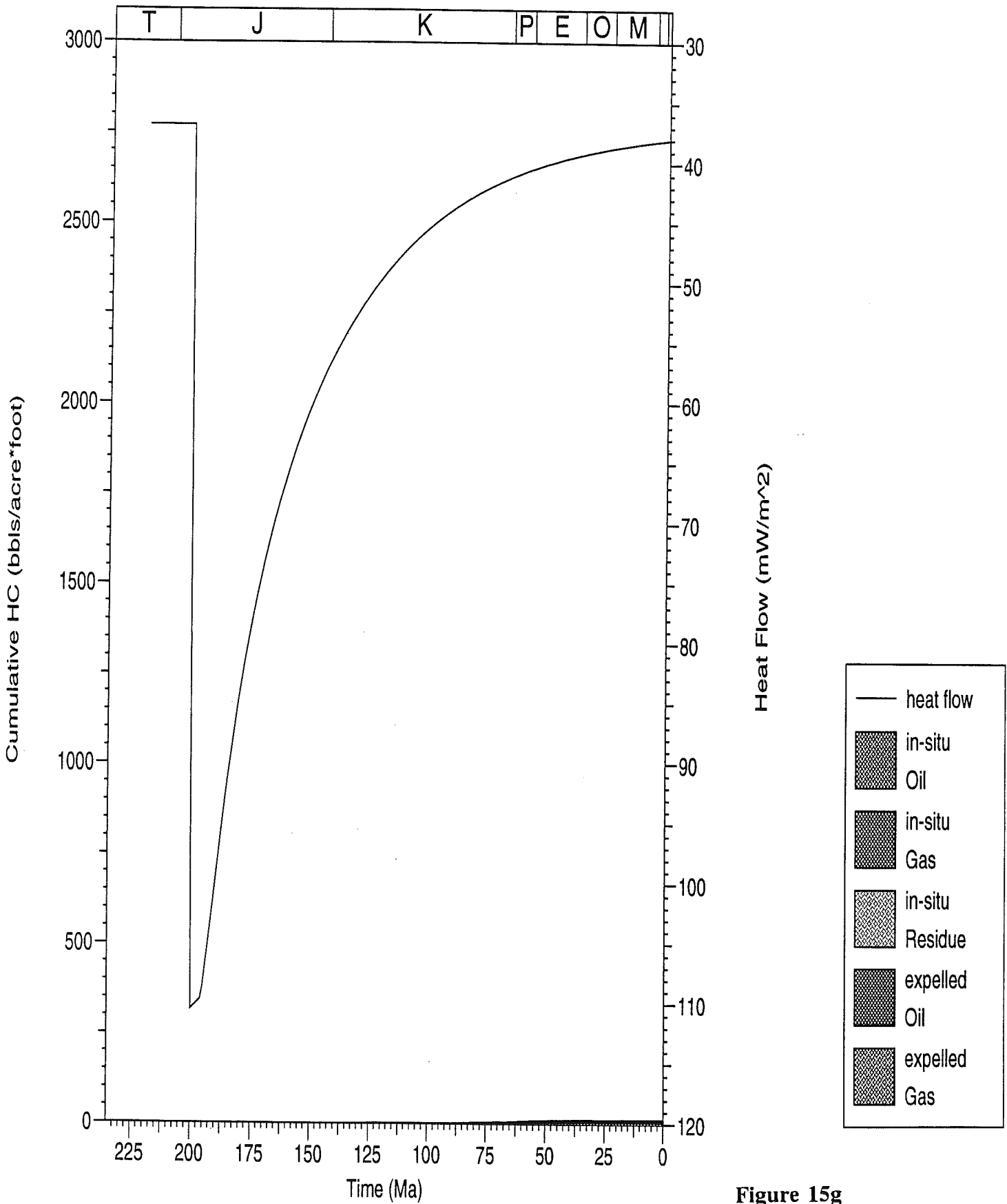


Figure 15g

Abenaki J-56

Scotian Shelf

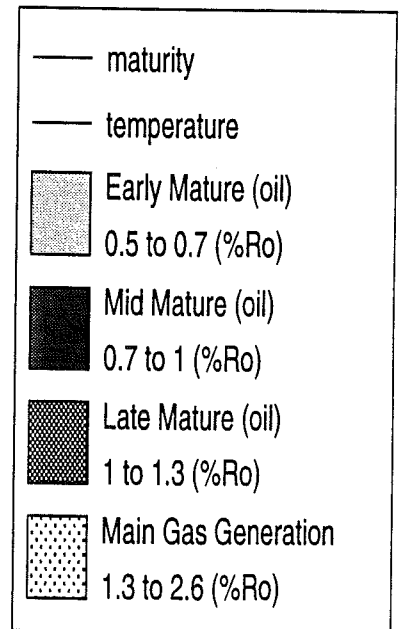
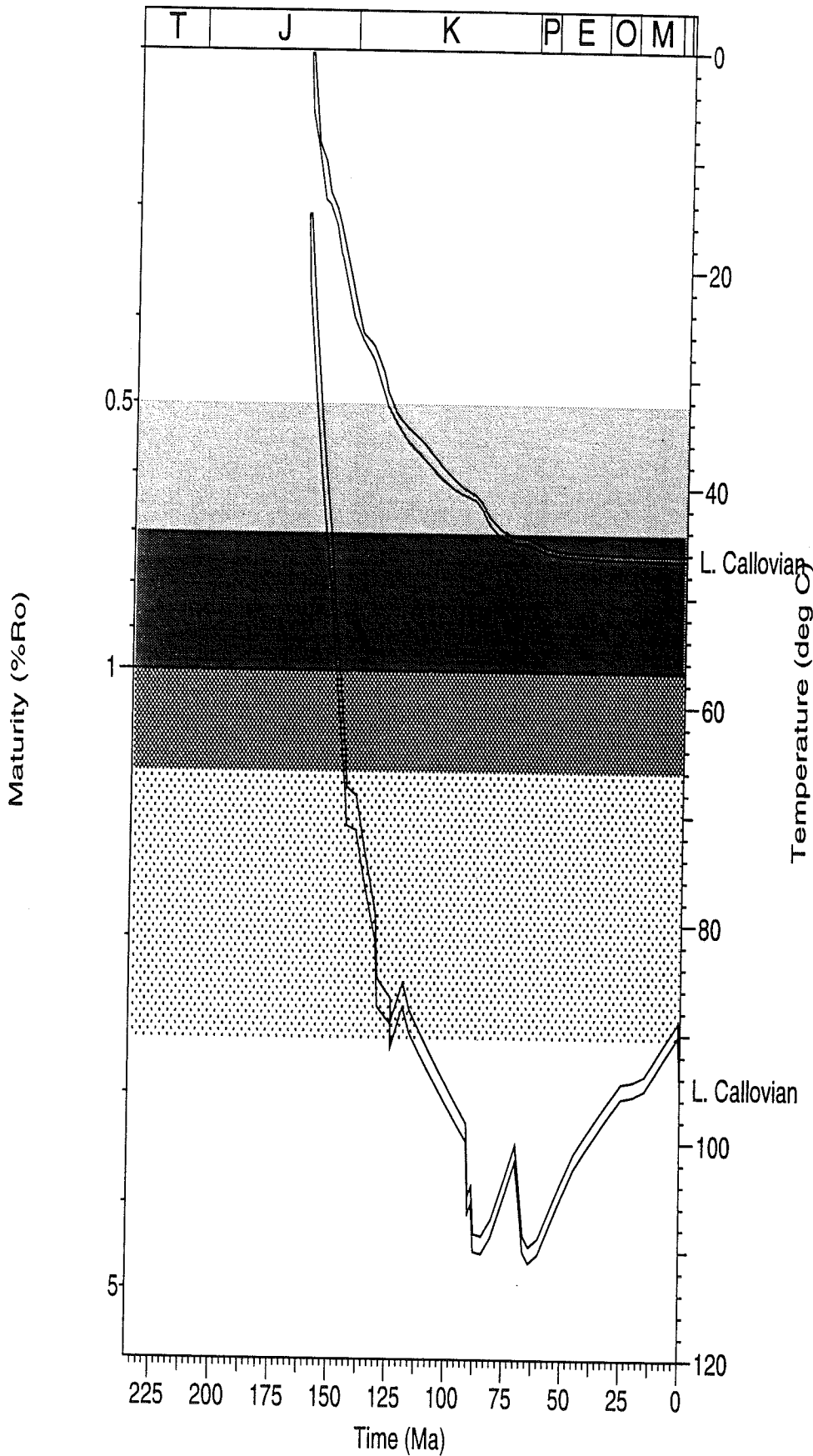


