# Physiographic Parameters: Estimation and 

 Application to Hydrologic RegionalizationR.M. Leith

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Physiographic Parameters: Estimation and Application to Hydrologic Regionalization

R.M. Leith<br>Planning and Studies Section<br>Water Resources Branch<br>Vancouver, B.C.<br>May 1986

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## ABSTRACT

Mean basin physiographic quantities have been estimated from various sampling schemes. The major schemes are a 2 kilometre by 2 kilometre grid on a $1: 50,000$ scale topographic map and a 10 kilometre by 10 kilometre grid on a $1: 250,000$ map. Basin estimates have been compared and have been used as predictors (independent variables) in regional regression equations for mean annual flow. Results indicate that /f the $2 \times 2$ data provide estimates closer to fine sampling estimates of basin quantities than $10 \times 10$ data and that 2 X 2 data provide better predictors in that the regression equations developed using them have lower standard errors than do equations using 10 X 10 data.
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RESUME

# PHYSIOGRAPHIC PARAMETERS: ESTIMATION AND APPLICATION TO HYDROLOGIC REGIONALIZATION 

## 1. INTRODUCTION

### 1.1 Preface

The 1971 Shawinigan Engineering Company report "Hydrometric Planning Study for Western and Northern Canada" pointed to the need for a data bank and an analysis system in order to increase the usefulness of hydrometric data and to analyze the efficiency of data gathering. Such a data bank should contain not only hydrometric data but also other data that could be correlated with hydrometric data. These correlations would allow the hydrometric data to be extrapolated to ungauged basins, thus increasing its usefulness. The correlations and extrapolations as functions of the number of hydrometric stations used in their development provide accuracy measures for the assessment of hydrologic data gathering.

The data bank developed by Shawinigan Engineering Company contained physiographic parameters extracted using a 10 kilometre by 10 kilometre grid on $1: 250,000$ topographic maps. In this report this grid is referred to as the 10 X 10 grid. One of the recommendations of the Shawinigan report was that, in areas of complex hydrologic response, the analysis be expanded with physiographic data extracted. on a finer grid. In British Columbia, an area of complex, highly varied hydrology, a 2 kilometre by 2 kilometre grid ( 2 X 2 ) was chosen and the extraction of physiographic data from l:50,000 scale maps was undertaken. In this report this grid is referred to as the $2 \times 2$ grid and data extracted from this grid is referred to as $2 \times 2$ data. The
title for this report refers to estimation of physiographic quantities from 2 X 2 and 10 X 10 grids for southern British Columbia. The application of physiographic parameters to hydrologic regionalization refers to their use as predictors in regression equations for mean annual flow.

As of October 1985 the extraction of data for the province of British Columbia was complete south of Latitude 55 degrees north. To the north of 55 degrees the extraction is approximately $65 \%$ complete. However, only $80 \%$ of the extracted data has been transferred onto magnetic tape. Figure 1 indicates the availability of the physiographic data extracted from the $2 \times 2$ grid for the Province (Note: The figures and tables are in separate sections at the end of the report).

As the computer programs for storing and accessing the 2 X 2 data are now being tested, this time appears propitious to compare physiographic estimates made from the two grids and to examine the effects of these physiographic estimates upon hydrologic regionalization by the use of regression. An outline of objectives for this report showing the development of the two main objectives is as follows:

Objectives

1. Comparison of Basin 2. Examination of Contribution Quantities Estimated
From Various Spatial Sampling Schemes

Basin Quantities
Drainage Area
Elevation
Slope
Aspect
Area of Lakes
of Various Physiographic Quantities to Regional Hydrologic Equations

Regression Equations Developed for

2 X 2 Descriptors
10 X 10 Descriptors
10 X 10 Locators

Compare standard errors

### 1.2 Effects of Sampling Different Grids

In this study basin mean quantities are estimated from physiographic data extracted from different grids. The basin mean quantites compared are: elevation, slope, aspect and areas of lakes and swamps. Estimates of basin drainage area by various methods related to sampling grids are also compared. The question of accuracy of the estimates has considerable importance to both present and future hydrologic studies. The question is raised: does physiographic data extracted from a finer grid (higher sampling density) provide better estimates than those produced from a coarser grid? In this context, "better" implies producing regional regression equations with lower standard errors given a similar number of independent variables.

