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FLOOD OF JULY 1982

ON THE WAPITI RIVER

BRITISH COLUMBIA

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FLOOD OF JULY 1982
ON THE WAPITI RIVER
BRITISH COLUMBIA

A.G. SMITH

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Planning and Studies Section
Water Resources Branch
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TABLE OF CONTENTS

	Page
ABSTRACT	1
1. INTRODUCTION	1
Physiography	2
Climate	3
Precipitation	4
Snowfall	4
Temperature	6
Geology	6
Soil Types	6
Vegetation	7
Wapiti River Basin in British Columbia	7
2. DESCRIPTION OF FLOOD	9
Rainstorm of July 13-15, 1982	11
Meteorological description of the rainstorm of July 13-15, 1982, over Smoky River Basin (Wapiti River)	11
3. ANALYSIS OF FLOOD DATA	14
Wapiti River above Mistanusk Creek (07GCC001)	14
Estimation of flood discharge	14
4. FREQUENCY CHARACTERISTICS OF THE FLOOD	16
Interpretation of frequency curves	17
5. CONCLUSIONS AND RECOMMENDATIONS	18
Causes of the flood	18
Peak flow	18
BIBLIOGRAPHY	20

TABLE OF CONTENTS

Page

TABLES

1. Snow course data for selected stations	5
2. Mean annual and maximum runoff for selected basins	8
3. Physiography of selected streamflow stations draining eastern slopes of Rocky Mountains	10
4. Daily precipitation for selected stations - July 1-31, 1982	12

ILLUSTRATIONS

Figure 1. Map of British Columbia with Wapiti River Basin	21
Figure 2. Physiographic divisions of the Canadian Cordillera	22
Figure 3. Profiles of (a) Wapiti River; (b) Red Deer Creek, and (c) Belcourt Creek	23
Figure 4. Snow course stations	24
Figure 5. Wapiti River Basin in British Columbia	25
Figure 6. Streamflow hydrograph - 1979 snowmelt peak Pouce Coupe River below Henderson Creek	26
Figure 7. Streamflow hydrograph - 1982 rainfall peak Sukunka River at the Mouth	27
Figure 8. Streamflow hydrograph - 1980 rain on snow peak Wapiti River above Mistanusk Creek	28
Figure 9. Streamflow hydrograph - 1980 winter rainfall peak Beaverlodge River near Beaverlodge	29
Figure 10. Mass curves of rainfall July 1-31, 1982	30
Figure 11. Isohyetal chart - Smoky River Basin - July 13-15, 1982	31
Figure 12. Surface low and 500-Millibar low chart - July 14-15, 1982	32
Figure 13. Wapiti River above Mistanusk Creek after flood of July 14, 1982	33
Figure 13a. Wapiti River above Mistanusk Creek after flood of July 14, 1982.	34
Figure 14. Frequency curve, Wapiti River above Mistanusk Creek	35
Figure 15. Frequency curve, Wapiti River near Grande Prairie	37
Figure 16. Frequency curve, Kakwa River near Grande Prairie	39
Figure 17. Frequency curve, Smoky River above Hells Creek	41
Figure 18. Frequency curve, Cutbank River near Grande Prairie	43
Figure 19. Envelope curve of extreme floods - eastern slopes of Rocky Mountains in British Columbia	45

Crue de Juillet 1982, Rivière Wapiti, Colombie-Britannique

RESUME

Le 14 juillet 1982, une forte précipitation a provoqué une crue, ayant une période de récurrence estimée à 35 ans, sur la rivière Wapiti en Colombie-Britannique. La division des relevés hydrologiques du Canada y opérait une station de jaugeage appelée "Wapiti River above Mistanusk Creek". Il a été estimé, par la méthode "pente-surface", que le débit maximum instantané a atteint $700 \text{ m}^3/\text{sec}$ à cet endroit. Cette estimation doit être considérée approximative, parce que la pente de l'écoulement a été obtenue à l'aide des marques laissées par les hautes eaux et le coefficient de rugosité a été estimé, par similarité, avec d'autres cours d'eau ayant une rugosité connue. Lors du débit de pointe, des conditions de refoulement ont été causées soit par un affluent en aval, un embâle, ou la configuration de la rivière.

Les installations de la division des relevés hydrologiques ainsi que la graphique indicatif du niveau d'eau ont été perdus.

De sévères inondations et érosion se sont produites à plusieurs endroits, en aval, dans la province de l'Alberta.

Flood of July 1982 on the Wapiti River, British Columbia

ABSTRACT

On July 14, 1982 runoff from heavy precipitation produced an estimated 35 year flood on the Wapiti River in British Columbia. An instantaneous maximum discharge of $700 \text{ m}^3/\text{sec}$ was estimated by the slope-area method at the Water Survey of Canada gauging station "Wapiti River above Mistanusk Creek". The estimate is at best an approximation because the water surface slope is obtained from highwater marks and a roughness coefficient is estimated from similarity with other streams of known roughness. Backwater conditions were present at the time of peak discharge and were caused either by a downstream tributary, log jam or the channel configuration.

Water Survey's cableway and recorder were lost with the recorder chart of river stage. Severe flooding and erosion were caused in many locations downstream in Alberta.

1. INTRODUCTION

The Wapiti River rises on the eastern slopes of the Rocky Mountains in Central British Columbia northeast of Prince George and near the Alberta border. The gauging site is located 84 kilometers downstream from the source, and contains a basin with a drainage area of 2,400 square kilometers. After crossing into Alberta, the Wapiti River flows approximately 150 kilometers until it joins the

Smoky River near Grande Prairie. The location of the basin is shown on Figure 1.

Streamflow records have been collected in the British Columbia portion of the basin above Mistanusk Creek since September 1977. The station located downstream near Grande Prairie, Alberta was established in November 1960, the drainage area of this site is 11,300 square kilometers. The available data on high flows at this station are used in the flood estimates for the B.C. site.

The purpose of this report is to help provide an understanding of the causes of floods in this region, to estimate the peak discharge through indirect means, to estimate the frequency of the flood and to present an analysis of the basic data.

The report provides a brief description of the physiography and climate of the area, the events leading up to the flood, the extent of flooding and flood damage in British Columbia.

1.1 PHYSIOGRAPHY

The basin consists of three main physiographic subdivisions: the mountains, the foothills and the plains. Refer to Figure 2, Physiographic Regions. The front ranges of the Rocky Mountains extend from the divide east to the foothills. The gradient is steep, the rivers dropping from 2,150 meters to 1,200 meters in

less than 16 kilometers. The area contains many rugged peaks but none over 2,290 meters in elevation. There are few glaciers with only one supplying some meltwater to Belcourt Creek, a tributary of the Wapiti.

The foothills, composed mainly of rounded hills contain outlying mountain ridges nearly as high and rugged as the front range. The gradient changes drastically, the rivers dropping from 1,200 meters to 900 meters in 50 kilometers.

The plains region is characterized by buttes and mesas. The gradient is very flat, the rivers dropping less than 0.25 meters per kilometer. The drainage pattern consists of parallel stream systems flowing east and transecting the regional physiographic trends.

Stream profiles are shown on Figure 3, for the Wapiti River and its main tributaries upstream of the British Columbia-Alberta border.

1.2 CLIMATE

The climate of the basin is dominated by continental influences. The winter continental Arctic air masses move down from the Mackenzie Valley for periods of varying length. These cold anticyclone air masses create conditions which are unfavourable to precipitation, cold and dry in winter and warm and dry in summer.

The basin is characterized by semi-arid conditions in the plains and by progressively higher precipitation in the foothills and mountains.

The tracts of the various pressure systems are not fixed and vary considerably throughout the year. In winter many depressions, caused by the Aleutian Low, enter British Columbia from the ocean and drift eastward over the western mountains, and if dominant, on across the plains. As the land warms in the spring the frontal zones shift to the north and the depressions travelling across British Columbia are less affected by the cold Arctic air mass.

1.2.1 Precipitation

Annual precipitation is generally less than 500 millimeters to the east of the Rocky Mountain foothills but increases westward through the foothills and into the mountains to over 1,500 millimeters. Precipitation has an even distribution from September to May but has a pronounced summer maximum. The winters are relatively dry because of the low moisture content of the Arctic air which covers the area during this period. Summer precipitation is predominantly of the convection type so that spotty distribution within the area is characteristic as are high intensities for short periods. In general precipitation in the form of hail is infrequent.

1.2.2 Snowfall

The mean annual snowfall in the area ranges from 150 centimeters on

the plains to over 350 on the highest mountain ranges. Snowfall starts in October, reaches a peak in January and February, recedes in March and has a pronounced drop in April decreasing as summer approaches. Snow course locations are shown in Figure 4 and snow course data for 1982 are shown in Table 1.

TABLE 1 Snow Course Data for Selected Stations
(Snow Survey Bulletins, 1982)

Course Number	Snow Course Name	Elevation in Meters	Date (1982)	Snow Depth in cm	Water Equivalent in mm	Period of Record
4A20	Monkman Creek	1550	Mar 1	209	622	4 yrs
			Apr 1	NM	-	-
			May 1	NM	-	-
			May 15	NM	-	-
			Jun 1	148	689	8 yrs
4A22	Sukunka River	1140	Mar 1	109	235	5 yrs
			Apr 1	131	347	7 yrs
			May 1	102	327	8 yrs
			May 15	NM	-	-
4A02	Pine Pass	1430	Mar 1	300	932	20 yrs
			Apr 1	296	1106	20 yrs
			May 1	282	1193	21 yrs
			May 15	251	1173	14 yrs
			Jun 1	203	1066	12 yrs
4A18	Mount Sheba	1490	Mar 1	258	803	12 yrs
			Apr 1	240	884	13 yrs
			May 1	238	923	13 yrs
			May 15	220	945	7 yrs
			Jun 1	183	901	7 yrs
4A25	Ft. St. John	690	Mar 1	84	191	8 yrs
			Apr 1	76	210	8 yrs
			May 1	0	0	8 yrs
4A26	Portage Mountain	760	Mar 1	75	156	8 yrs
			Apr 1	68	173	8 yrs
			May 1	39	107	8 yrs

NM - Not Measured

1.2.3 Temperature

Temperatures rise rapidly from winter to summer and fall again as fast to winter. The mean daily temperature is below 0°C in all five winter months from November to March and above 10° from May to September. The year may be divided into winter and summer with one month for spring and one month for autumn.

1.3 GEOLOGY

The Rocky Mountains consist of a series of parallel ridges and valleys roughly aligned in a northwest to southeast direction. They are underlain by faulted and folded sedimentary rocks. The mountain massifs are formed of the harder rocks, limestone, dolomites and quartzites while the valleys are principally underlain with the softer sediments, shales and sandstone. The Rocky mountain foothills are characterized by series of subparallel ridges and valleys situated between the plateau to the east and the Rocky Mountains to the west. These ridges and valleys are dissected by the larger streams flowing in a northeasterly direction. The streams and their tributaries have entrenched themselves in deeply incised canyons. The foothills are underlain by faulted and folded shales and sandstone.

1.3.1 Soil Types

Most of the soils are of glacial origin. Glacial-fluvial and glacial-lacustrine deposits are thin on the higher elevations but

tend to be of considerable depth in the larger stream valleys. Where river valleys are broad and subject to periodic flooding the formation of peat bogs and swamps has been promoted.

1.3.2 Vegetation

The foothills and Rocky Mountains contain mostly evergreens at the higher elevations and deciduous lower down. The forest cover consists of Fir, Spruce, Lodgepole Pine, Cottonwood, Birch and Aspen; some of these species can be found at elevations 1,800 meters above sea level.

1.4 WAPITI RIVER BASIN IN BRITISH COLUMBIA

The Wapiti River Basin at the Water Survey of Canada gauging site comprises three basins of comparable size originating in the Rocky Mountains and flowing through the foothills with their confluence near the edge of the Alberta plateau. Wapiti River itself extends approximately 84 kilometers above the measuring site; Red Deer Creek is 70 kilometers and Belcourt Creek is 84 kilometers long respectively above their confluence. Refer to Figure 5 for detailed map of basin.

The average annual runoff of the Wapiti River in British Columbia, for the four-year period of record, is 15.8 L/km^2 . This is the contribution from $2,400 \text{ km}^2$. The second station on the Wapiti located in Alberta near Grande Prairie approximately 120 km down-

TABLE 2 Mean Annual and Maximum Run off for Selected Basins

Station Number	Name	Drainage Area km ²	Years of Record	Mean Annual Flow m ³	Mean Annual Runoff Litres/km ²	Maximum Daily Flow m ³	Maximum Daily Runoff Litres/km ²
07FB001	Pine River at East Pine	12100	22	203	16.8	3960	327
07FB007	Sukunka River above Chamberlain Creek	927	5	23.9	25.8	266	287
07FB003	Sukunka River near the Mouth	2510	5	52.5	20.9	704	280
07FB004	Dickebusch Creek near the Mouth	85.5	5	0.665	7.8	16.1	188
07FB006	Murray River above Wolverine River	2410	5	57.1	23.7	423	176
07FB002	Murray River near the Mouth	5620	5	84.6	15.1	736	131
07FB005	Quality Creek near the Mouth	29.5	5	0.207	7.0	9.23	313
07FD001	Kiskatinaw River near Farmington	3570	23	11.4	3.2	413	116
07FD007	Pouce Coupe River below Henderson Creek	2850	12	9.72	3.4	368	129
07GC001	Wapiti River above Mistanusk Creek	2400	5	38.0	15.8	650	270
07GE001	Wapiti River near Grande Prairie	11300	22	107	9.6	3680	326
07GB001	Cutbank River near Grande Prairie	842	12	Seasonal	-	506	601
07GB002	Kakwa River near Grande Prairie	3290	7	Seasonal	-	1680	511
07GA001	Smoky River below Hell's Creek	3830	15	78.9	20.6	1230	321

stream, gauges runoff from 11,300 km². Average annual runoff at this location for the 20-year period of record is 9.5 L/km². The upper area has a contribution over double that of the lower area. Annual and maximum daily runoff for stations in the region are listed in Table 2.