Figure 15h

Abenaki J-56

Scotian Shelf

L. Callovian

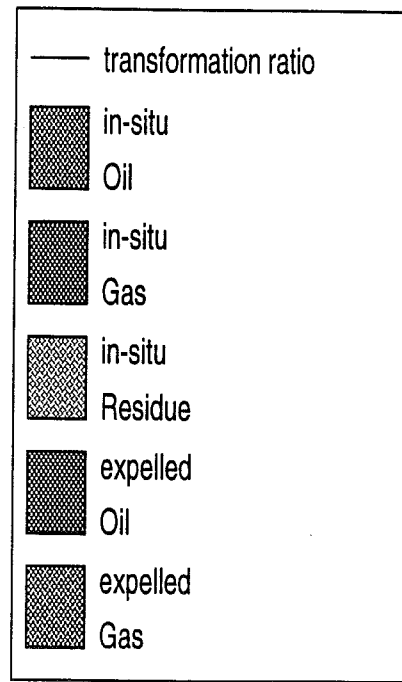
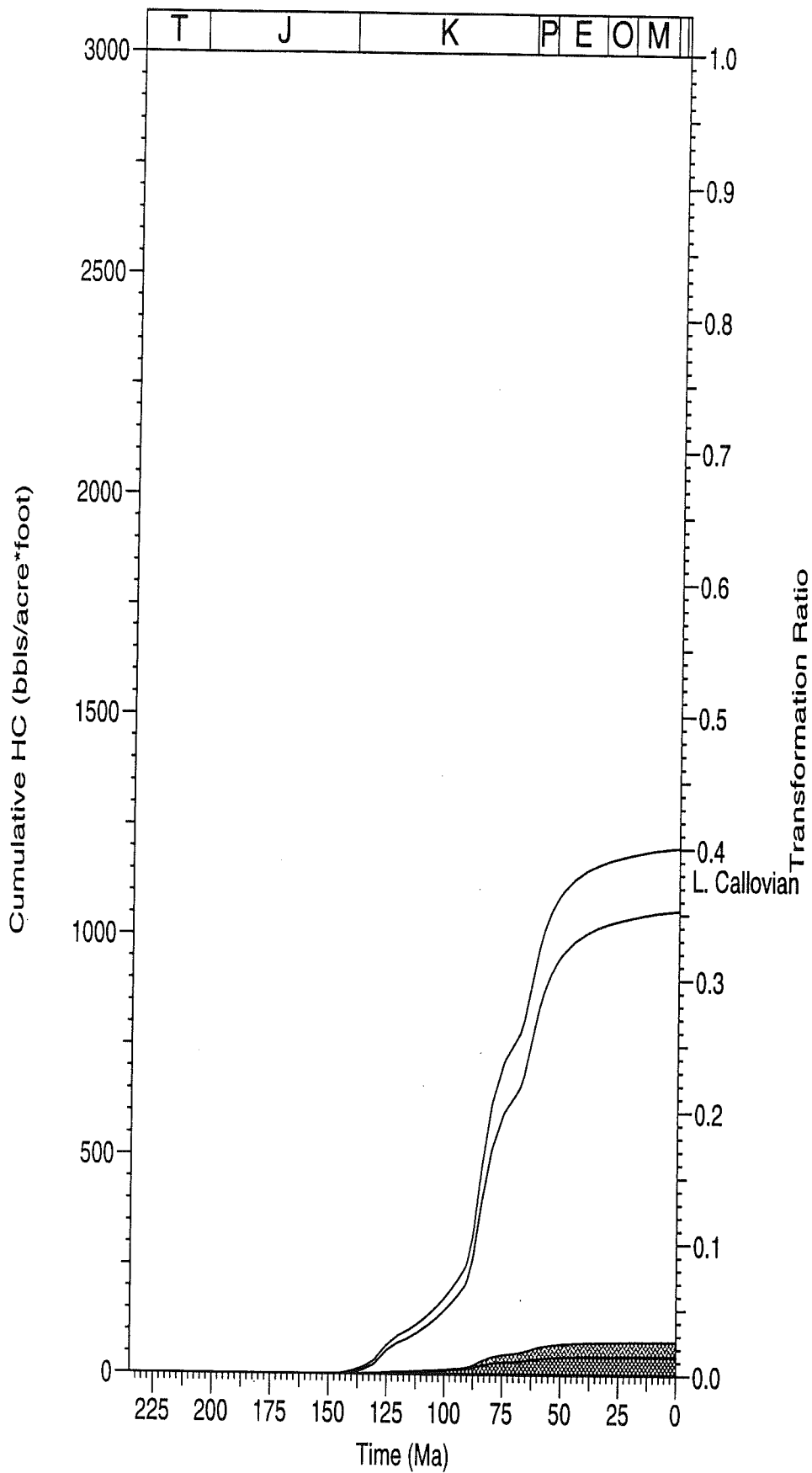