Physiographic data have been sampled on 10 X 10 and 2 X 2 grids. Other sampling schemes such as sampling at corner points and centre point of $2 \times 2$ squares and taking one sample elevation at the centre of a 6 km X 6 km square are used to fill in the range of sampling densities. In addition, elevations for the basin gauged by 08LG016 Pennask Creek near Quilchena have been sampled on a 100 metre by 100 metre grid. Several basins have had elevations sampled on a 1 kilometre by 1 kilometre grid. This allows comparison of mean physiographic quantities based on a range of spatial sampling.

### 1.3 Classification of Physiographic Parameters

In addition to the grid from which they are extracted, physiographic parameters can be further classified. In this report physiographic quantities such as elevation, land use, or slope are called descriptor parameters in that they describe a basin. Descriptor parameters can be classified as primary - those which are extracted directly from maps or derived - those which are derived from primary descriptors.

Elevation is a primary descriptor while slope and aspect are derived descriptors.

Physiographic quantities such as distance to the sea or barrier height in specified directions are termed locator parameters. As complete coverage is necessary to produce locator parameters, basin locator parameters are currently available only from the 10 X 10 grid. Complete coverage is the condition that there must be no missing data from the given square to the sea or other boundary of data coverage (such as the border between Canada and the United States).

### 1.4 Purposes of Study

This investigation has two purposes. The first is w comparison of basin quantities estimated from various sampling schemes. These schemes are primarily the 10 X 10 and $2 \times 2$ grids, but variations on the $2 \times 2$ grid and limited use of smaller resolution sampling schemes are employed in order to assess the variation in estimates of basin quantities in relation to the sampling densities. The second purpose is to asess the contribution of various physiographic quantities to regional hydrologic regression equations. Various physiographic quantities refer to physiographic quantities extracted from the 2 X 2 and 10 X 10 grids as well as basin descriptor and basin locator quantities. The effects of spatial resolution and type of physiographic quantity are investigated to try to provide answers to the following questions:
i) do 2 X 2 descriptors provide better equations than 10 X 10 descriptors?
ii) do derived basin descriptors provide improvement to equations?
iii) do basin locator variables make a significant contribution to regression equations?
iv) what are the results when parameters of different types and different resolutions are entered into regression studies?

## 2. PROCEDURE

### 2.1 Study Area

The study area chosen was southern British Columbia, which is densely gauged relative to other parts of the province, with approximately one station per 1200 square kilometres. Mean annual runoff ranges from less than 100 mm on the dry interior plateau to 4000 mm on the exposed west coast of Vancouver Island. In the study area, 71 stations operated by the Water Survey of Canada were selected. These stations record flow regimes classified as natural; that is, no significant man induced intervention to the flow regime is known. Drainage areas range from 20 square kilometres to 3000 square kilometres and are represented in heavy black silhouettes on Figure 2. Also shown on Figure 2 are three hydrologic regions which were identified in previous work by Leith (1983).

The three regions - coastal, wet interior and dry interior - were identified from a typicality-diversity plot, Wong and Liu (1975) and a bivariate plot of runoff versus mean basin response in Geostationary Operational Environmental Satellite (GOES) visible band. Further details on these procedures and results will be provided in a forthcoming publication. These methods produce objective delineation of regions which generally agree with experience. The boundaries of the regions are blurred, particularly from coastal to wet interior.

### 2.2 Basin Drainage Area

In this report basin drainage area is the area of the
surface drainage system and it is assumed that the surface and subsurface drainage systems coincide. Thus, the runoff recorded at hydrometric gauging stations is the response of both surface and subsurface drainage and can be ascribed to the surface drainage area. Note that the Water Survey of Canada number and name identifiers for a hydrometric station are used to identify a basin. Drainage areas estimated by three methods were compared.

The published Water Survey drainage areas are determined from a continuous delineation of drainage area on a topographic map. The basin boundary is delineated by joining the points of high ground around the basin with a continuous smooth curve as discussed in Linsley et al. (1949). Maps of both 1:50,000 and 1:250,000 scales are used for the delineation. The area enclosed by the curve is estimated by a program PAL (Point Area Length) which uses a digitized input of the basin boundary. Before the program became available, drainage areas were estimated by planimeter. The drainage area determined using this method will depend upon the skill of the person delineating the boundary and to some extent upon the relation between the size of the basin and the scale of the mapsheet. Basins with drainage areas between 200 and 400 square kilometres are usually considered to be estimated to within $2 \%$.