The physiography of the upper basin would indicate that for any storm centred over the basin the flood wave travel time would be similar for each tributary and would concentrate the flow at their confluence.

Table 3 gives the physiographic location and drainage area for selected stations in the area.

2. DESCRIPTION OF FLOOD

The annual maximum discharge of rivers draining the eastern slopes of the Rocky Mountains and foothills usually occurs from early May through July and is the result of snowmelt or rain, or rain on snow. Occasionally summer highs are caused by local convective disturbances after the snowmelt has dissipated. On rarer occasions in the fall, Pacific Lows manage to penetrate the mountain barriers and deposit high amounts of precipitation.

In the Wapiti basin spring runoff normally commences during the latter part of April and peak snowmelt flows are reached in May or early June. Secondary peaks often occur late in June and July as a

TABLE 3 Physiography of Selected Streamflow Stations
Draining Eastern Slopes of Rocky Mountains

Station Number	Name	Drainage ² Area km	Physiographic Region Represented in Watershed
07FB001	Pine River at East Pine	12100	Rocky Mountains and Rocky Mountain Foothills
07FB007	Sukunka River above Chamberlain Creek	927	Rocky Mountains
07FB003	Sukunka River near the Mouth	2510	Rocky Mountains and Rocky Mountain Foothills
07FB004	Dickebusch Creek near the Mouth	85.5	Rocky Mountain Foothills
07FB006	Murray River above Wolverine River	2410	Rocky Mountains and Rocky Mountain Foothills
07FB002	Murray River near the Mouth	5620	Rocky Mountains and Rocky Mountain Foothills
07FB005	Quality Creek near the Mouth	29.5	Rocky Mountain Foothills
07FD001	Kiskatinaw River near Farmington	3570	Alberta Plateau
07FD007	Pouce Coupe River below Henderson Creek	2850	Alberta Plateau
07GC001	Wapiti River above Mistanusk Creek	2400	Rocky Mountains and Rocky Mountain Foothills
07GE001	Wapiti River near Grande Prairie	11300	Rocky Mountains, Rocky Mountain Foothills and Alberta Plateau
07GB001	Cutbank River near Grande Prairie	842	Rocky Mountain Foothills
07GB002	Kakwa River near Grande Prairie	3290	Rocky Mountains, Rocky Mountain Foothills and Alberta Plateau
07GA001	Smoky River below Hells Creek	3830	Rocky Mountains and Rocky Mountain Foothills

result of heavy rains. Occasionally these peaks are greater than the snowmelt peaks. In some years consistent summer rains keep the runoff high throughout the summer period. Streamflow recession is rapid once the cold temperatures appear.

The very high summer flows can be the result of heavy rain that follows a high runoff caused by previous snowmelt or rainfall. Hydrographs illustrating the different causes of peaks are shown on Figures 6 through 9.

2.1 Rainstorm of July 13-15, 1982

The flood of July 14, 1982 was caused by heavy rainfall which covered the entire Wapiti River basin. The distribution was uneven, as the storm was centred over the eastern edge of the basin. Table 4 lists the precipitation for the period July 1 to 31, 1982 for selected stations in the area. Figure 10 shows the mass rainfall curves for the month of July for these stations.

2.1.1 Meteorological Description of the Rainstorm of July 13-15, 1982 over Smoky River Basin (Wapiti River)

A chronological description of the meteorological occurrences, prepared by the Scientific Support Unit of the Atmospheric Environmental Service (AES), Edmonton, Alberta follows:

1. Most of the precipitation fell over British Columbia on the 13th and 14th of July and over Alberta on the 14th and 15th.
2. Two axes of precipitation, one from the south and one from the

TABLE 4 Daily Precipitation in Millimeters - July 1-31, 1982
(B.C. Air Studies, Victoria & A.E.S. Edmonton)

Day	Murray Townsite	Wolverine	Quintette 2	Monkman Camp	Stoney Lake	Duke Mountain	Nose Mountain
Elev.	835 m	823 m	1541 m	1000 m	1085 m	1593 m	1574 m
1	4.0	6.0	2.5	4.2	5.0	0.0	0.3
2	3.0	3.0	6.0	1.0	5.5	2.0	0.5
3	16.5	M	12.0	22.0	32.5	18.2	28.2
4	17.5	M	15.0	18.2	17.0	39.5	27.0
5	0.5	M	20.0	3.7	3.0	7.7	3.6
6	3.0	M	9.0	3.0	4.0	10.2	7.5
7	1.0	M	1.0	1.0	2.0	1.0	T
8	0.0	M	0.0	0.0	0.0	0.0	T
9	0.0	M	0.5	0.5	2.0	0.0	1.3
10	0.0	M	0.5	0.0	0.0	0.0	0.0
11	0.5	M	0.0	0.3	0.0	0.0	3.0
12	0.0	M	0.0	0.0	0.0	0.0	1.0
13	34.5	42.0	68.0	11.7	25.0	11.0	5.1
14	66.0	56.0	86.0	76.0	90.0	83.0	138.9
15	0.5	0.0	2.0	19.7	0.5	28.5	27.9
16	0.0	0.0	0.0	0.0	0.5	1.0	T
17	0.0	0.0	0.0	0.0	0.0	0.0	0.0
18	0.0	0.0	4.0	1.2	1.0	0.0	0.0
19	4.0	1.0	0.5	1.2	0.0	3.0	0.0
20	7.5	12.0	9.5	4.0	5.0	5.0	0.5
21	42.0	37.0	40.0	12.7	5.0	21.7	6.3
22	2.0	2.0	1.5	2.7	1.5	1.7	T
23	0.5	0.0	0.5	0.8	1.5	2.2	T
24	0.0	0.5	0.0	0.0	4.5	0.0	0.0
25	0.0	0.0	0.0	0.0	0.0	0.0	0.0
26	0.0	0.0	0.0	0.0	0.0	0.0	0.0
27	0.0	0.0	0.0	1.2	0.0	0.3	0.0
28	0.5	0.0	0.0	0.0	1.0	0.0	0.0
29	0.0	2.0	0.0	0.0	0.2	0.0	0.0
30	0.0	0.0	0.0	0.0	0.0	0.0	0.0
31	26.0	4.0	20.0	13.2	13.5	8.5	15.2
TOTAL	229.5	-	298.5	198.3	220.2	244.5	266.3

M - Missing

T - Trace (less than 0.2 mm)

west, converged in an area to the west of Nose Mountain Lookout. It is estimated that amounts in excess of 150 millimeters of precipitation could have fallen over British Columbia within a 50 to 75 kilometer radius of Nose Mountain. There are no reporting stations to substantiate these amounts (Figure 11).

3. Main contributors to precipitation include the following:

- a) 500 millibar (cold low) which moved from Georgia Strait at July 14/12Z to 50 n.mi. south of Vernon (15/00Z) to 120 east of Vernon (15/12Z) to 60 southwest of Calgary (16/00Z) to 50 northwest of Lethbridge by 16/12Z (Figure 12).
- b) surface low which moved from near Revelstoke (14/00Z) to 60 n.mi. west of Calgary (14/12Z) to 50 southwest Edmonton (15/00Z) then curled north-westward to 75 west Grande Prairie (15/12Z) and thereafter weakening and moving rapidly eastward to central Saskatchewan by (16/00Z).
- c) combined easterly to northeasterly upslope flow from low to high levels across the Smoky basin as a result of the tracks of the upper and surface low pressure systems.
- d) high surface dew points (12-16°C) across the area for the duration of the storm.
- e) tongue of warm air aloft stretching across the basin.
- f) associated instability.

3. ANALYSIS OF FLOOD DATA

3.1 Wapiti River above Mistanusk Creek

This station was originally equipped with an automatic recorder, cableway and rain gauge. All of this equipment was lost on July 14, 1982. Figures 13 & 13a show effects of the flood at the gauging site.

A slope area survey was done on October 12, 1982 and established the peak gauge height at an elevation (local levels) of 5.220 meters and a slope of 0.00082.

The highest measured flow prior to July 1982 was $205 \text{ m}^3/\text{s}$ at a gauge height of 1.920 metres. The maximum daily discharge at this station prior to the flood was $371 \text{ m}^3/\text{s}$, on June 4, 1980, which corresponds to a gauge height of 2.480 meters. The maximum instantaneous gauge height of 2.620 meters was observed also on June 4, 1980, corresponding to an estimated discharge of $418 \text{ m}^3/\text{s}$.

3.2 Estimation of Flood Discharge

An examination of the hydrographs of neighbouring streams and the precipitation isohyetal map shows that the storm was centred over the extreme lower portion of the upper Wapiti basin in British Columbia and extended over the lower portion of Narraway River in Alberta.

The slope of the water surface at the upper measuring site on the Wapiti River indicates that there was considerable backwater at the time of the flood. The backwater conditions could have been caused from a log jam, from channel constriction, or from the downstream tributary. The only observed phenomenon was that of the downstream tributary being extremely high.

Preliminary information at the downstream station in Alberta indicates a very large ratio between the maximum daily and the instantaneous discharge. The average value for the period of record being 1.14 compared to a value of 1.79 obtained for this flood. The previous high ratio was 1.35 in 1962.

The normal method of estimating peak flow at a stream gauging station when the stage is known is by a log-log extension of the existing rating curve. However this method cannot be used in the case where a high degree of backwater is present. In these situations the only reliable method is that of the slope-area relationship.

After the fact surveys require an accurate interpretation of high water marks to obtain water surface slope and an estimate of the coefficient of roughness. Error can occur in the estimation of both of these values. A roughness factor could be obtained from a measurement at a relatively high stage but at the present time this is impossible. However, reference can be made to Water Supply

Paper 1849 which contains many examples of stream beds for which the roughness coefficients have been established.

The water surface slope as obtained from the high water marks was calculated to be 0.00082 and the coefficient of friction was selected as 0.03. The maximum instantaneous discharge as calculated using the above values is $700 \text{ m}^3/\text{s}$ and adjusted to $650 \text{ m}^3/\text{s}$ for the maximum daily.

A more accurate study would require relationships established between stage and the following parameters: area, slope and roughness coefficient.

It is interesting to note that the average channel slope at the measuring site as obtained from contour maps was 0.0043 compared to that of 0.0033 calculated from the highest measurement taken at the site.

4. FREQUENCY CHARACTERISTICS OF THE FLOOD

A frequency curve is a graphic representation of the average occurrence of events, derived from past experience. Its intended use is to estimate how often, on the average, certain events will occur in the future.

The plotting position for each flow is determined from the

formula: $T = (N + 0.5)/(m - 0.24)$. Maximum daily discharges and frequency graphs are shown in figures 14-18 for Wapiti River above Mistanusk Creek, Wapiti River near Grande Prairie, Kakwa River near Grande Prairie, Smoky River above Hells Creek, and Cutbank River near Grande Prairie.

The maximum daily flows for the above named stations have been plotted on Gumbel extremal probability paper. The log-Pearson Type III distribution has been fitted to the flows by the method of moments. This distribution appears to fit the data very well and indicates a trend which can be extrapolated to cover a larger range.

4.1 Interpretation of Frequency Curves

The period of record for peak flow discharges is very short on the Wapiti River above Mistanusk Creek and would not be expected to be a good representation of the population of events. However, some information can be derived from the lower station, Wapiti River near Grande Prairie which has 22 years of record. Assuming that the curve is a good representation of the population distribution, the following interpretation can be made. On the average the annual maximum daily flow may exceed the flood of July 16, 1982 at intervals of 35 years in length or a probability of exceedence of 1 in 35. Assuming that records at the two stations are similar the above would hold for the upper Wapiti station also.

An envelope curve of extreme floods is shown in Figure 19.

5. CONCLUSIONS AND RECOMMENDATIONS

5.1 Causes of the Flood

The flooding on the Wapiti River at the stream gauging station above Mistanusk Creek was due to several factors:

1. heavy precipitation - rainfall from the storm was concentrated over the lower portion of the basin.
2. backwater - the presence of backwater which could have been caused by one or all of the following conditions: a log jam, an extreme high stage on a downstream tributary (Mistanusk Creek), and channel constriction.

5.2 Peak Flow

The survey of the water surface slope has provided the only means of estimating the flood peak at this station and it is only as good as the estimate of the highwater marks and the estimate of the coefficient of friction. The maximum instantaneous discharge was calculated to be $700 \text{ m}^3/\text{s}$ with probable magnitude of error of $\pm 25\%$.