Figure 15i

Abenaki J-56

Scotian Shelf

Scatarie

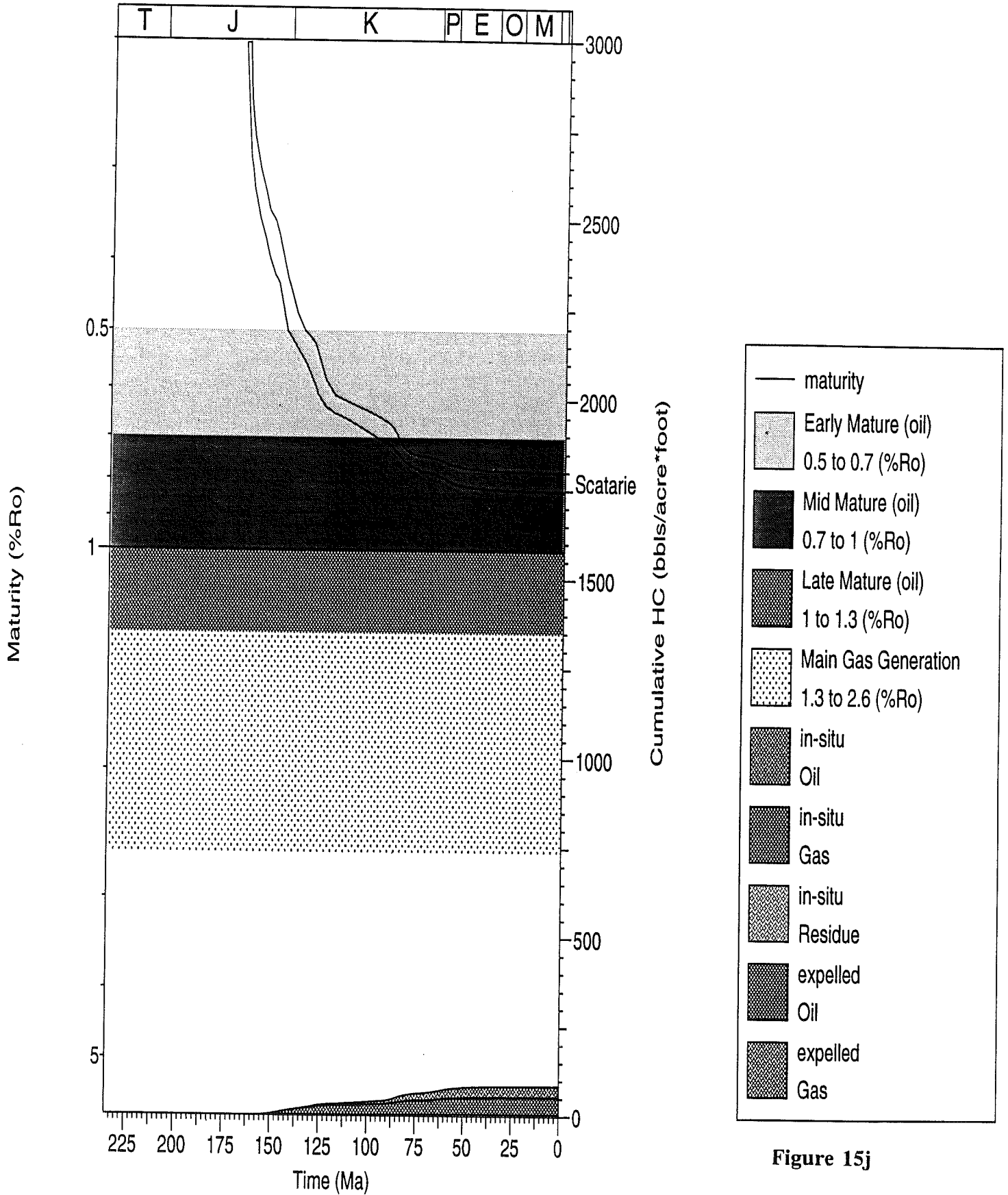


Figure 15j

Well: W. Chebucto K-20

Stratigraphy Table

Formation or Event Name	Type	Begin Age (Ma)	Well Top (m)	Present Thick (m)	Lithology	Kerogen Name
Quaternary	F	1.6	74	176	90/10/0	
Unconformity1	H	3				
Pliocene	F	5	250	550	30/70/0	
Unconformity2	H	10.5				
Miocene	F	20	800	275	Shale	
Unconformity3	H	30				
E. Oligocene	F	49.5	1075	425	10/80/10	
Unconformity4	H	58.5				
L. Paleocene	F	62	1500	180	0/90/10	
Unconformity5	H	70				
E. Maastricht	F	74.5	1680	29	0/60/40	Type III
Camp/Wyan	F	85	1709	94	0/90/10	IIA-IIB
Con/San/DC	F	92	1803	185	40/60/0	IIA-IIB
Log Can	F	98.5	1988	334	60/40/0	Type III
Sable	F	101	2322	168	50/50/0	IIB
Cree	F	117	2490	1241	60/40/0	IIB
Naskapi	F	120	3731	254	30/70/0	IIB
Miss 1	F	123.5	3985	415	60/40/0	Type III
Miss 2	F	127	4400	420	40/60/0	IIA-IIB
Miss 3	F	144	4820	1980	40/60/0	IIB
Top Jura	F	152	6800	500	20/80/0	Type III
MicMac/VC	F	181	7300	4300	20/70/10	Type III
Mohican	F	200	11600	3000	20/70/10	Type III
BU	H	201				
Argo	F	204	14600	2000	0/5/95	Type III
Eurydice	F	235	16600	1500	30/60/10	Type III