The $2 \times 2$ delineation of the basin depends upon the continuous basin outline and the 2 X 2 grid on l:50,000 mapsheets. Each square through which the continuous outline passes is included in the basin if the square contains $50 \%$ or more of the basin. Figure 3 provides an example of the continous curve in heavy black and the resulting 2 X 2 approximation to the boundary appears as the polygon figure. Figure 3 also provides an example of the coding of $2 \times 2$ physiographic data on a mapsheet. Kreuder (1979) provides a description of the development of the $2 \times 2$ data bank.

The drainage areas extracted from the $10 \times 10$ square grid were accepted as provided by Water Resources Branch in Ottawa.

### 2.3 Mean Basin Elevation

Elevations were taken at the centre point of each $2 \times 2$ square. To examine the effects of more frequently sampled data, elevations were extracted at the corner points of each square. This sampling scheme produced a slightly higher sampling density than the centre point sampling scheme, the increase depending upon the number of basin boundary squares. Basins were initially chosen to cover a range of drainage basin areas from 20 square kilometres to 600 square kilometres and a range of steepness from relatively flat to very steep basins. Relatively flat basins had a range of elevations less than 500 feet while the relatively steep basins had elevation ranges of 5000 feet. Several basins showed large differences in mean basin elevation estimated from 10 X 10 and $2 \times 2$ data. For these basins elevation was extracted on a $1 \mathrm{~km} X 1 \mathrm{~km}$ grid, the elevation being extracted at the centre point, thus providing four samples of elevation in each 2 km X 2 km square or a sampling density of 1 sample per square kilometre. For the basin gauged by station 08LG016 Pennask Creek near Quilchena, the mean basin elevation was estimated from elevations taken to produce a digital elevation model (DEM). Elevations were sampled on a 100 metre by 100 metre grid from stereoscopic air photos. Accuracy of elevations estimated using this technique depends upon ground reference points; for the Pennask basin which is relatively open (not heavily treed). each elevation is expected to be be accurate to $\pm 5 \%$. The digital elevation model work was done by Pacific Survey (Spar Aerospace) on contract to the Water Survey of Canada.

10 km by 10 km elevation is established by sampling the elevation at each corner point and the centre point of the

10 km by 10 km square. The usual contour interval on 1:250,000 maps in British Columbia is 500 feet, while on the 1:50,000 maps the contour interval is usually 100 feet although it can be less. These differences in scale and resolution should be kept in mind in the comparison.

### 2.4 Mean Basin Slope

Previous studies on hydrologic regionalization (Leith, 1978) have shown that basin slope is a statistically important parameter, particularly for low flows. Ragan (1985) has pointed to the importance of basin slope to the timing of hydrologic response and to the sensitivity of slope to the sampled elevations which go into its estimation.

On a square by square basis slope is estimated in two ways from the 2 X 2 data: first from contour crossing data, reference National Handbook of Recommended Methods for Water Data Acquisition, and second from the largest gradient between the centre point elevation of the square under consideration and the nearest neighbour elevations. For the first method the slope (s) for a square is found from

$$
s=\frac{N h \sec (c c)}{1}
$$

where $N$ is the number of contour crossings both east-west and north-south, $h$ is the contour interval, for example, loo feet, cc is the angle between the contour lines on the map and the grid line, 1 is the length of the grid, in this: case 4 kilometres.

Linsley et al. (page 250) suggest that any attempt to determine an average secant is at best uncertain, but Horton's value of 1.571 has been accepted for the estimation of slope.

The second method for slope estimation from $2 \times 2$ data is illustrated in Figure 4; the data are elevations in metres taken for 08 NH 084 Arrow Creek near Erickson. The slope for a square is found by scanning for the greatest change in elevation from the centre point of the square to the centre point of the nearest neighbours. For an interior square there are 8 nearest neighbours. For a boundary square only squares within the basin are scanned. The slope is the greatest change in elevation divided by 2 km .

The procedure for estimating slope for each 10 X 10 square is shown in Figure 4. The slope in the north-south direction is estimated from the difference between the elevation in the northwest corner and the elevation in the southwest corner divided by 10 km . The slope in east-west direction is estimated from the difference between the elevation in the southwest corner and the elevation in the southeast corner divided by 10 km . The slope for the square is found by taking the square root of the sum of the slope squared in the north-south direction and the slope squared in the east-west direction.