6. ACKNOWLEDGMENTS

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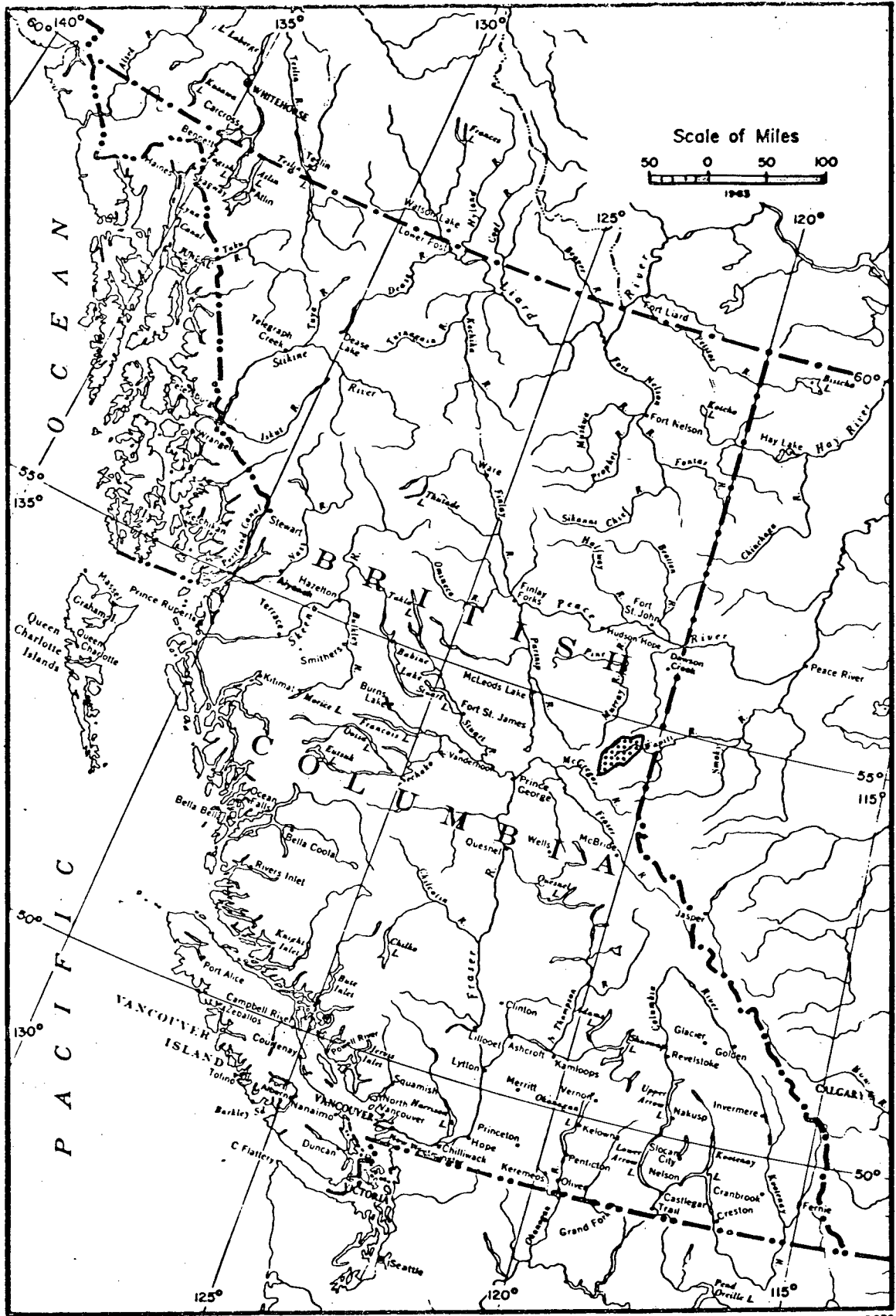