Table 9

W. Chebucto K-20

Scotian Shelf

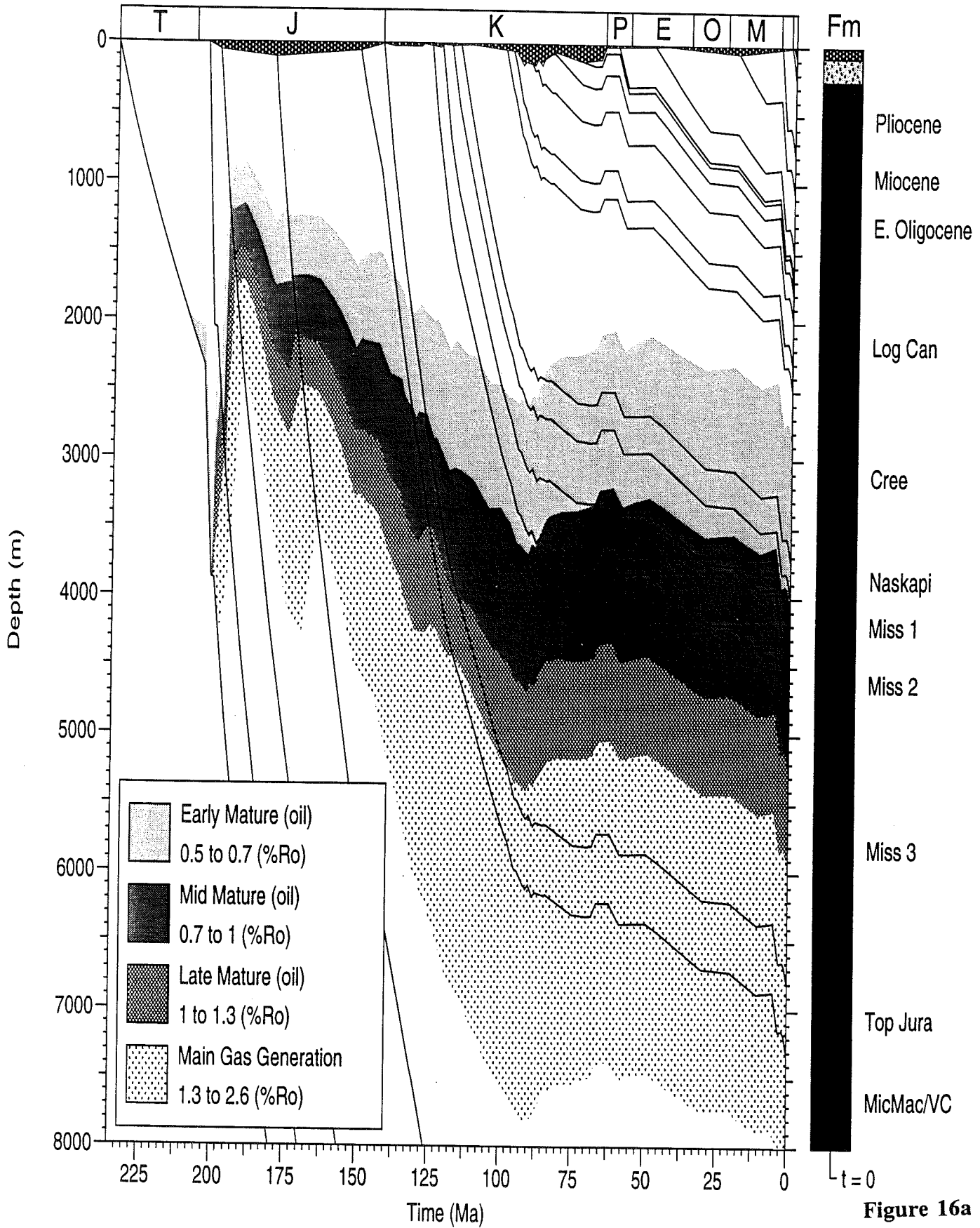


Figure 16a

W. Chebucto K-20

Scotian Shelf

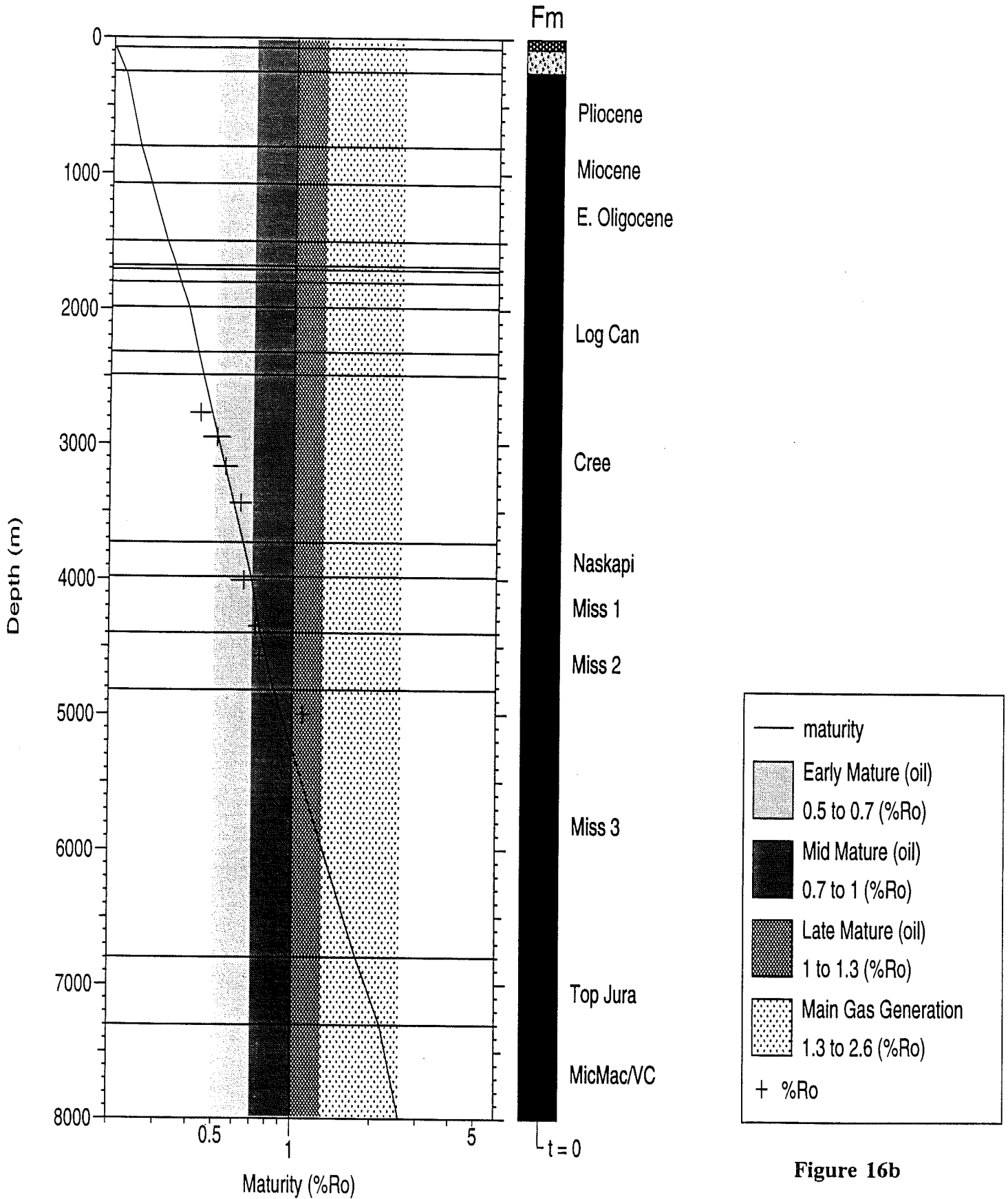
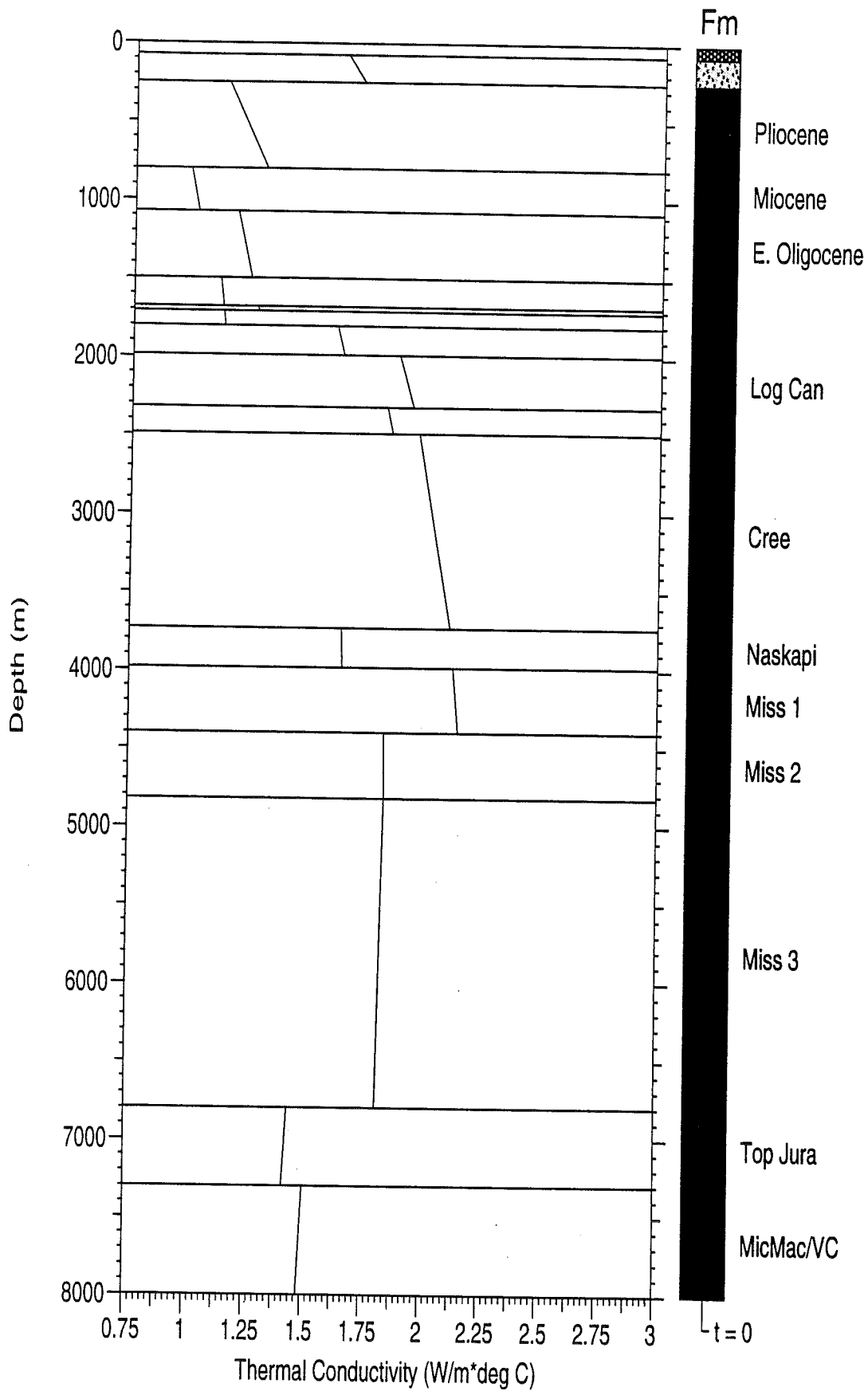