An estimate of slope is also provided by the Digital Elevation Model for Pennask Creek basin. This estimate is based upon greatest rate of change from an observation elevation to its nearest neighbours.

### 2.5 Relative Area of Lakes and Swamps

For the $2 \times 2$ grid, the area of lakes and swamps is estimated by overlaying each square with a transparency having a random array of dots. The number of dots covering lakes and swamps were counted, no distinction being made between lakes and swamps. With reference to Figure 3, the dot count is the number in the lower left quadrant. The area in square inches is the number of dots divided by 100. For mapsheets with a scale of $1: 50,000$, the area of lakes
and swamps in square miles is found by dividing the dot count by 100 and multiplying by 0.62 . The derivation of the 0.62 factor is developed in Appendix 2.

The dot overlay claims a precision of $97 \%$ if a map area of at least 12 square inches is estimated. The error for smaller areas could be large, small areas of lakes and swamps could be missed. For dot counts of the order of 25 the error should be less than $15 \%$. This error estimate depends upon the accuracy of the map.

### 2.6 Basin Orientation or Aspect

The measure of basin orientation or basin aspect is provided by taking the most frequently occurring orientation for the 2 km by 2 km squares. The orientation of each square is taken to be the direction of greatest rate of change of elevation from the slope of the square (see Figure 4 previous section for an example). If the gradient from the centre point of a square is positive, i.e. the neighbouring square is lower, then the aspect is taken to be the direction to the neighbouring square. If the gradient is negative, i.e. the neighbouring square is higher, then the aspect is taken as gradient direction plus 180. The directions for the aspect of a square are quantized to every 45 degrees. An aspect of zero degrees is north. An example of the aspects of the squares of the Arrow Creek near Erickson basin together with their distribution is shown in Figure 5. From this distribution the basin was assigned an aspect of 180 degrees, i.e. south.

For each 10 km by 10 km square a plane is fitted through the elevations of the northwest, northeast and southwest corner points of the square, then the projection of the orienation of this fitted plane on the horizontal plane is found and the azimuth is taken to be the angle clockwise from north to the projection. The basin aspect is the mean value for the
squares belonging to the basin.

### 2.7 Annual Runoff

The response (dependent) variable for the regression studies was mean runoff over the 48 month period from November 1978 to October 1982. In this case, mean is determined by dividing by four, as if there were four years of data. The runoff in millimetres (mm) is obtained by dividing the flow volume in cubic decametres by the drainage areas in square kilometres.

In order to see if the results were applicable to longer term runoff, mean runoff from 1950 or when the station opened until 1984 was computed and compared with the 48 month mean runoff.

The use of the 48 month mean runoff makes the results capable of comparison with previous studies (Leith, 1983).

Only 66 of the 71 basins had sufficient flow records to be used in the regression portion of the study. The other five were missing too much record, that is, $20 \%$ or more, or were missing high flow records.

### 2.8 Regression Equations

Regression equations for runoff (response) were developed in terms of physiographic parameters (predictors) using the Triangular Regression Package (TRP) at the University of British Columbia. The equations were developed using backward elimination. Equation building calibration-validation exercises were undertaken in the wet interior region in order to assess the variability in standard error and to get an indication of importance of various physiographic parameters in terms of the frequency with which they occurred.

## 3. RESULTS

### 3.1 Drainage Area

For drainage basins with published drainage areas greater than 40 square kilometres the $2 \times 2$ drainage area estimates agree with published drainage areas to within $4 \%$. The 10 X 10 drainage area estimates are within $4 \%$ of published areas for basins greater than 100 square kilometres. The results are summarized in Table 1 . In this table the mean absolute difference refers to the mean value of the absolute differences, that is, there is no consideration given to whether the difference is positive or negative. For example, for $2 \times 2$ estimates of drainage areas for basins with areas between 41 and 100 square kilometres, the mean of the signed differences is $-0.1 \%$, while the mean of the absolute differences is $4 \%$. The mean of the absolute differences is assumed to be useful for developing an envelope curve.