Figure 1 Map of British Columbia with Wapiti River Basin

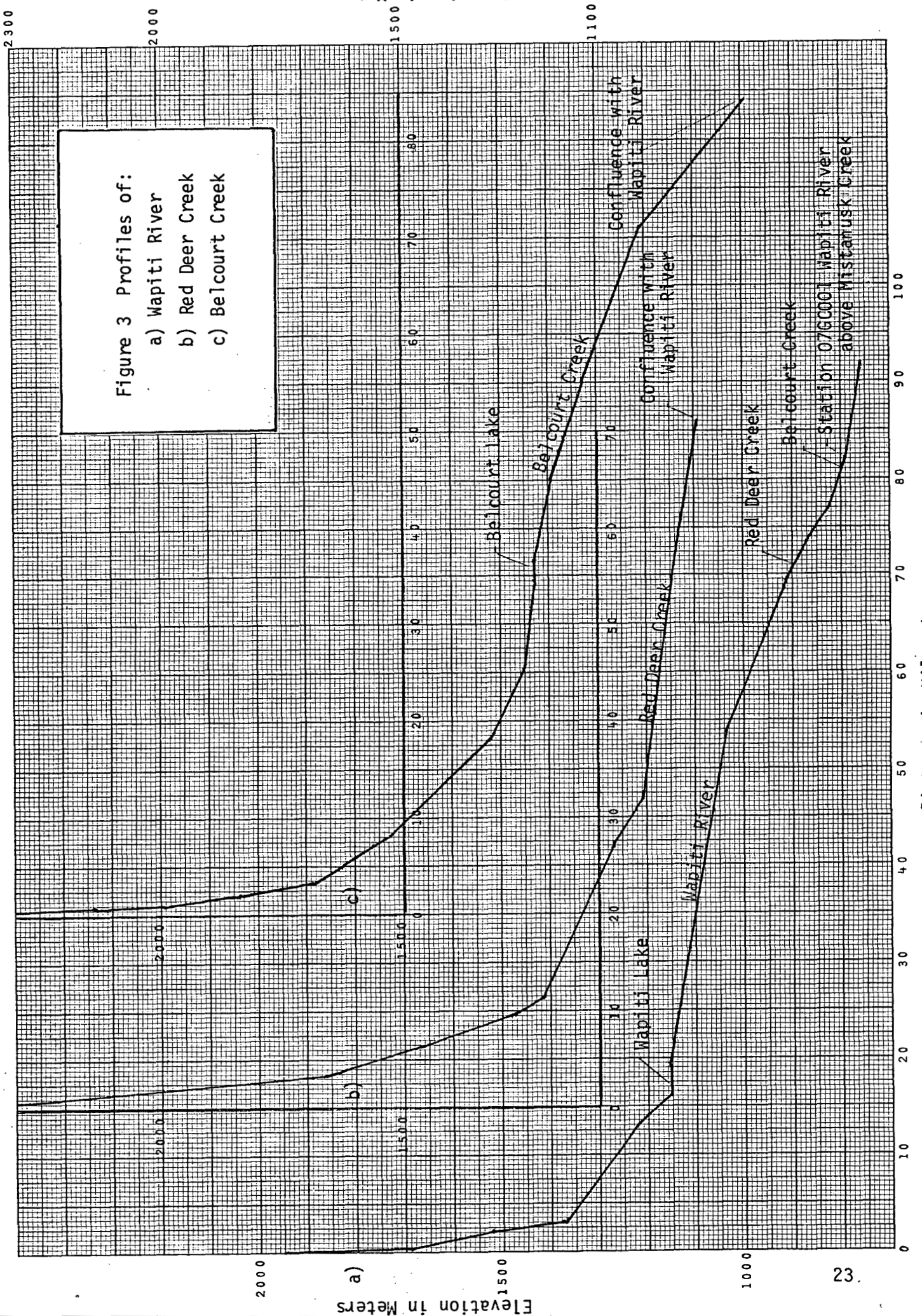


Figure 3 Profiles of:
 a) Wapiti River
 b) Red Deer Creek
 c) Belcourt Creek

Elevation in Meters

Distance in Kilometers

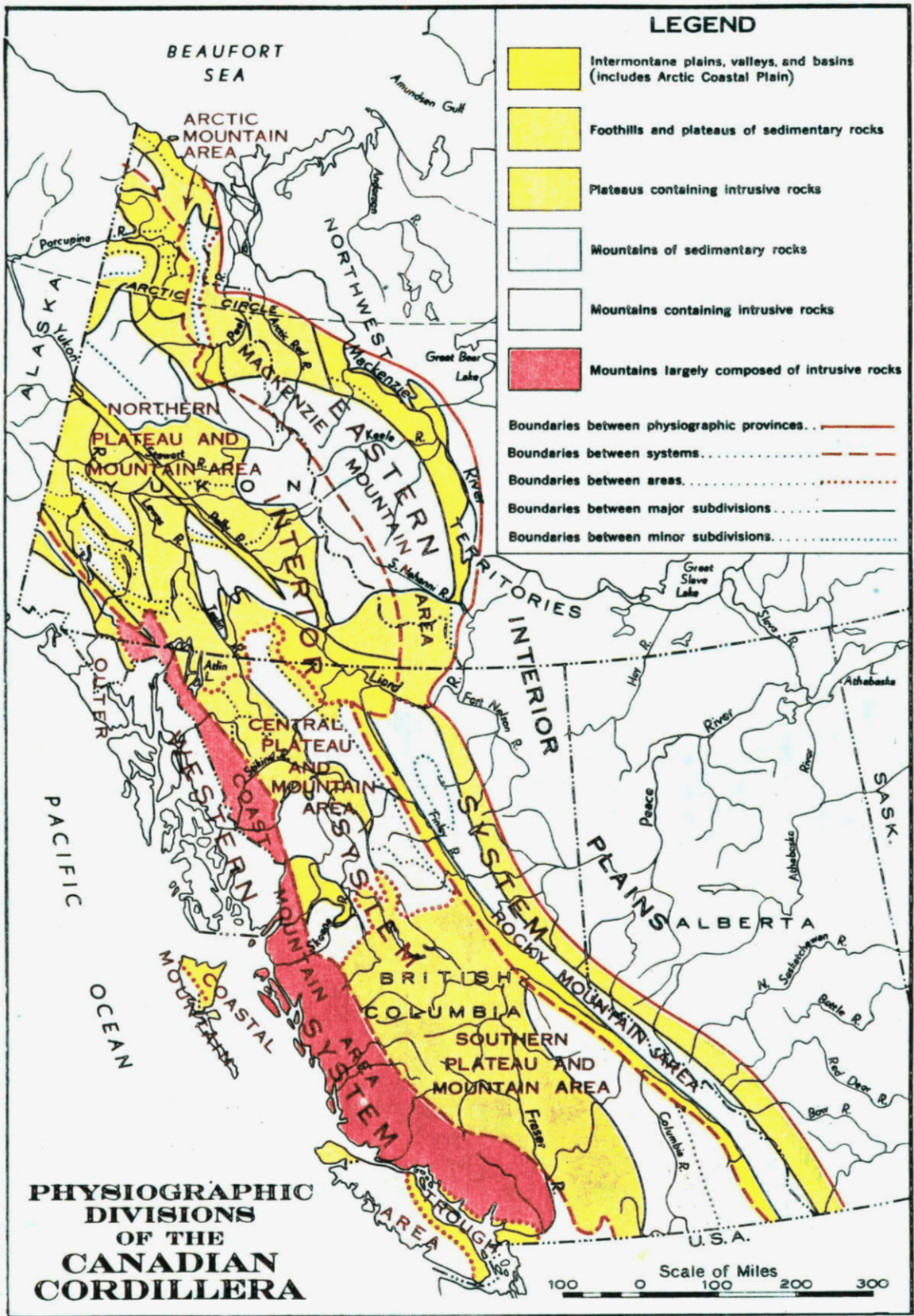


Figure 2 Geological Survey of Canada Memoir 247

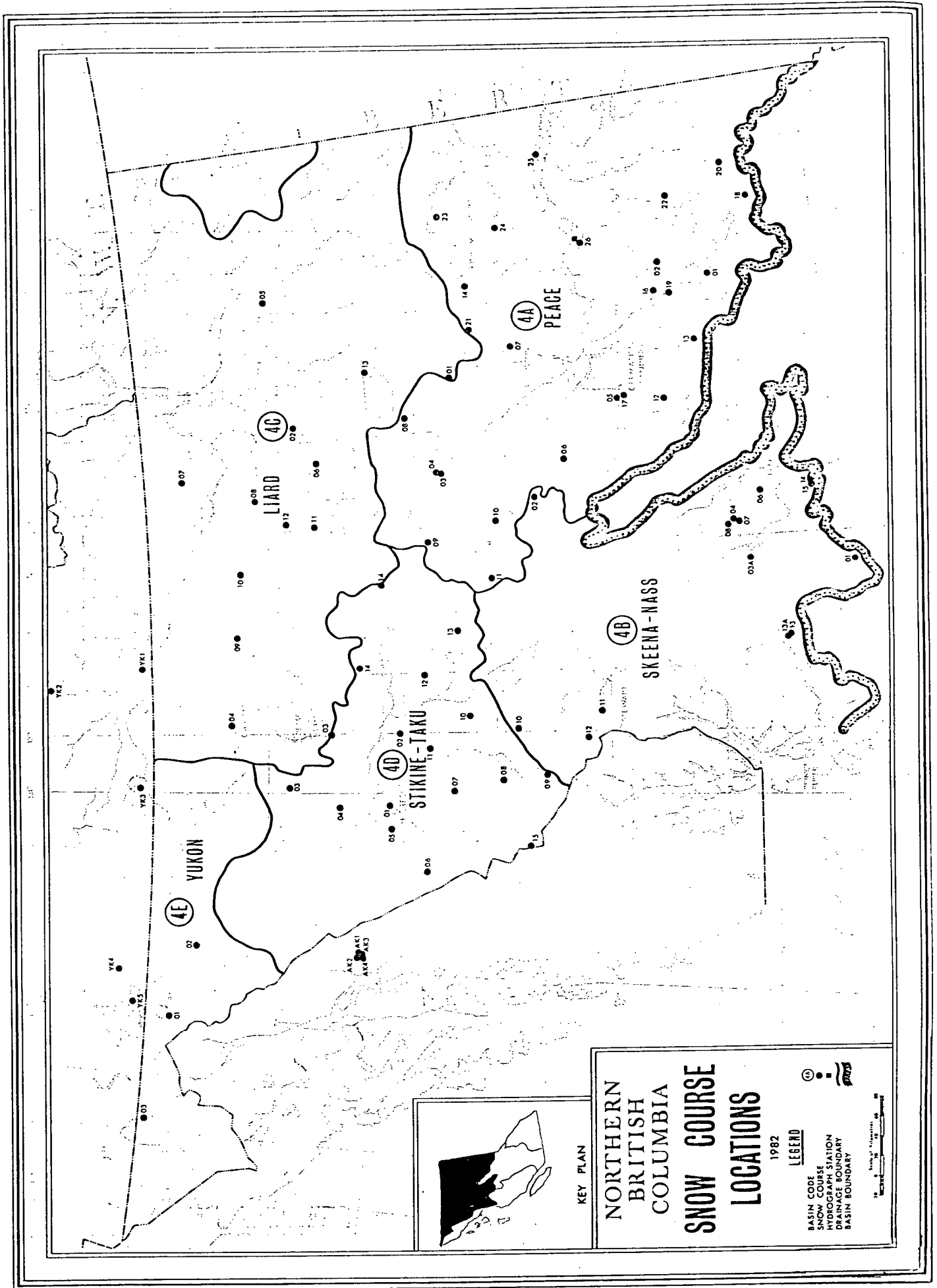


Figure 4 Snow Survey Bulletin - Ministry of Environment, Victoria, B.C.

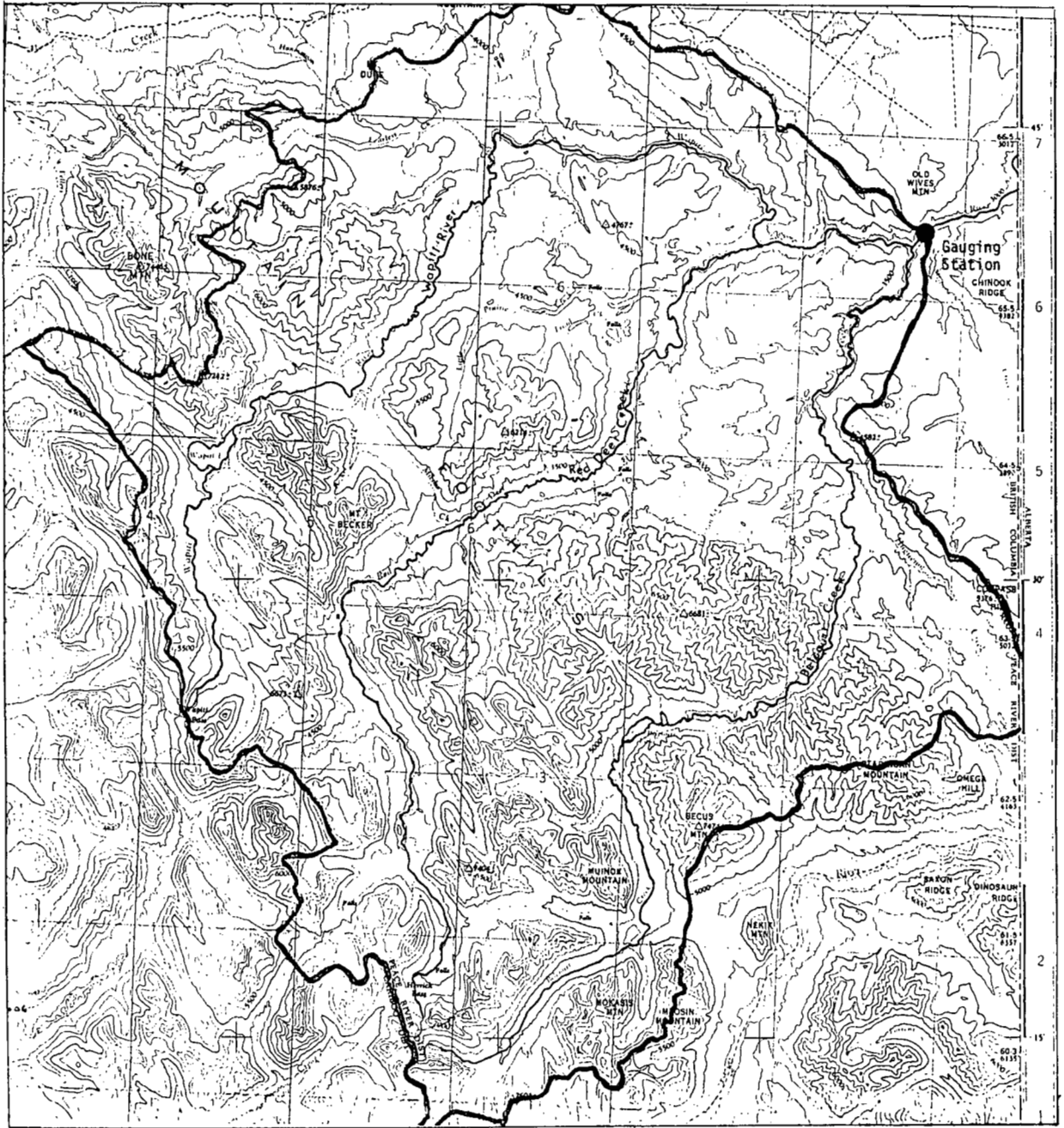
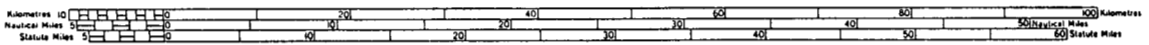


FIGURE 5. WAPITI RIVER BASIN IN BRITISH COLUMBIA



SCALE 1:500,000

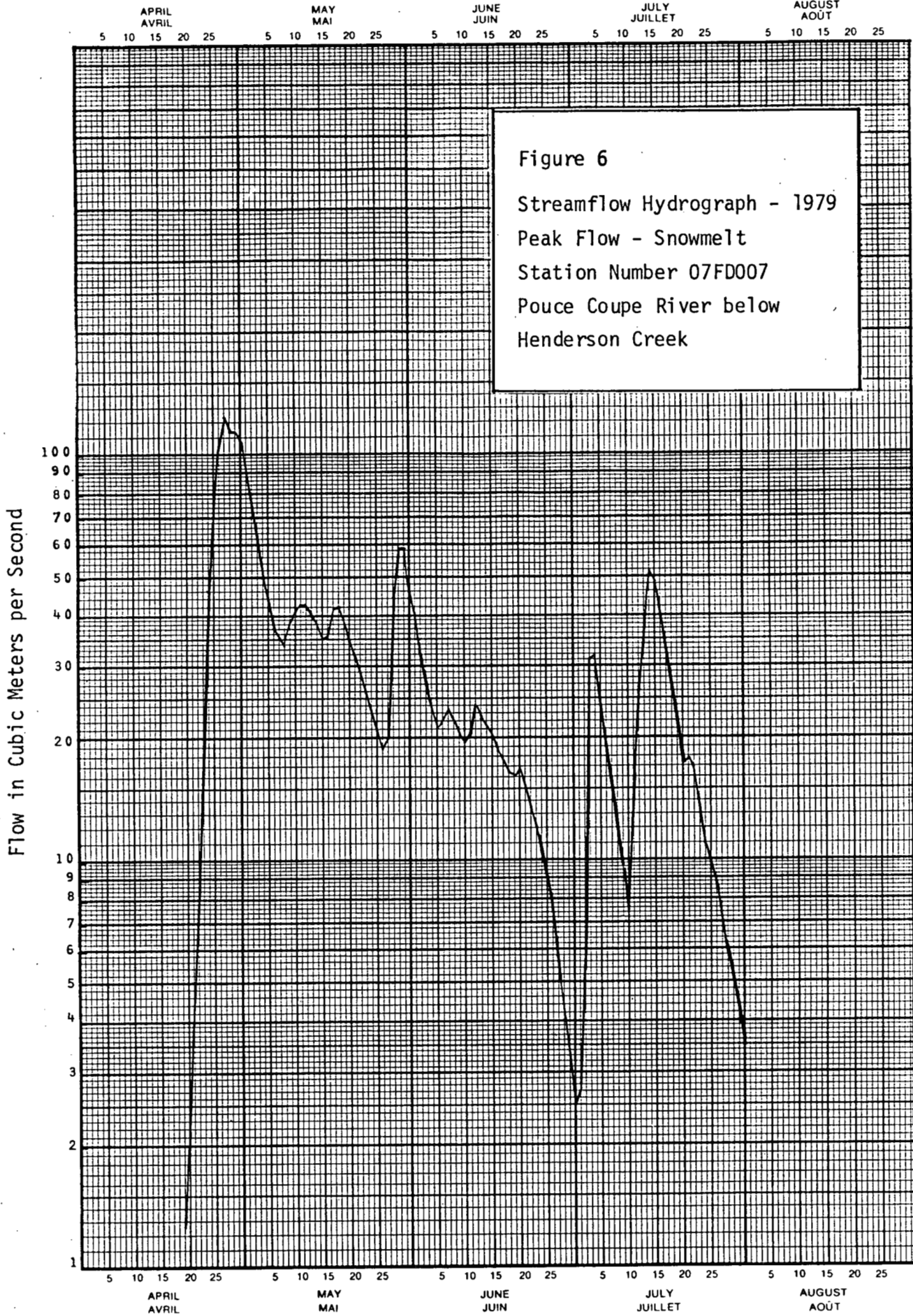


Figure 6
 Streamflow Hydrograph - 1979
 Peak Flow - Snowmelt
 Station Number 07FD007
 Pouce Coupe River below
 Henderson Creek

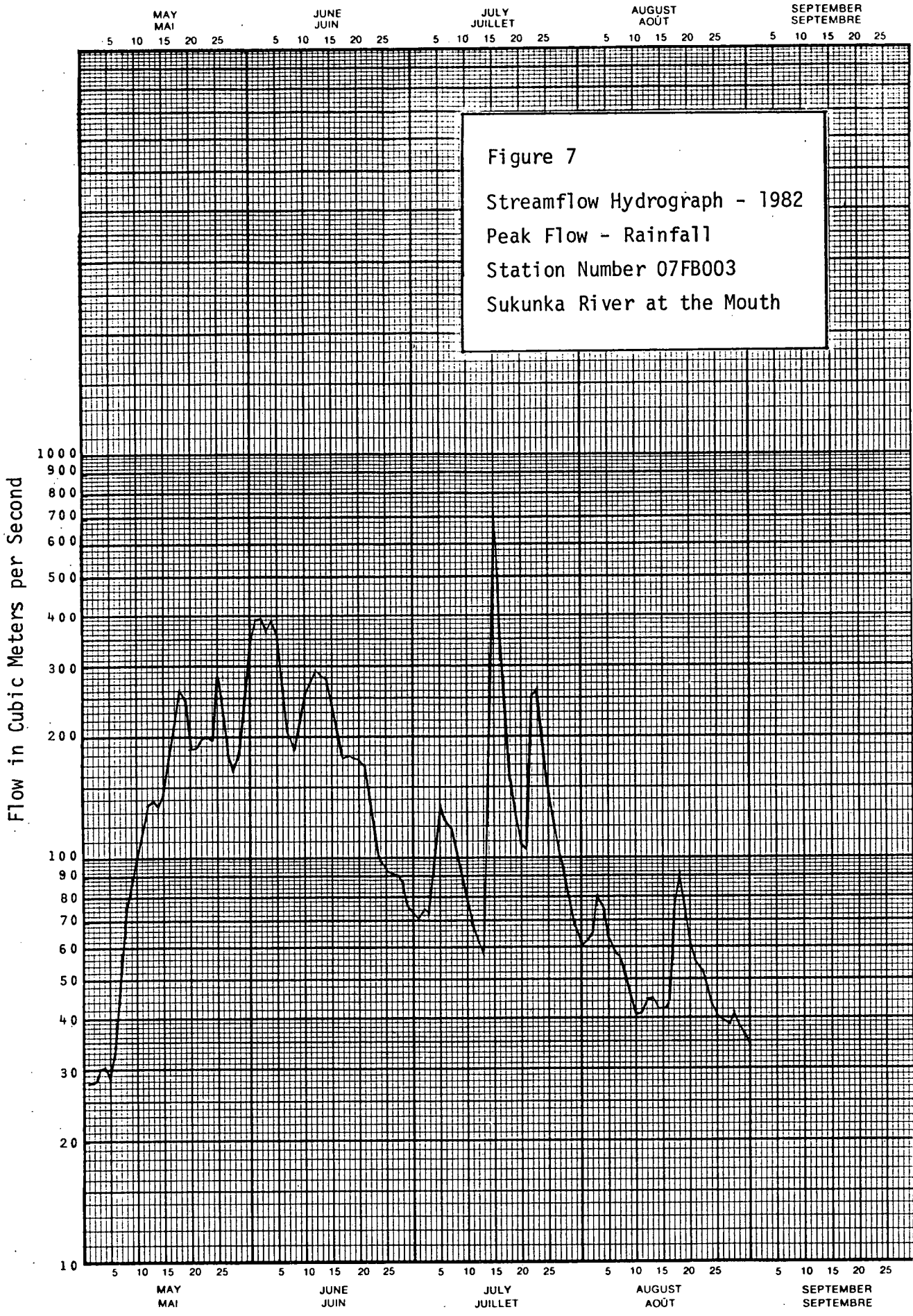


Figure 7
 Streamflow Hydrograph - 1982
 Peak Flow - Rainfall
 Station Number 07FB003
 Sukunka River at the Mouth