Figure 16b

W. Chebucto K-20

Scotian Shelf



— conductivity

Figure 16c

W. Chebucto K-20

Scotian Shelf

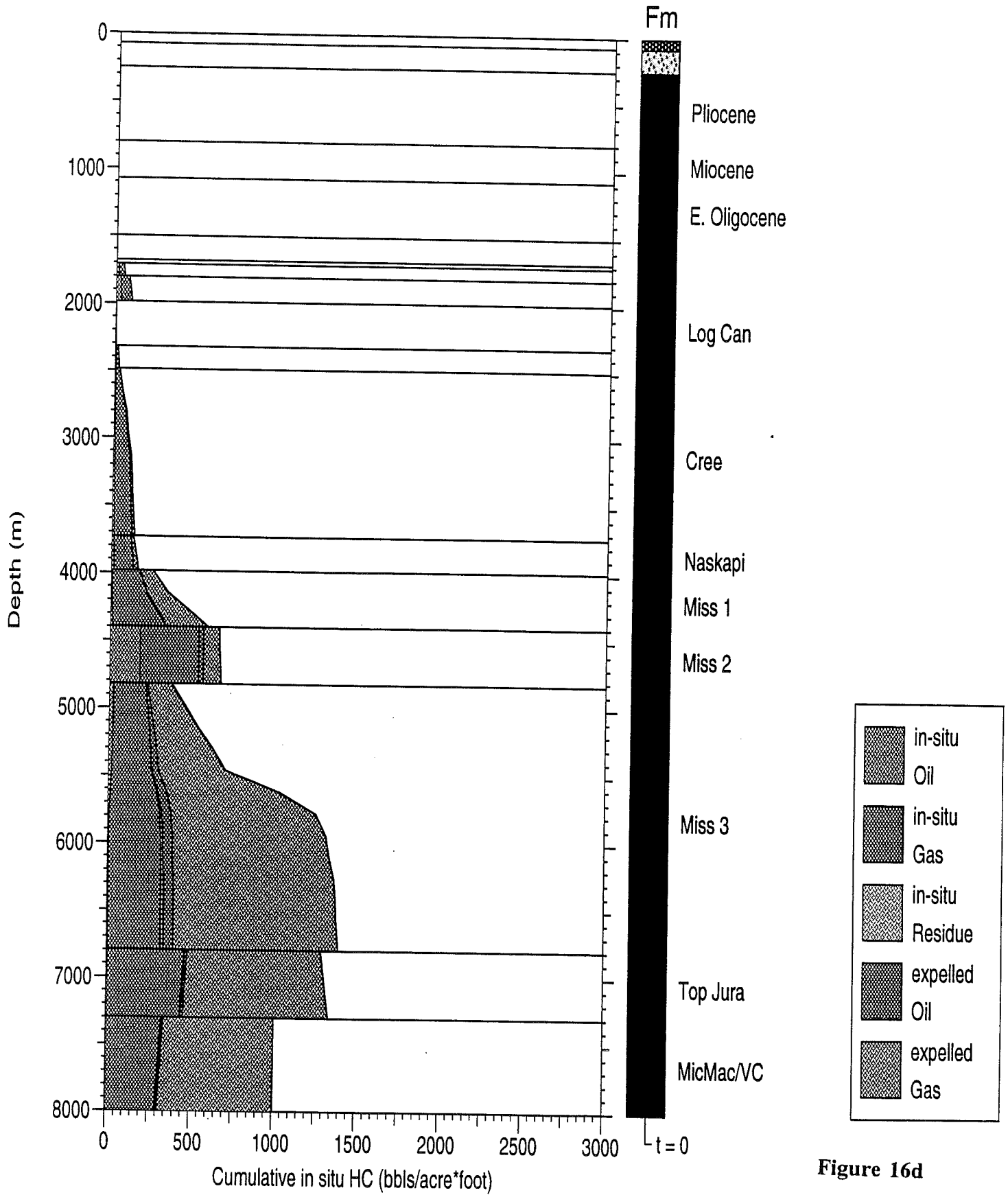


Figure 16d

W. Chebucto K-20