With reference to Table 1 , for the six basins with published drainage areas less than 40 square kilometres, five have had their drainage areas overestimated by 2 X 2 . Overestimation yields a positive difference between published and estimated drainage areas. The mean signed difference is $9 \%$. This indicates a significant overestimation of drainage area. For other drainage area ranges, the number of overestimates and underestimates are nearly equal and there is no indication of bias.

A noticeably large difference of $28 \%$ between the $2 \times 2$ estimate of drainage area and the published area is provided by station 08NH130 Fry Creek below Carney Creek. For this basin the basin boundary input to PAL has been questioned and a larger basin thought to be more realistic was used for the 2 X 2 delineation.

### 3.2 Mean Basin Elevation

The differences in mean basin elevation estimated from 2 X 2 and 10 X 10 data are summarized in Table 2 and plotted in Figure 6. The percentage differences are taken relative to the 2 X 2 estimate. It is interesting to observe that for basins with drainage areas between 40 and 1000 square kilometres the mean unsigned difference is relatively constant - between 10 and $13 \%$. This mean is largely controlled by a small number of large differences. Even for the San Juan River basin (08HA010), a relatively large area of 580 square kilometres, the difference is $25 \%$.

In Figure 6 a positive difference indicates that the 2 X 2 estimate of mean basin elevation is greater than that for the $10 \times 10$. For $69 \%$ ( 49 of 71 ) of the basins the $2 \times 2$ data produce higher estimates of mean basin elevation. While the largest differences and greatest spread are observed for small basins, the size of differences and spread has not reduced much for basins of 200 to 300 square kilometres. Differences, both positive and negative, greater than $20 \%$ occur in 15 out of 62 basins with drainage areas of less than 500 square kilometres and for 1 out of 9 basins with drainage areas greater than 500 square kilometres. Figure 7 shows the plot of the difference between the estimates of 2 X 2 and 10 X 10 mean basin elevations against the range in elevations in the basins. The range in elevations is the difference between the maximum and minimum $2 \times 2$ centre point elevations. There may be a suggestion of narrowing in the difference in mean basin elevation with range of elevation but there is no strong relation.

The sixteen basins with largest differences in mean basin elevation estimates are listed in Table 3. The differences are not surprising for basins smaller than 200 square kilometres but are surprising for larger basins. It
is worth noting that all of the sixteen basins except for $08 N N 016$ Sutherland Creek near Fife are in the Vancouver Island and Coastal Mountain area of British Columbia.

Table 4 provides a summary of the consistency of mean basin elevation estimates. In this study the best estimate for the mean basin elevation is taken to be that provided by 1 km X l km data. In general, the 2 X 2 centre point estimate of mean basin elevation is found to be closer to the 1 X l estimate that is the $10 \times 10$ estimate and is therefore termed more consistent.

The $2 \times 2$ centre point and $2 \times 2$ four corner point estimates are generally close, that is within one standard error of the mean. For example, for Shawnigan basin (OBHA033), the standard error for the $2 \times 2$ centre point is 61 feet based on 22 observations. The four corner point estimates of mean basin elevation are based on more observations than the centre point estimates, for the Koksilah basin (OBHA003) 80 four corner point observations, to 55 centre point observations. In this case the 80 observations are independent, that is allowance has been made for a corner point belonging to more than one square. For one basin, gauged by 08HA010 San Juan near Port Renfrew, the 10 km by 10 km sampling of elevation was checked by taking six squares which partially covered the basin. Mean basin elevation was estimated to be 1600 feet. The 10 X 10 estimate supplied by Ottawa was 1366 feet.

### 3.3 Mean Basin Slope

The mean basin slope has been estimated for each basin by the three methods detailed in Section 3.2.4. Figure 8 shows that estimates made from $2 \times 2$ elevation data are generally about $1 / 2$ of the estimates made from contour crossings. The distributions about the 1:2 line is amazingly narrow for physiographic quantities, the exception
being basin 22, station 08HA034 Craigflower Creek below Burnside Road (drainage area 13.5 square kilometres).

The contour crossing method was assumed to produce more reliable estimates of slope, although estimates are probably going to be high as some contours cross a grid line more than once. The assumption was supported by estimates of basin slope for station 08LG016 Pennask Creek above Quilchena. These estimates are summarized in Table 5 and show that the $2 \times 2$ contour crossings estimate was the closest to the DEM estimate.