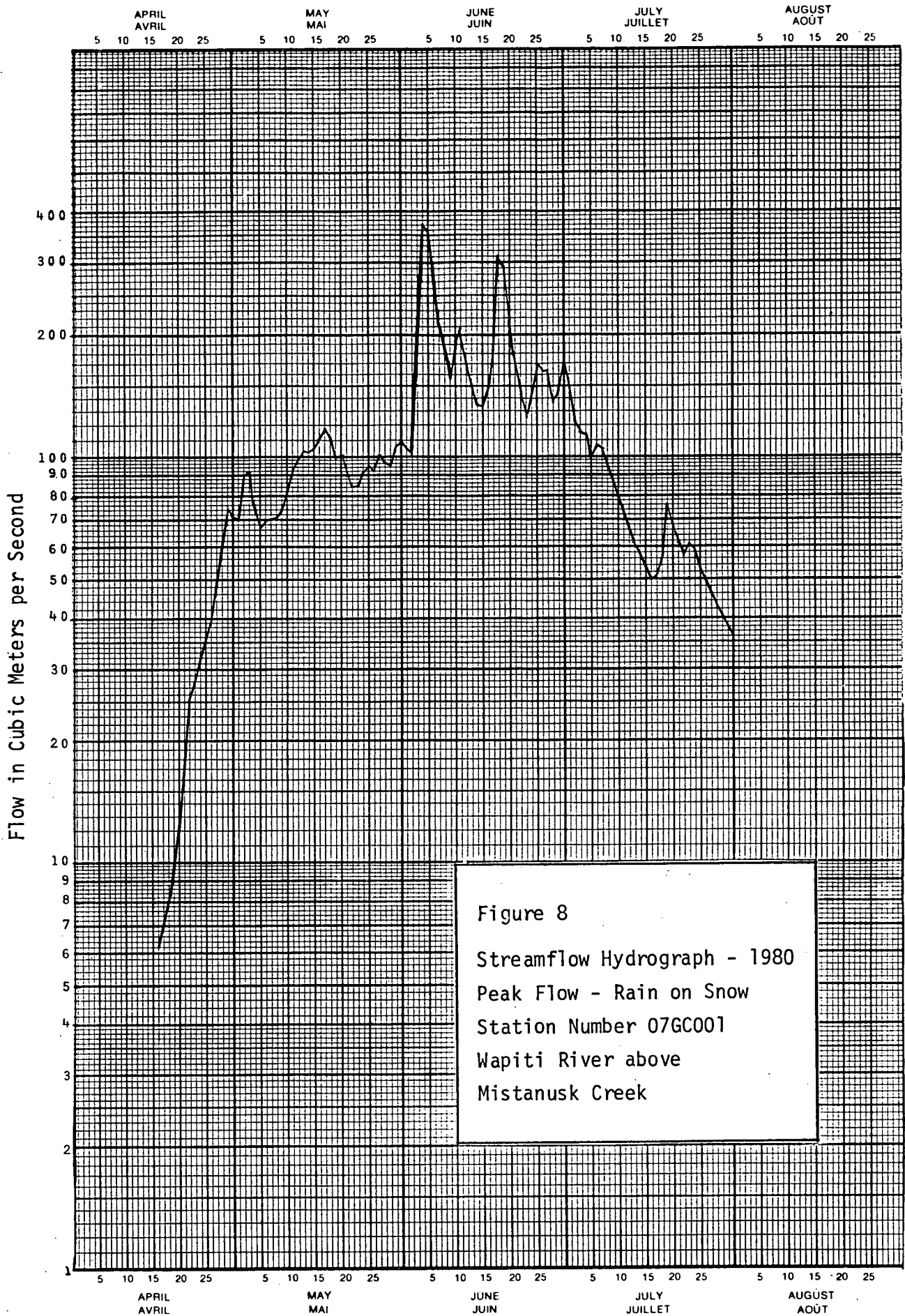
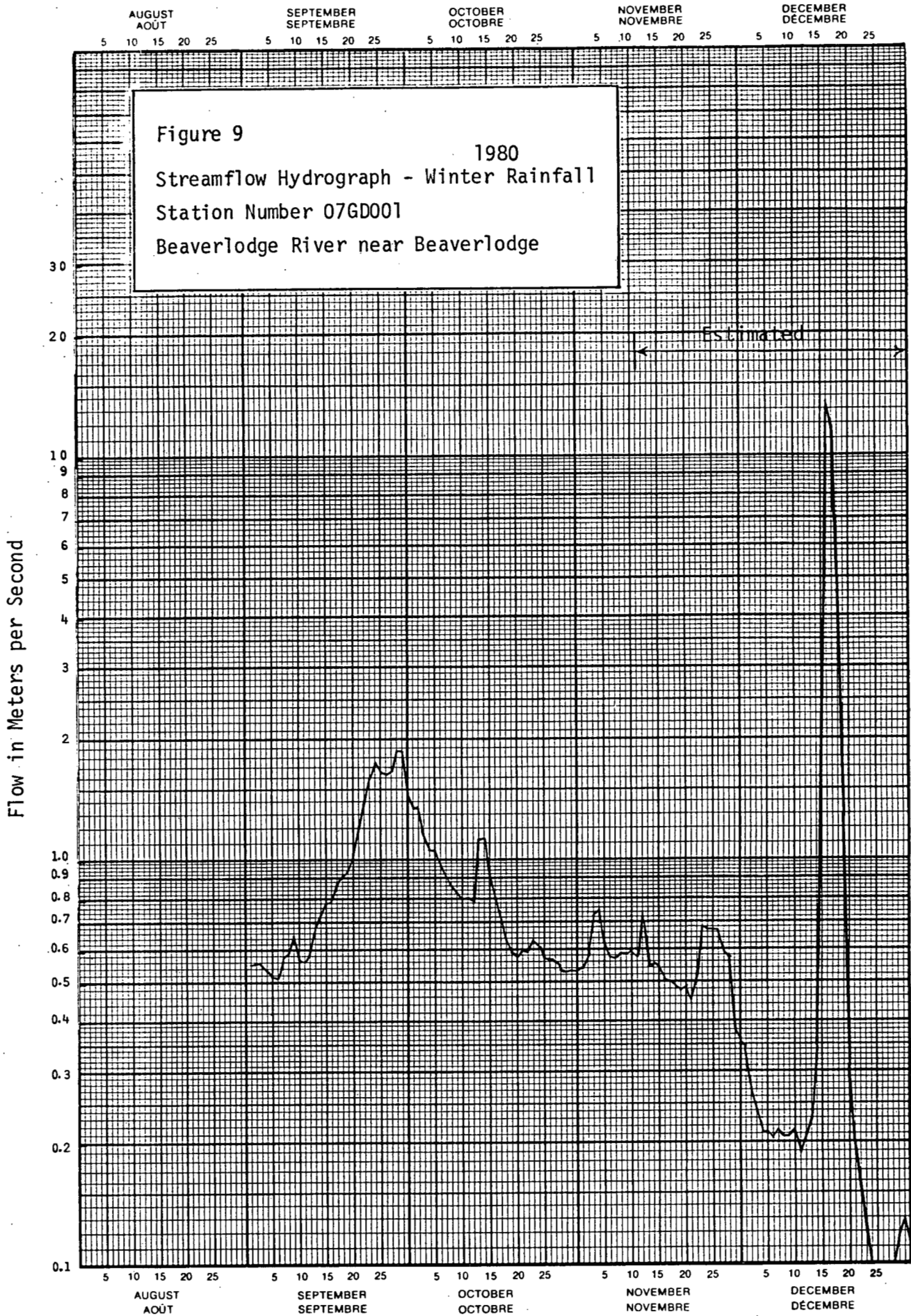


Figure 8
 Streamflow Hydrograph - 1980
 Peak Flow - Rain on Snow
 Station Number 07GC001
 Wapiti River above
 Mistanusk Creek