Scotian Shelf

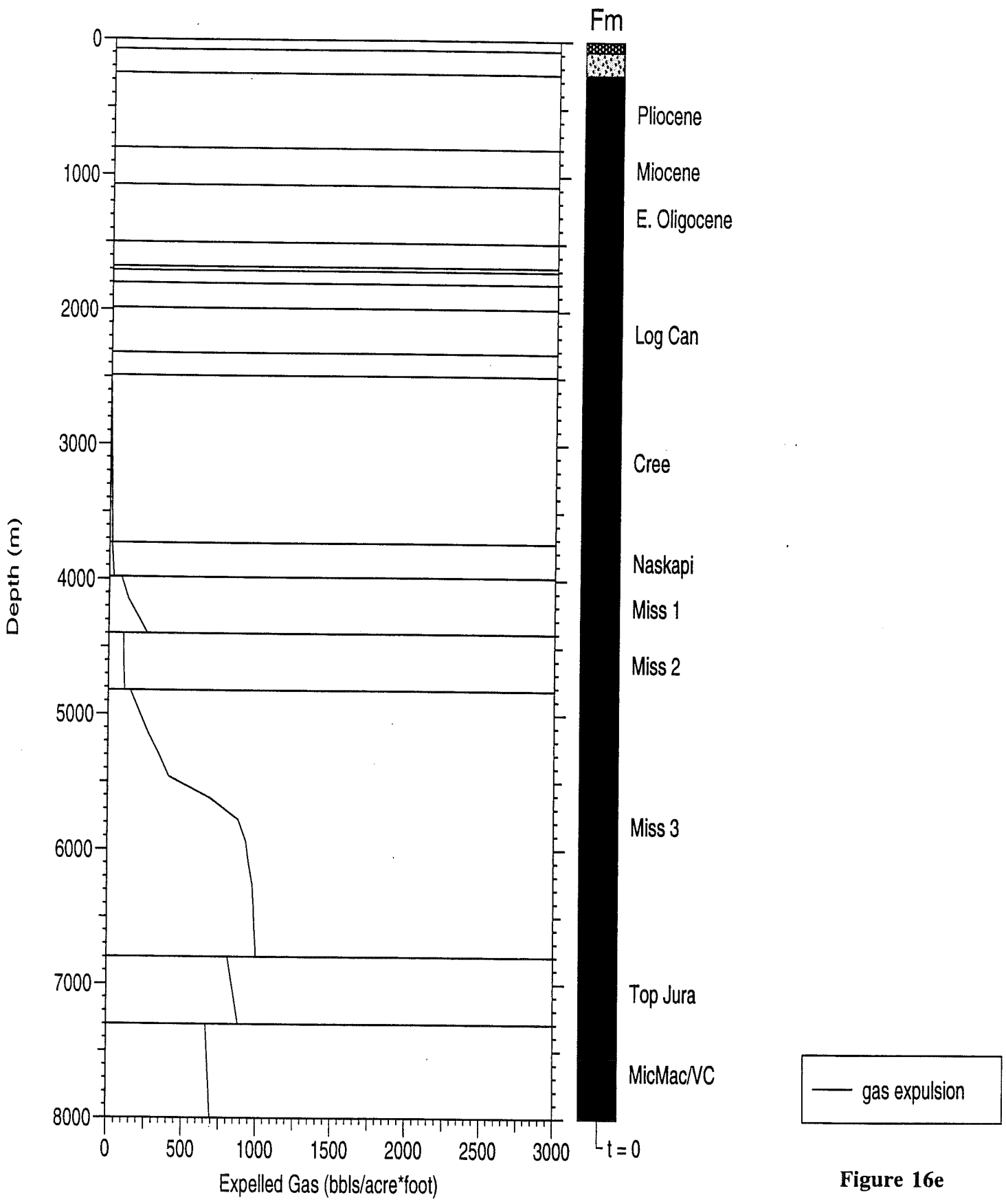


Figure 16e

W. Chebucto K-20

Scotian Shelf

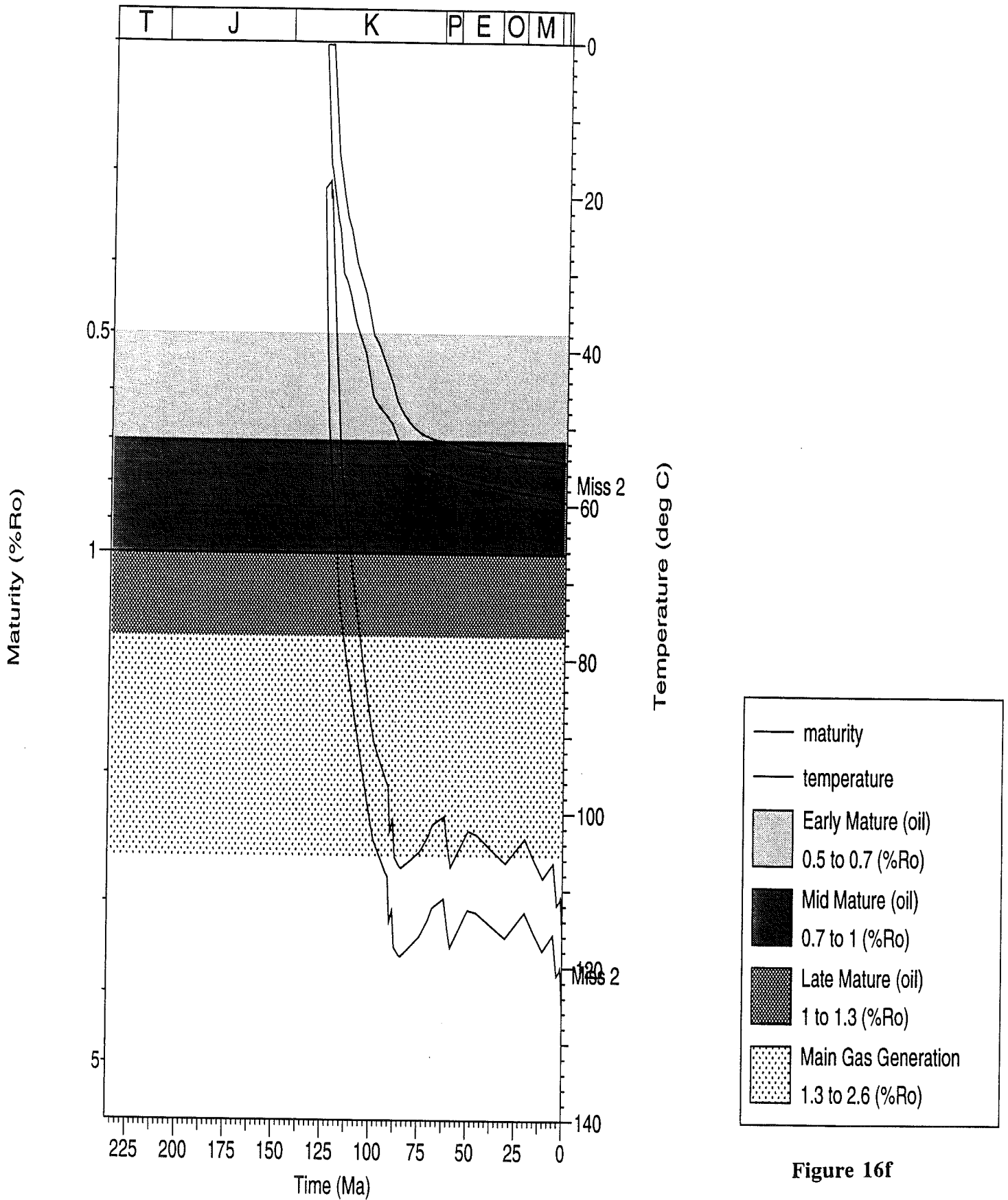
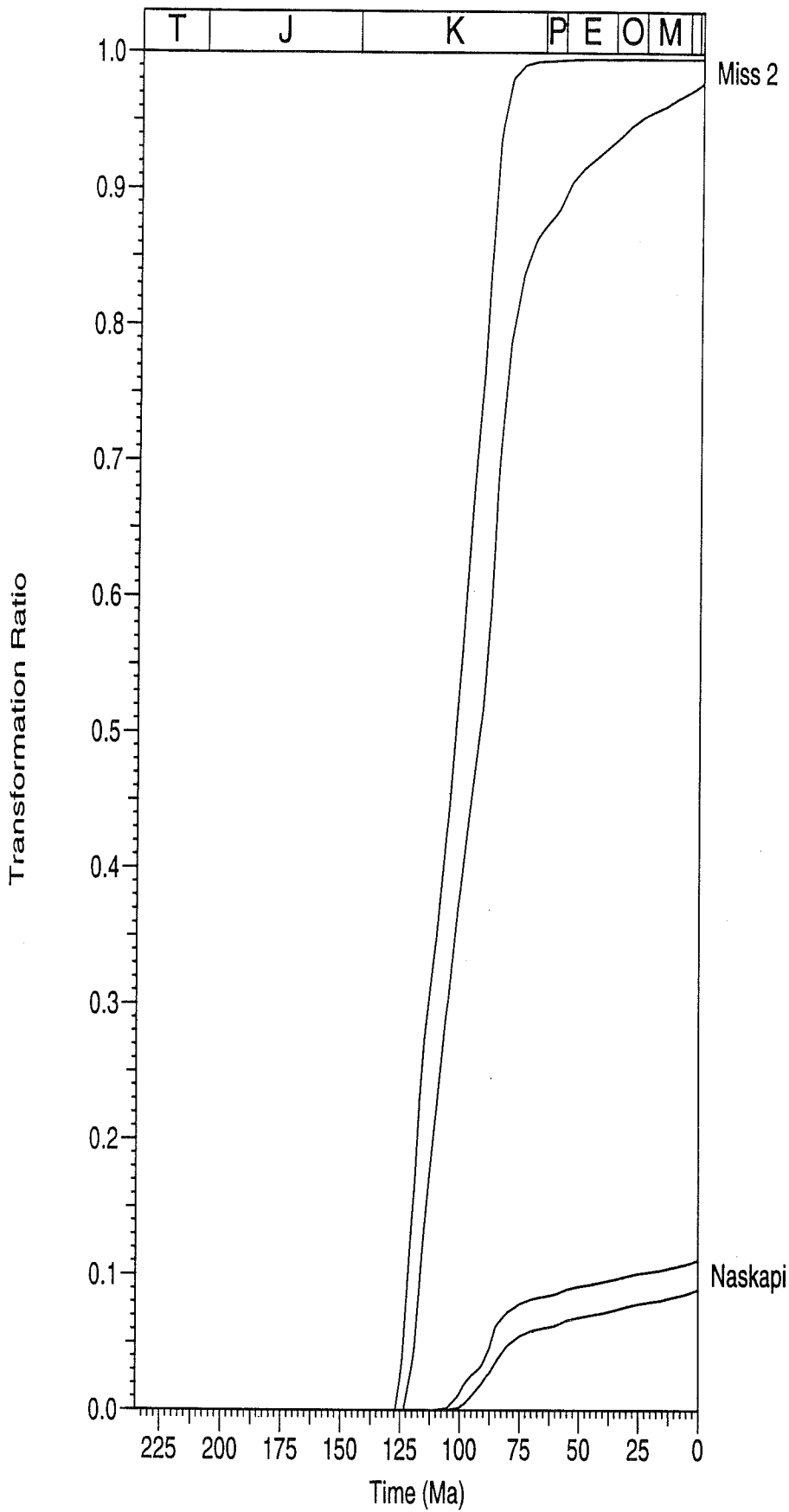