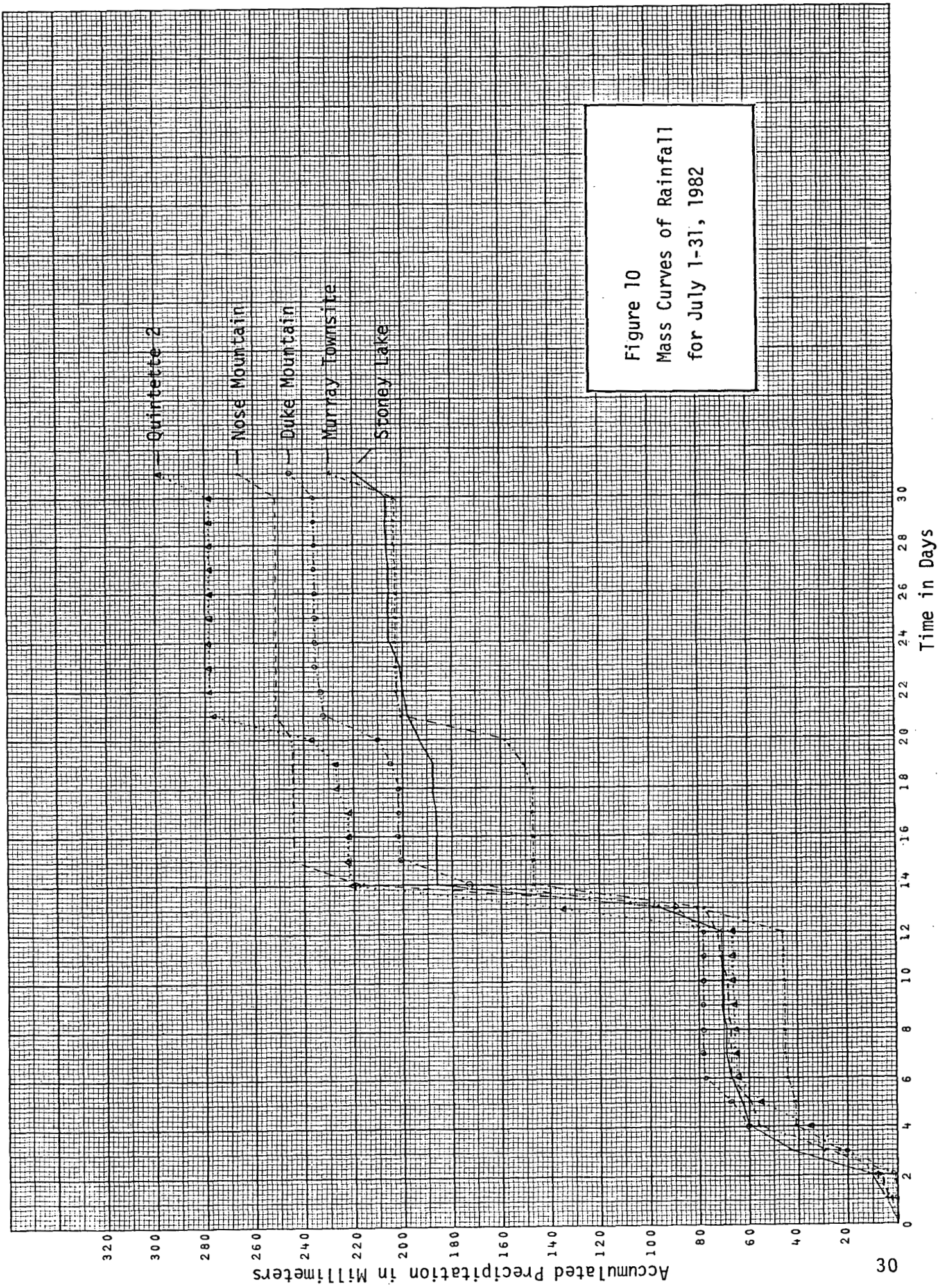


Figure 10
 Mass Curves of Rainfall
 for July 1-31, 1982

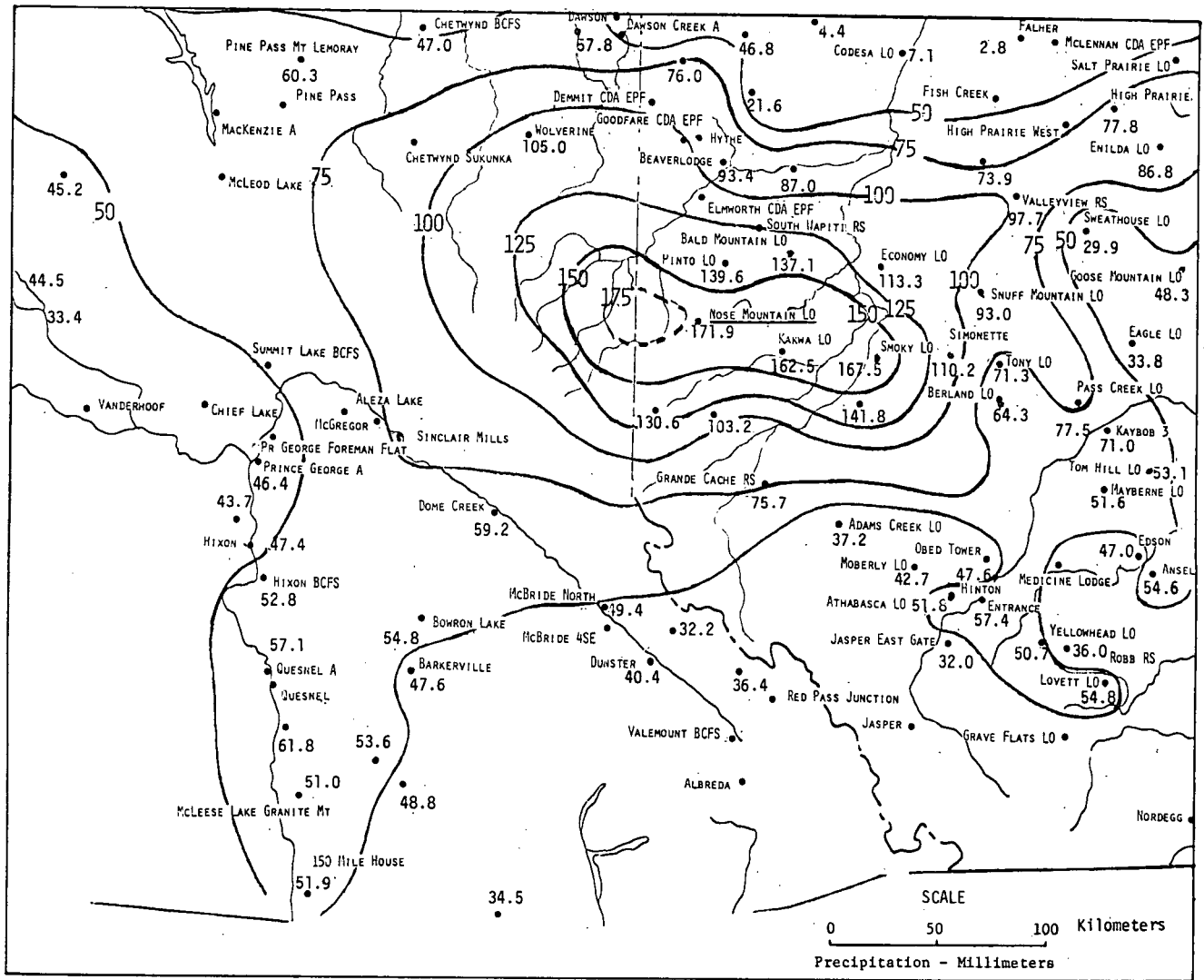


FIGURE 11 ISOHYETAL CHART - SMOKY RIVER BASIN (WAPITI RIVER) - JULY 13-15, 1982 (A.E.S. - EDMONTON)

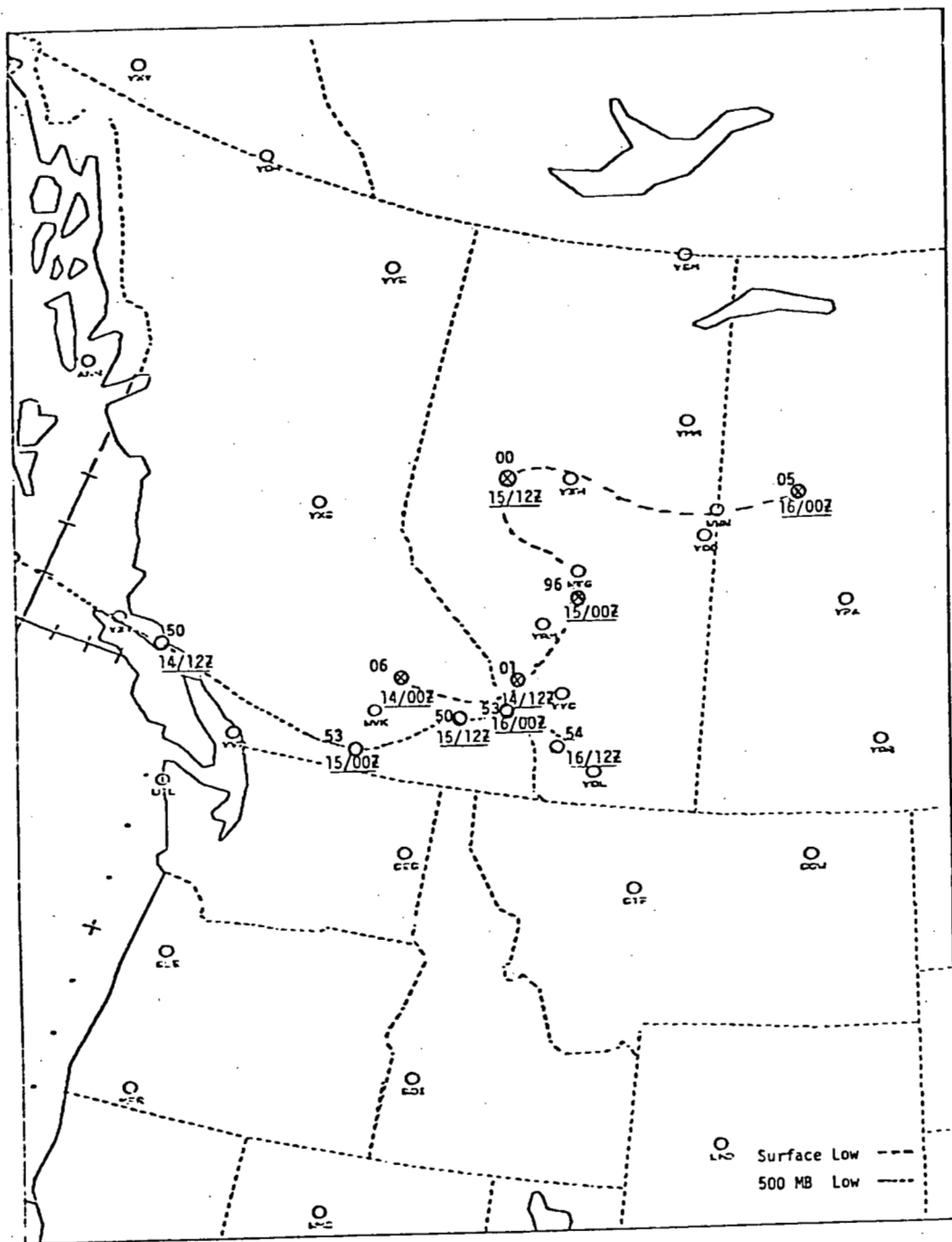
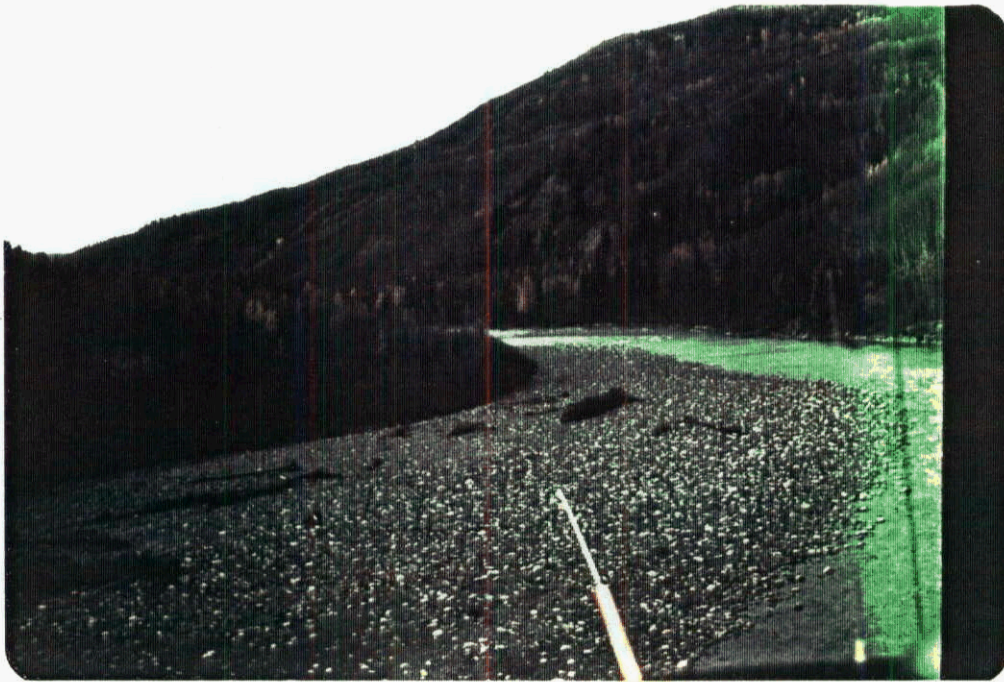


Figure 12 Surface Low and 500-Millibar Low Chart - July 14-15, 1982
(A.E.S. Edmonton)

#10 exposure



Crest Stage Gauge Located on Post



Looking Upstream from Station

Figure 13. Wapiti River above Mistanusk Creek after Flood of July 14, 1982

#4
Expos we



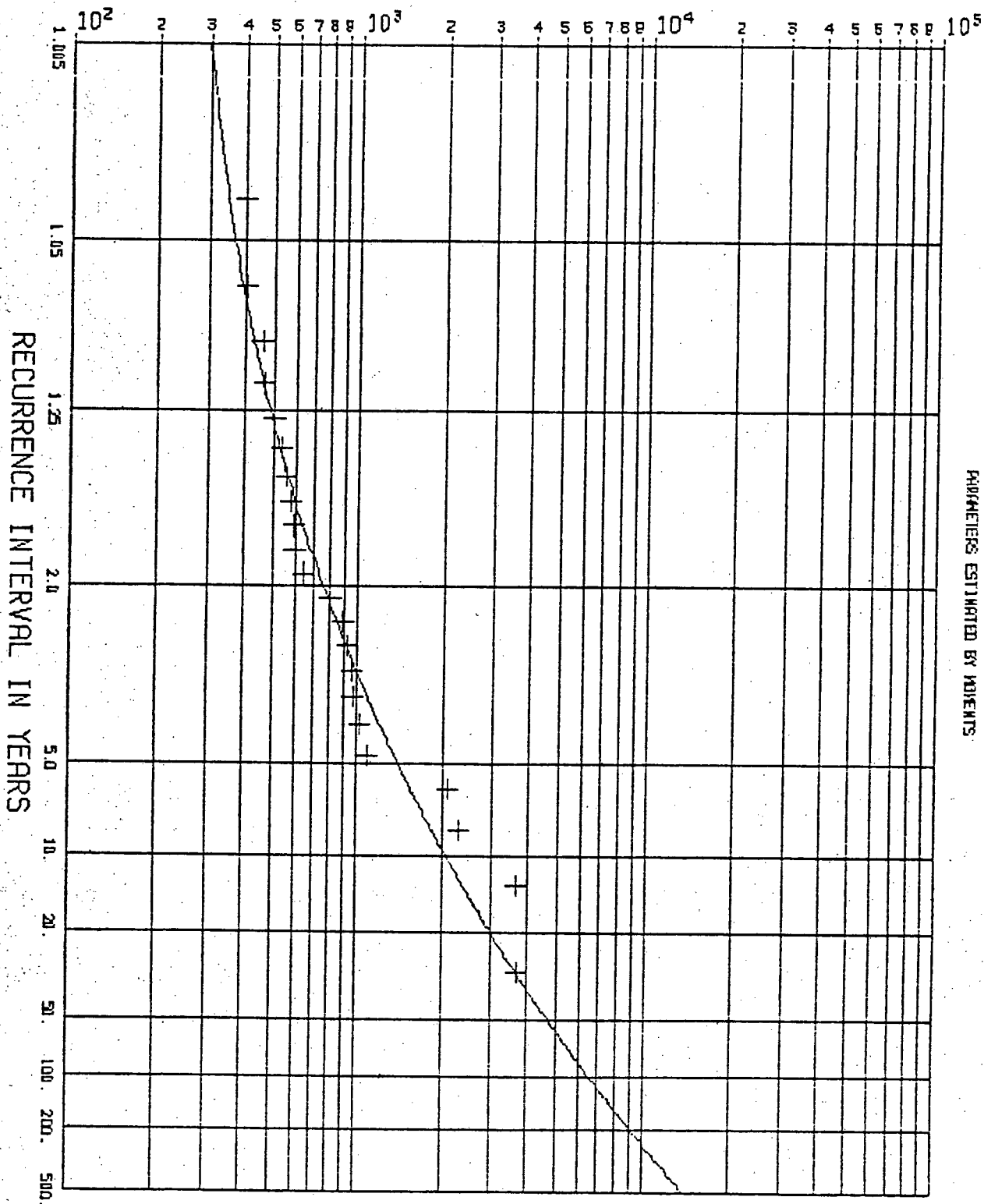
Highwater Mark on Left Bank and Downstream of Control