Figure 16f

W. Chebucto K-20

Scotian Shelf



— transformation ratio

Figure 16g

W. Chebucto K-20

Scotian Shelf

Miss 2

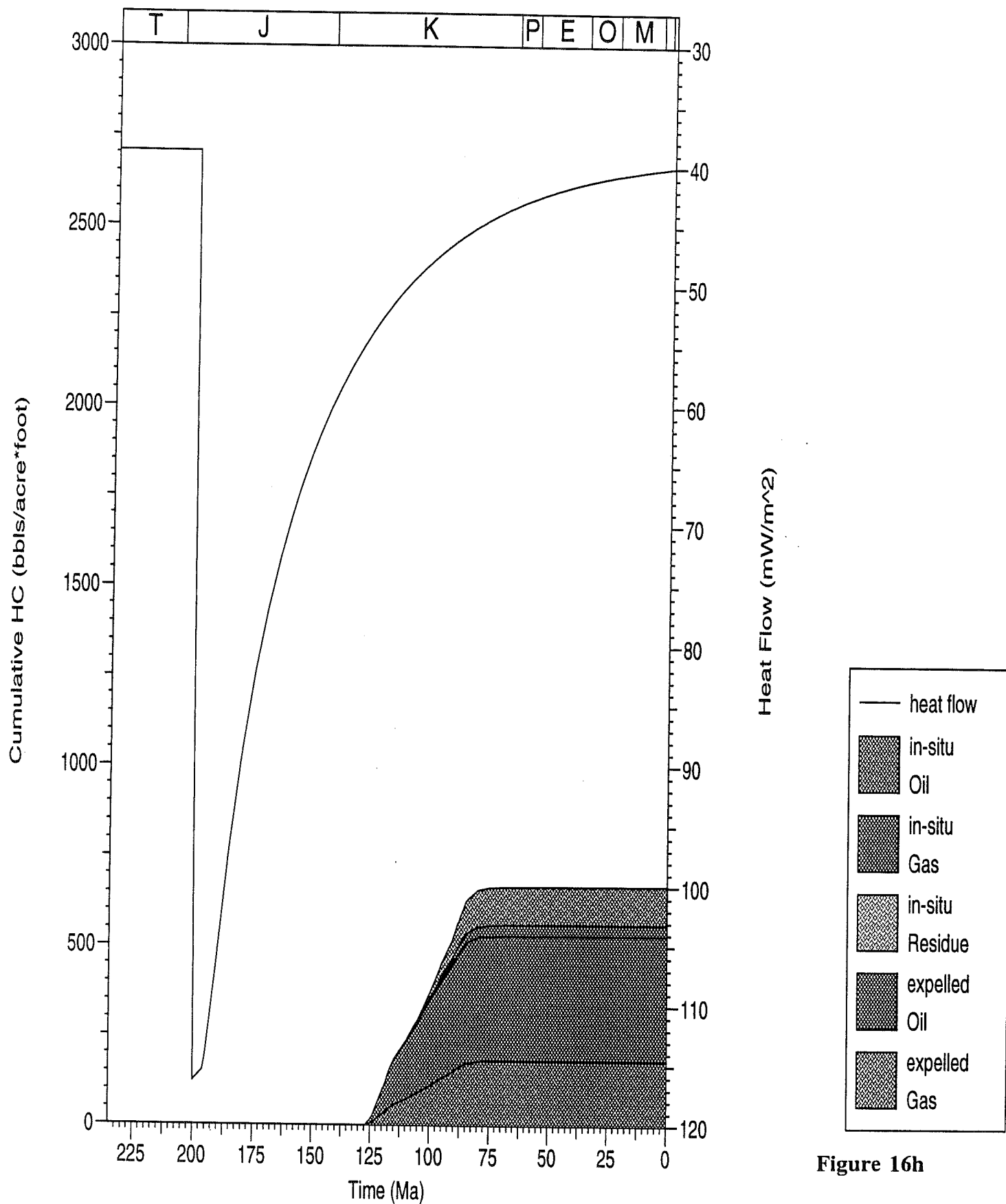
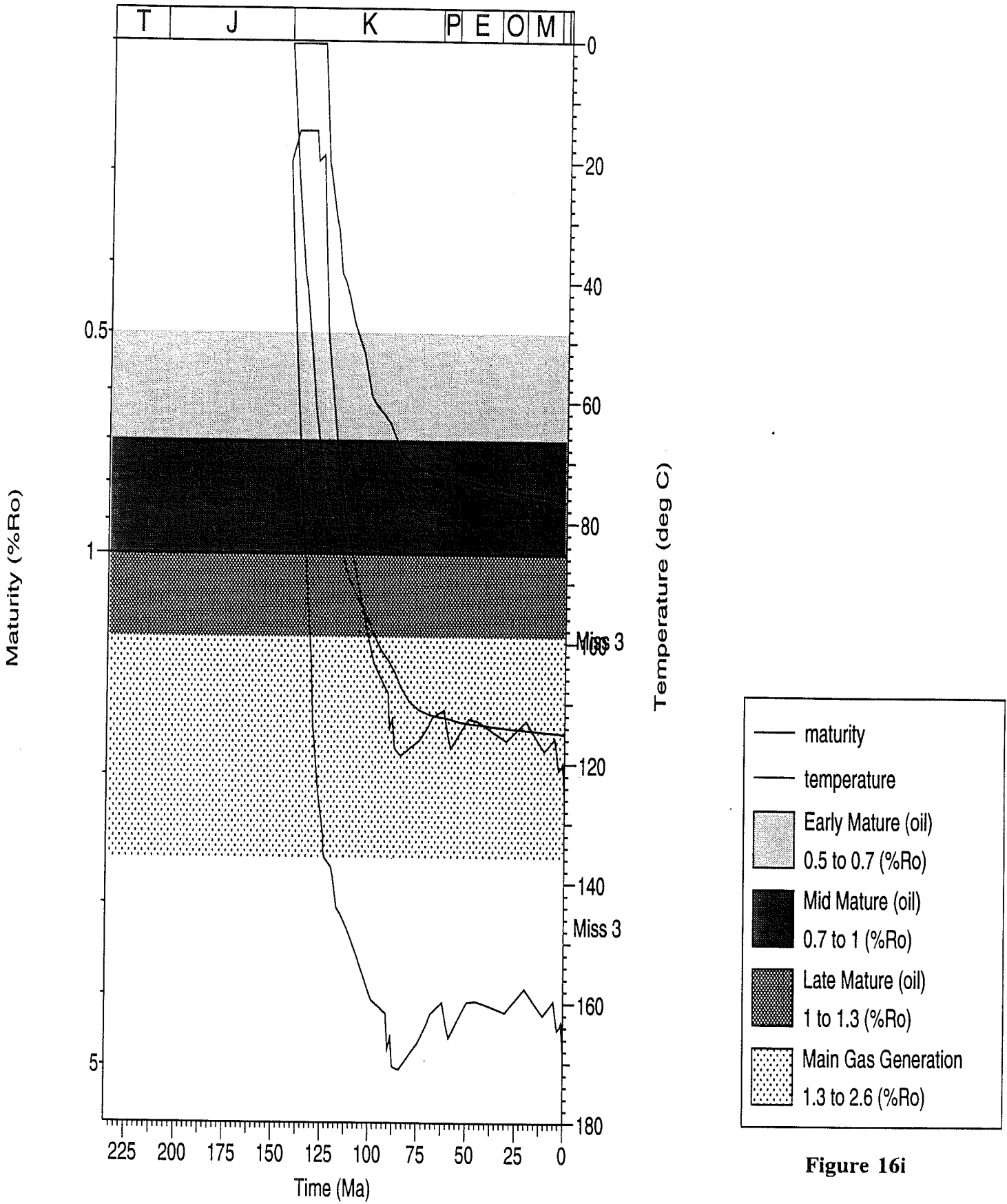


Figure 16h

W. Chebucto K-20

Scotian Shelf



W. Chebucto K-20

Scotian Shelf

Miss 3

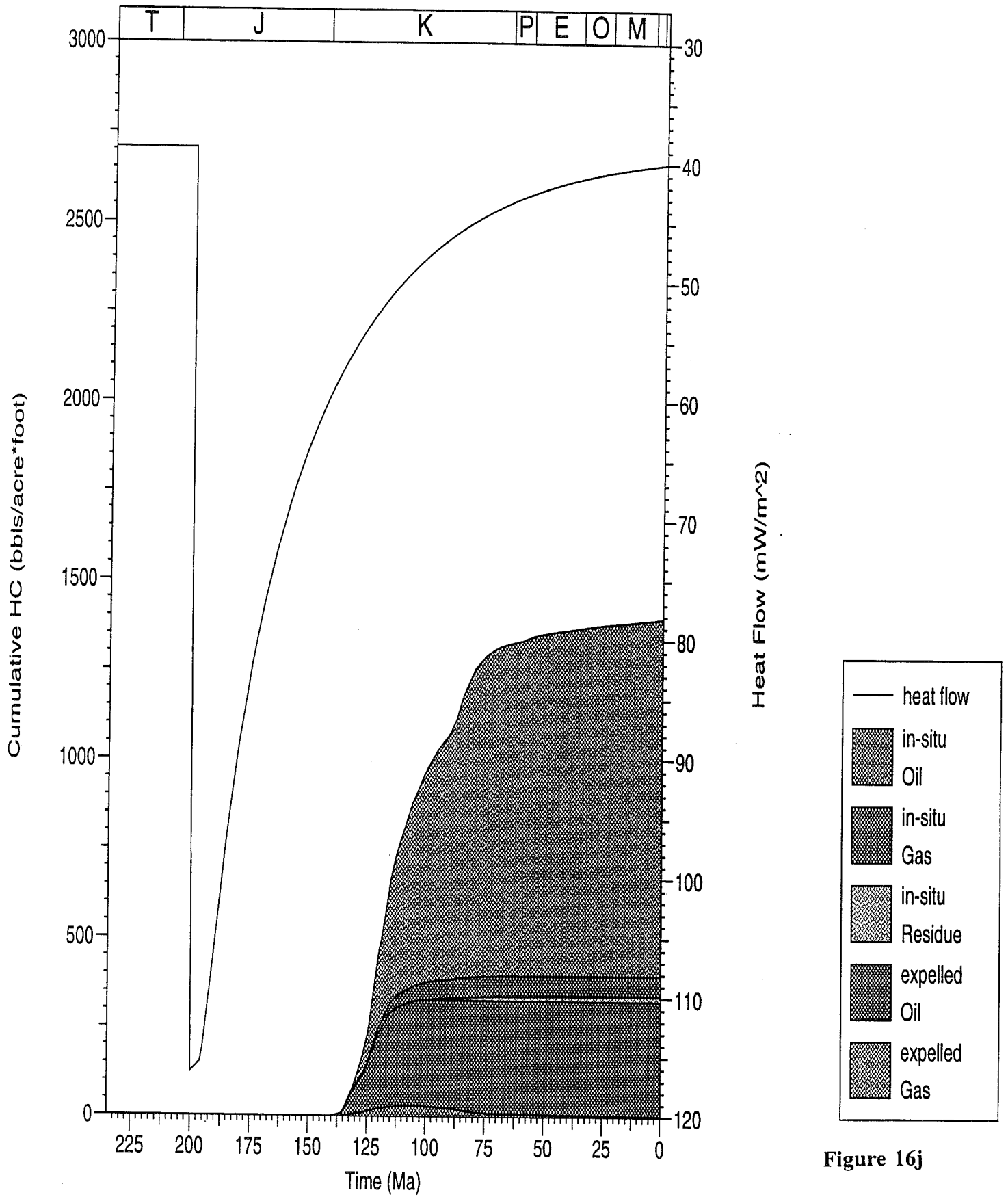


Figure 16j

W. Chebucto K-20

Scotian Shelf

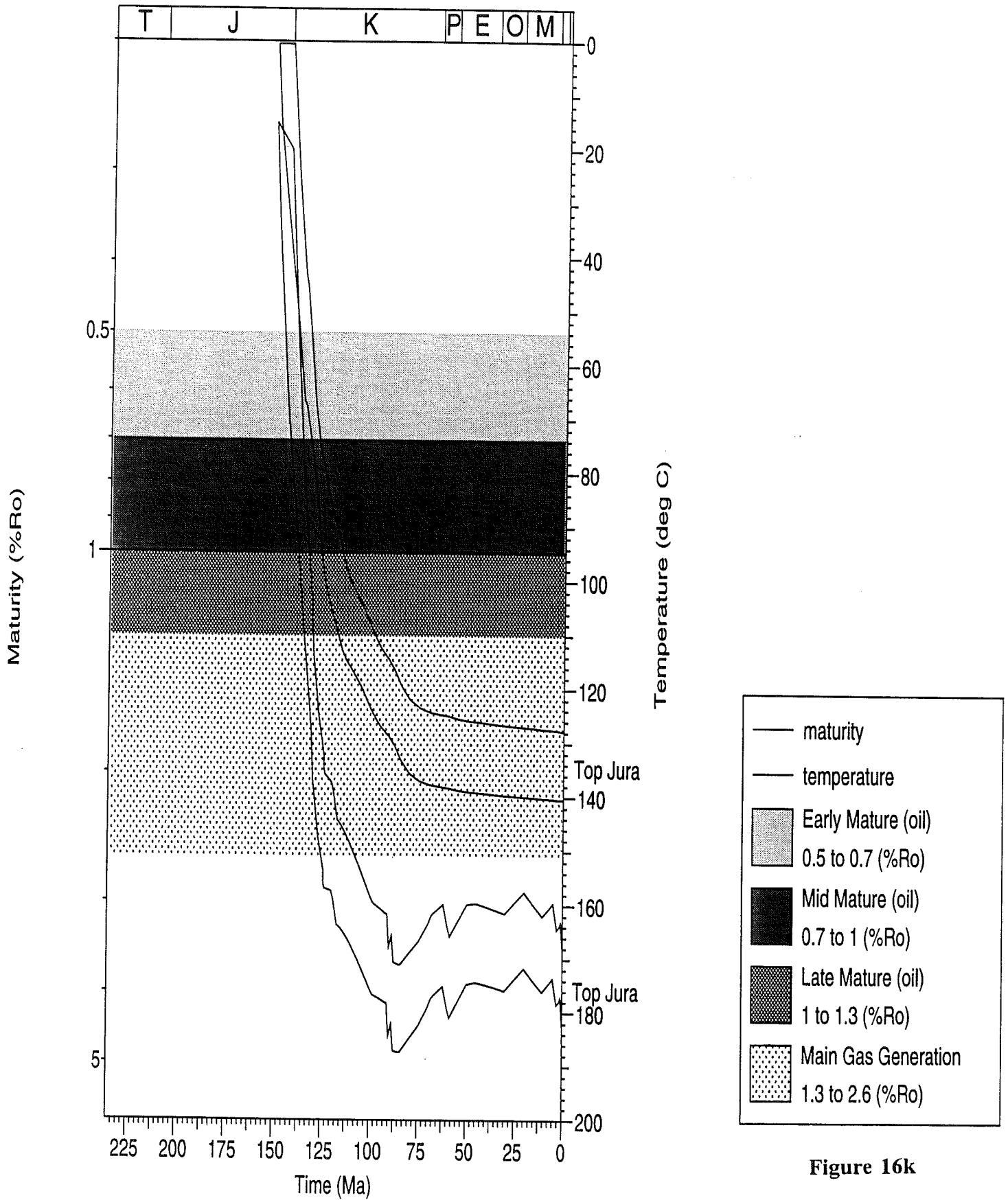


Figure 16k

W. Chebucto K-20

Scotian Shelf

Top Jura

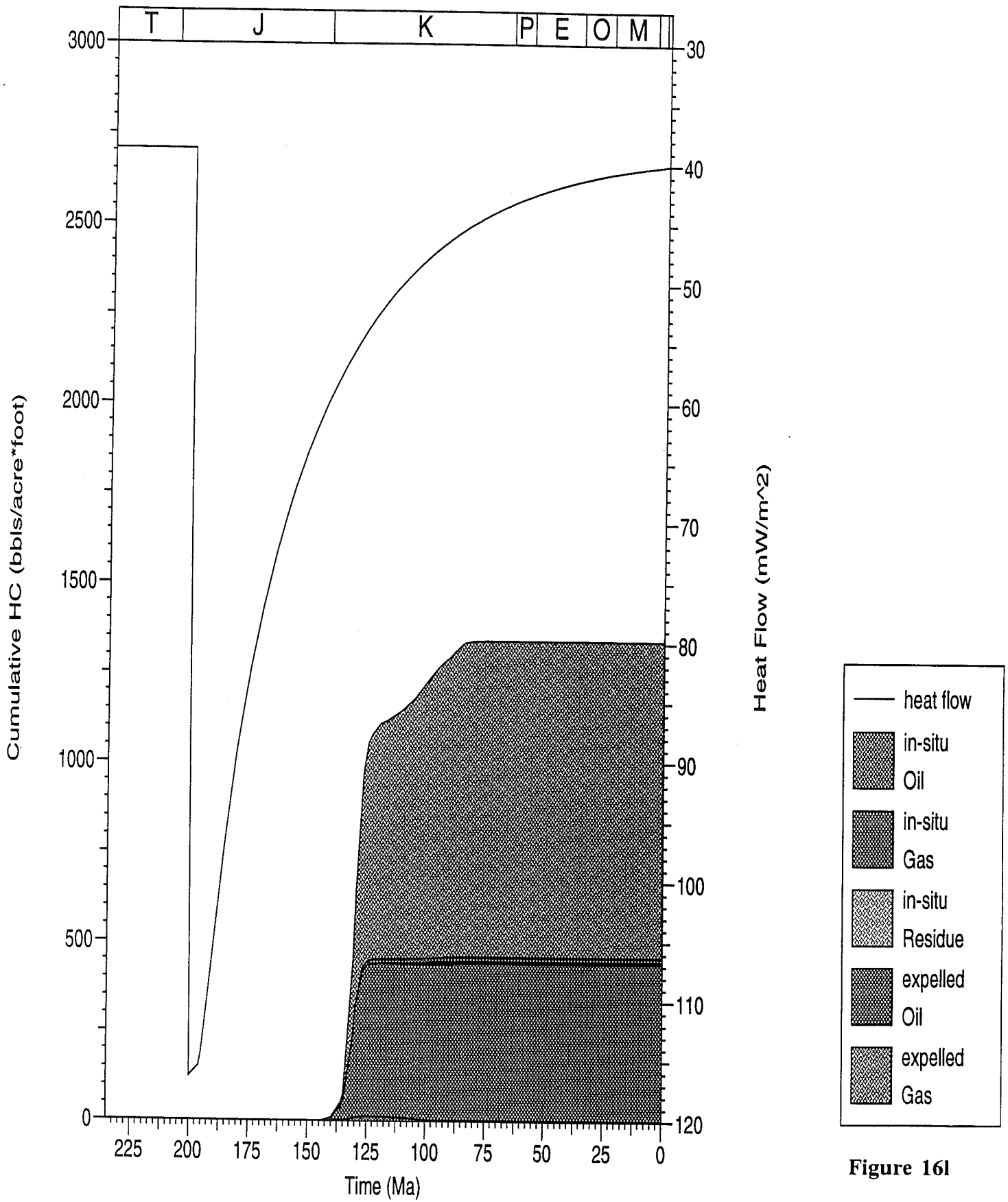


Figure 161