Highwater Mark Left Bank

Figure 13a. Wapiti River above Mistanusk Creek after Flood of July 14, 1982

DISCHARGE (M3/S)



07GE001 WAPITI RIVER NEAR GRANDE PRAIRIE
LOG PEARSON TYPE III DISTRIBUTION
PARAMETERS ESTIMATED BY MOMENTS

M A X I M U M D A I L Y M E A N F L O W S

07GC001 WAPITI RIVER ABOVE MISTANUSK CREEK

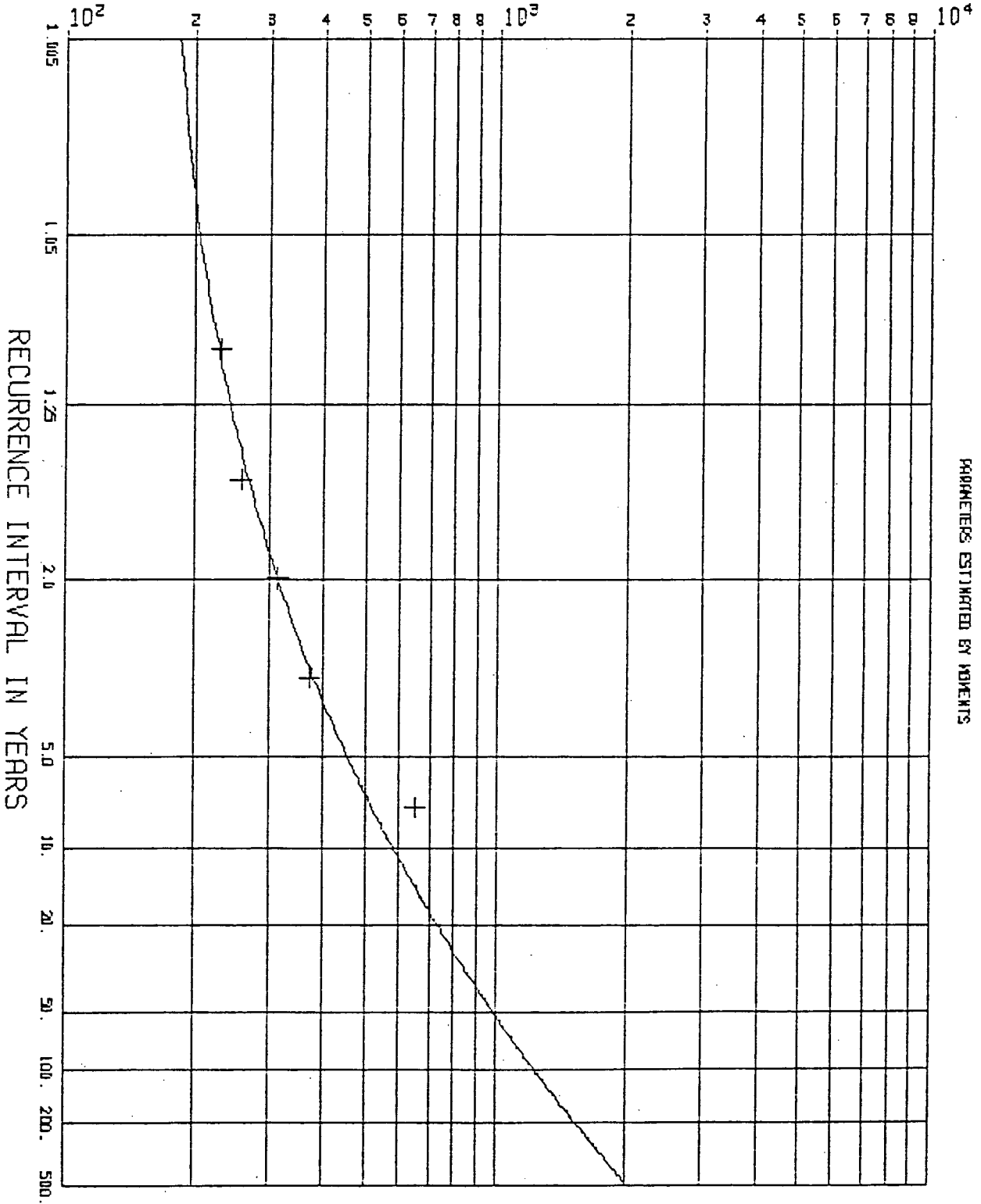
DATE	MAXIMUM DAILY FLOW IN M3/S	RANK	RECURRENCE INTERVAL IN YEARS	MAXIMUM DAILY FLOW IN M3/S	YEAR
JUL 11 1978	231.0	1	7.237	650.0	1982
JUN 05 1979	314.0	2	3.125	371.0	1980
JUN 04 1980	371.0	3	1.993	314.0	1979
MAY 26 1981	258.0	4	1.463	258.0	1981
JUL 14 1982	650.0	5	1.155	231.0	1978

MEAN ANNUAL FLOOD: 364.8 M3/S

DRAINAGE AREA: 2400 SQ KM

STANDARD DEVIATION: 168.276 M3/S

DISCHARGE (M3/S)



07G001 WAPITI RIVER ABOVE MISTANUSK CREEK
LOG PEARSON TYPE III DISTRIBUTION
PARAMETERS ESTIMATED BY MOMENTS

M A X I M U M D A I L Y M E A N F L O W S

07GE001 WAPITI RIVER NEAR GRANDE PRAIRIE

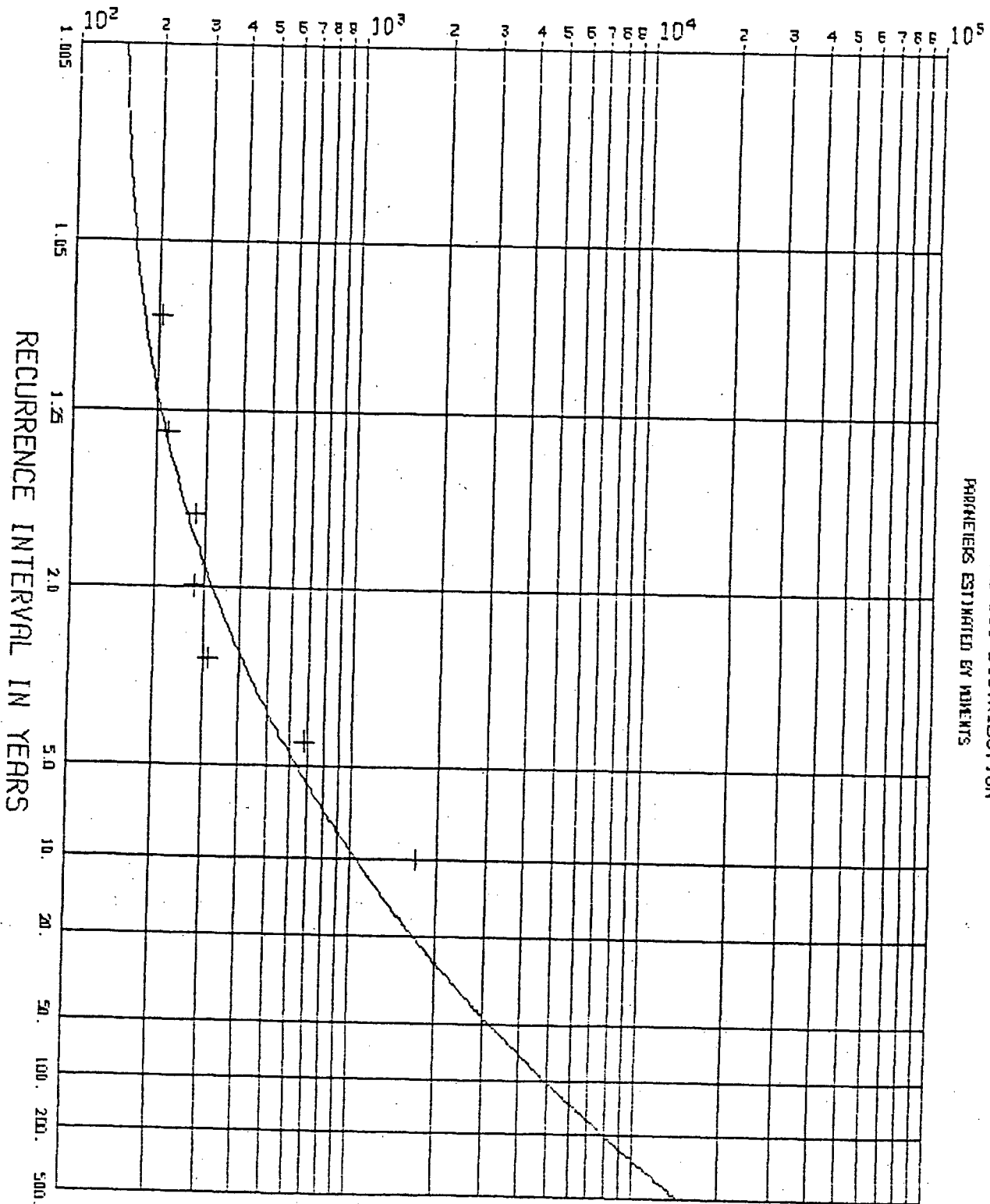
DATE	MAXIMUM DAILY FLOW IN M3/S	RANK	RECURRENCE INTERVAL IN YEARS	MAXIMUM DAILY FLOW IN M3/S	YEAR
MAY 24 1961	396.0	1	29.605	3680.0	1972
JUL 19 1962	1020.0	2	12.784	3630.0	1982
MAY 23 1963	535.0	3	8.152	2280.0	1965
JUN 29 1964	2090.0	4	5.984	2090.0	1964
JUL 9 1965	2280.0	5	4.727	1100.0	1980
MAY 11 1966	606.0	6	3.906	1020.0	1962
JUN 3 1967	960.0	7	3.328	968.0	1971
JUN 14 1968	807.0	8	2.899	960.0	1967
MAY 4 1969	402.0	9	2.568	926.0	1976
JUN 4 1970	467.0	10	2.305	889.0	1977
JUL 13 1971	968.0	11	2.091	807.0	1968
JUN 13 1972	3680.0	12	1.913	646.0	1978
MAY 19 1973	459.0	13	1.763	606.0	1966
APR 30 1974	595.0	14	1.635	595.0	1974
JUL 17 1975	561.0	15	1.524	585.0	1979
AUG 8 1976	926.0	16	1.428	561.0	1975
MAY 7 1977	889.0	17	1.342	535.0	1963
JUL 12 1978	646.0	18	1.267	505.0	1981
JUN 4 1979	585.0	19	1.199	467.0	1970
JUN 5 1980	1100.0	20	1.139	459.0	1973
MAY 27 1981	505.0	21	1.084	402.0	1969
JUL 16 1982	3630.0	22	1.034	396.0	1961

MEAN ANNUAL FLOOD: 1095.8 M3/S

DRAINAGE AREA: 11300 SQ KM

STANDARD DEVIATION: 960.209 M3/S

DISCHARGE (M3/S)



07GB002

KAKKA RIVER NEAR GRANDE PRAIRIE

LOG PEARSON TYPE III DISTRIBUTION

PARAMETERS ESTIMATED BY MOMENTS

M A X I M U M D A I L Y M E A N F L O W S

1

07GB002 KAKWA RIVER NEAR GRANDE PRAIRIE

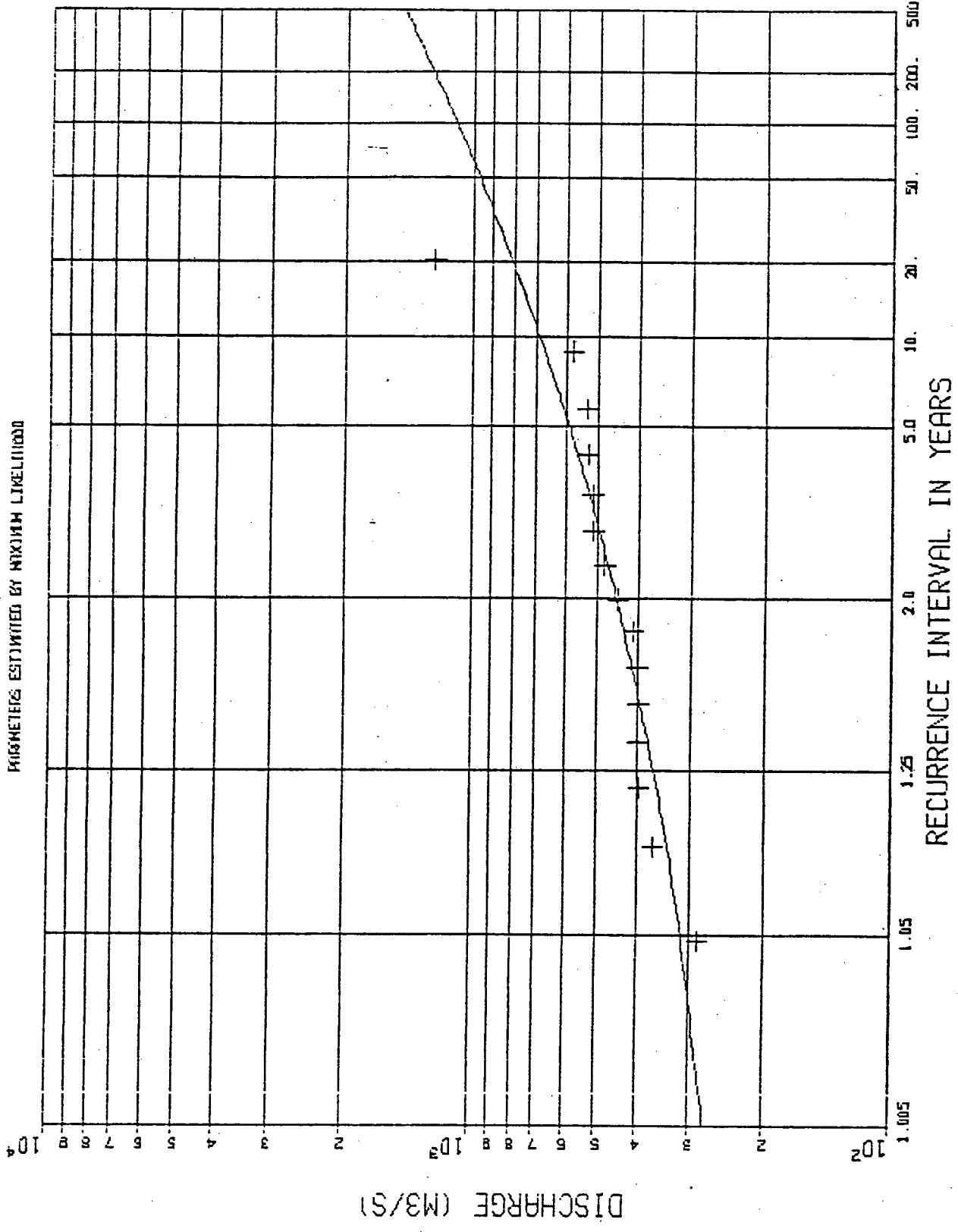
DATE	MAXIMUM DAILY FLOW IN M3/S	RANK	RECURRENCE INTERVAL IN YEARS	MAXIMUM DAILY FLOW IN M3/S	YEAR
JUN 27 1976	275.0	1	9.868	1680.0	1982
MAY 06 1977	309.0	2	4.261	683.0	1980
JUN 23 1978	220.0	3	2.717	309.0	1977
MAY 27 1979	275.0	4	1.995	275.0	1976
JUN 04 1980	683.0	5	1.576	275.0	1979
MAY 27 1981	205.0	6	1.302	220.0	1978
JUL 15 1982	1680.0	7	1.109	205.0	1981

MEAN ANNUAL FLOOD: 521.0 M3/S

DRAINAGE AREA: 3290 SQ KM

STANDARD DEVIATION: 536.339 M3/S

076A001 SMOKY RIVER ABOVE HELLS CREEK
LOG PEAKSON TYPE III DISTRIBUTION
PARAMETERS ESTIMATED BY MAXIMUM LIKELIHOOD



M A X I M U M D A I L Y M E A N F L O W S

07GA001 S M O K Y R I V E R A B O V E H E L L S C R E E K

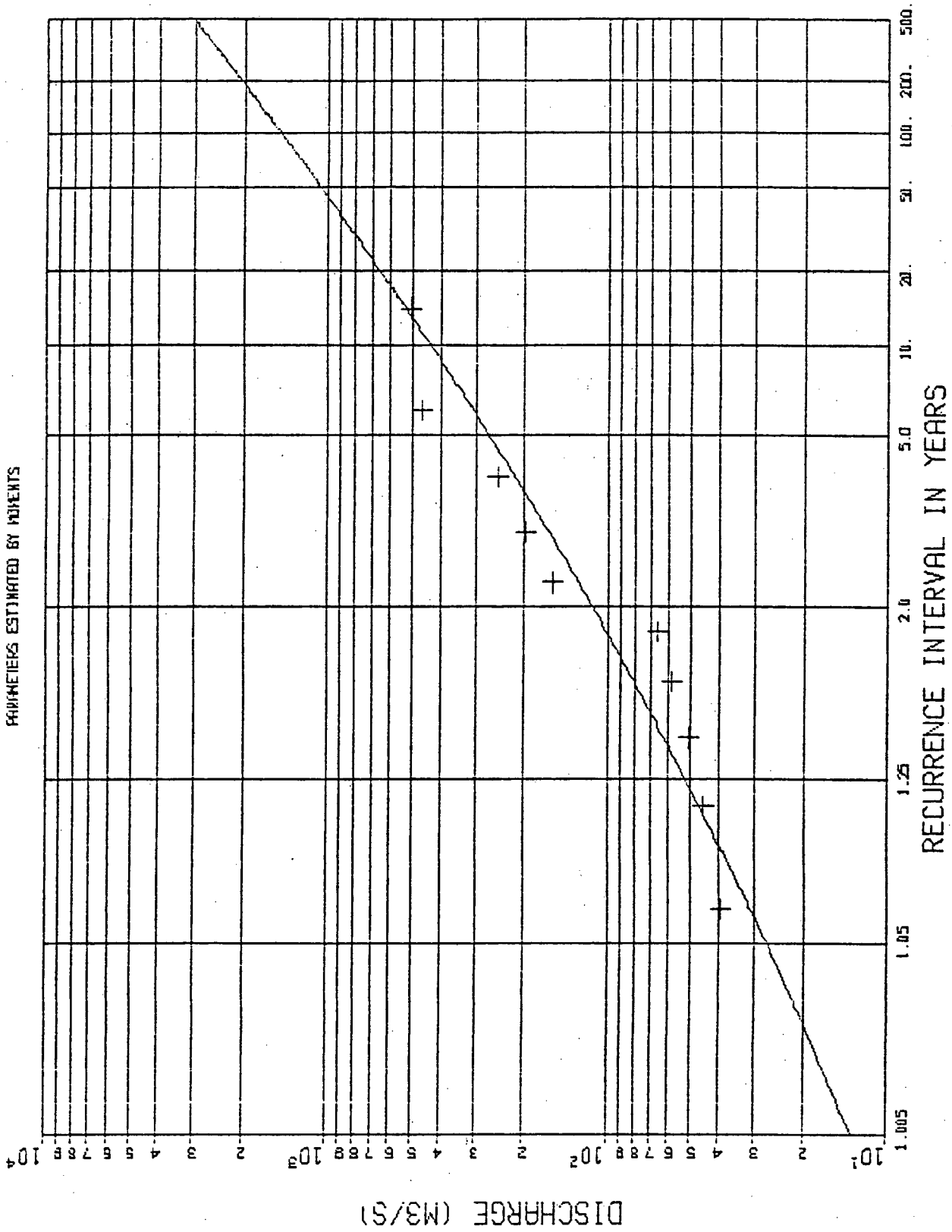
DATE	MAXIMUM DAILY FLOW IN M3/S	RANK	RECURRENCE INTERVAL IN YEARS	MAXIMUM DAILY FLOW IN M3/S	YEAR
MAY 21 1968	286.0	1	20.395	1230.0	1972
JUN 04 1969	399.0	2	8.807	575.0	1973
JUN 04 1970	527.0	3	5.616	532.0	1974
JUN 09 1971	515.0	4	4.122	527.0	1970
JUN 12 1972	1230.0	5	3.256	515.0	1971
JUN 24 1973	575.0	6	2.691	513.0	1982
JUN 16 1974	532.0	7	2.293	484.0	1977
JUL 06 1975	396.0	8	1.997	446.0	1980
JUL 01 1976	365.0	9	1.769	409.0	1979
JUN 08 1977	484.0	10	1.588	402.0	1981
JUL 11 1978	399.0	11	1.441	399.0	1969
JUN 05 1979	409.0	12	1.318	399.0	1978
JUN 04 1980	446.0	13	1.215	396.0	1975
MAY 26 1981	402.0	14	1.126	365.0	1976
JUN 15 1982	513.0	15	1.050	286.0	1968

MEAN ANNUAL FLOOD: 498.5 M3/S

DRAINAGE AREA: 3830 SQ KM

STANDARD DEVIATION: 216.619 M3/S

07GB001 CUTBANK RIVER NEAR GRANDE PRAIRIE
 LOG PEARSON TYPE III DISTRIBUTION
 PARAMETERS ESTIMATED BY MOMENTS



M A X I M U M D A I L Y M E A N F L O W S

1

07GB001 CUTBANK RIVER NEAR GRANDE PRAIRIE

DATE	MAXIMUM DAILY FLOW IN M3/S	RANK	RECURRENCE INTERVAL IN YEARS	MAXIMUM DAILY FLOW IN M3/S	YEAR
JUL 11 1971	198.0	1	13.816	506.0	1982
JUN 12 1972	464.0	2	5.966	464.0	1972
MAY 05 1973	39.2	3	3.804	249.0	1980
MAY 01 1974	51.3	4	2.793	198.0	1971
MAY 06 1977	159.0	5	2.206	159.0	1977
SEP 18 1978	45.0	6	1.823	66.0	1979
MAY 16 1979	66.0	7	1.553	59.2	1981
JUN 04 1980	249.0	8	1.353	51.3	1974
JUL 30 1981	59.2	9	1.199	45.0	1978
JUL 15 1982	506.0	10	1.076	39.2	1973

MEAN ANNUAL FLOOD: 183.7 M3/S

DRAINAGE AREA: 842 SQ KM

STANDARD DEVIATION: 174.651 M3/S

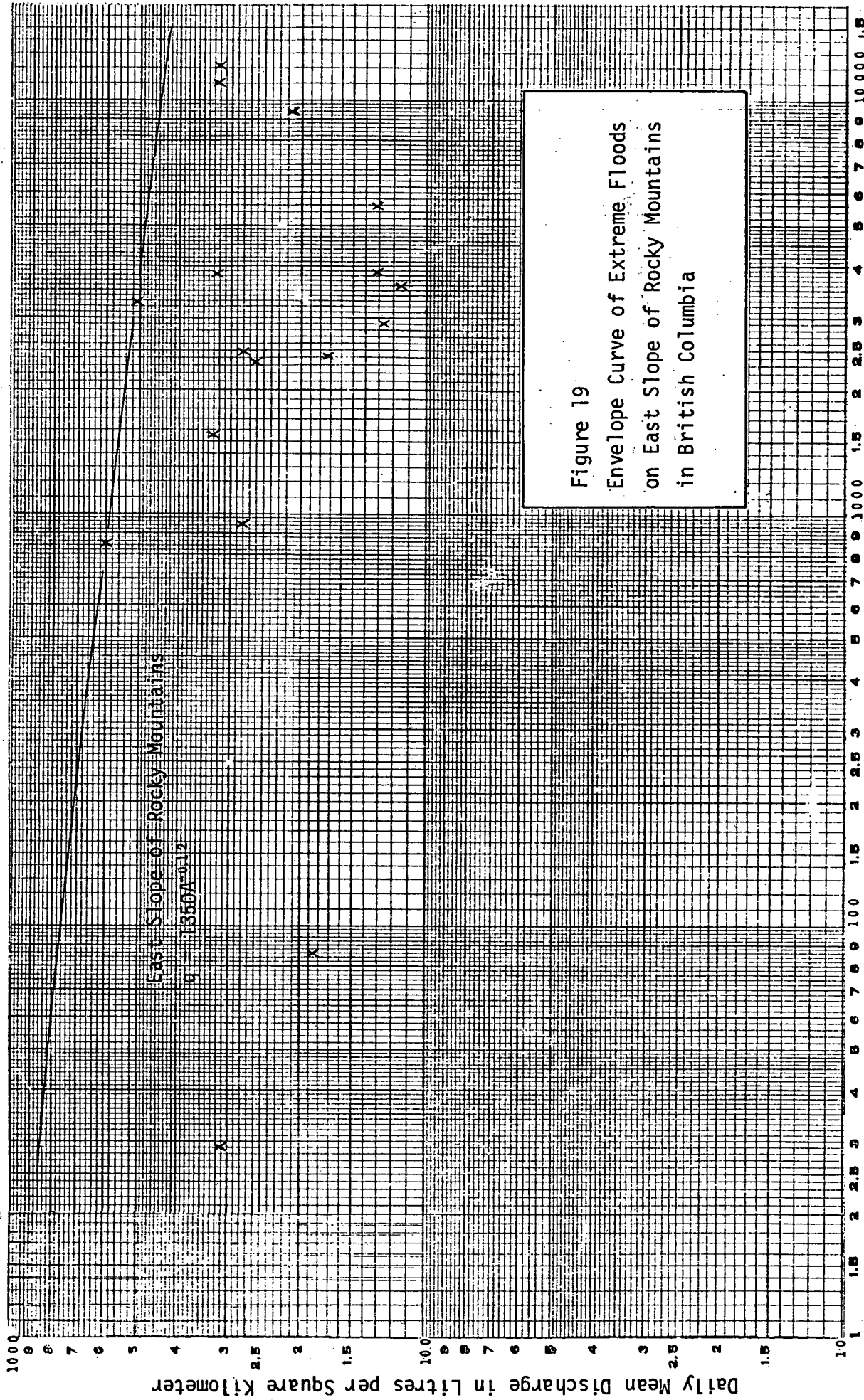


Figure 19
Envelope Curve of Extreme Floods
on East Slope of Rocky Mountains
in British Columbia

Drainage Area in Square Kilometers

Daily Mean Discharge in Litres per Square Kilometer