Table 6 lists basins with large differences between 2 X 2 contour crossing estimates and 10 X 10 estimates. The last column of the table suggests factors which may have produced the differences. Most of the basins have shown large differences in either drainage area or mean basin elevation beween $2 \times 2$ and 10 X 10 estimates. Basin 08MG005 Lillooet River at Pemberton contains extensive glaciers which may affect the estimation of elevations and contour crossings.

### 3.4 Relative Area of Lakes and Swamps

In general, the area of lakes and swamps estimated for basins on 2 X 2 and 10 X 10 grids agreed. The 10 x 10 estimates seemed quantized in units of 2.0 for lakes and 1.0 for swamps, so basins with relative areas less than $1 \%$ as estimated from 2 X 2 registered as $0 \%$ from 10 X 10 data. Several smaller basins showed large differences in estimates made from $2 \times 2$ and l0X 10 data. From Table $708 N N 023$ Burrel Creek showed a large discrepancy which was possibly brought about by the inclusion of the Lower Arrow Lake in the basin on the 10 X 10 grid.

### 3.5 Mean Basin Aspect

The mean basin aspect might better be termed the mode for the 2 X 2 . In general, the orientation of the basin estimated from basin aspect $2 \times 2$ agrees with the $10 \times 10$ estimates given that there is a 45 degree quantization in the $2 \times 2$ (Table 8). Notable exceptions are small basins (less than 100 square kilometres) and several basins with large differences (greater than $10 \%$ ) in estimated drainage area from between $10 \times 10$ or $2 \times 2$ and the published source. The basins are identified in Table 8.

The discrepancies may be due to inclusion in the basin of 10 X 10 squares which do not cover a large percentage of the basin and those not representative of the basin. The difference observed for station 08MG005 Lillooet River near Pemberton may be associated with the difference in slope (see Section 3.3).

### 3.6 Annual Runoff

Figure 9 shows the close association between the 48 month mean runoff and the long term mean annual runoff. The majority of stations are within $10 \%$, suggesting that the regression results of this study are not limited to the period November 1978 to October 1982 but have wider applicability.
3.7 Contribution of Various Physiographic Quantities to Regional Equations

Tables 9, 10, and 11 summarize the regression equations for the Wet Interior, Dry Interior and Coastal Regions. These regions are outlined in Figure 2. Assuming that a reduction in standard error of estimate provides an improved equation in terms of lower errors in prediction, then in Table 9 the equation developed from 2 X 2 basin descriptor parameters is superior to the equation from 10 X 10
descriptors. The introduction of $10 \times 10$ locators produced an equation with lower standard error than that developed from basin descriptors either 2 X 2 or 10 X 10.

Figure 10 shows the results of an equation building calibration-validation exercise for Wet Interior Region basins. The horizontal axis shows the number of basins used in developing the calibration equation. These basins were random samples. The equation developed from each random sample was then applied to the remaining basins. With calibration samples of twenty basins the range of calibration errors has narrowed from that for 10 basins. This provides an indication of the range of calibration error to be expected for 26 basins. Using the equation for prediction in the Wet Interior Region will probably involve larger errors than those suggested by calibration standard error. The magnitude of this error is indicated by the range of validation errors for 20 basins. All the above results are for regression equations developed by ordinary least squares.

In the buildup procedure the occurrence in various physiographic parameters in the equations was noted and is summarized in Table 12. The basin slope estimated from contour crossings on the 2 X 2 grid (SLPCC2) occurred six times in the nine equations developed.

From Table 10, in the Dry Interior Region, basin descriptors from either grid were statistically insignificant. 2 X 2 descriptors combined with basin locators produced an equation with a slightly. lower standard error than that for the equation developed from locator parameters.

From Table 11 , in the coastal region $2 \times 2$ descriptors were superior to $10 \times 10$ descriptors and basin locators were required to produce the best equation.

### 4.0 CONCLUSIONS AND RECOMMENDATIONS

With reference to Section 1.4 Purposes of Study, the comparison of sampling on estimating physiogaphic parameters indicates that, in general, $2 \times 2$ data do provide more reliable estimates of basin properties than do 10 X 10 data. "Reliable" here refers to the evidence that the 2 X 2 estimates are closer to those made with more densely sampled data than are estimates made with 10 X l0. This consistency is particularly evident in basins between 40 and 100 square kilometres and for derived parameters such as basin slope. Drainage area estimates made from the 2 X 2 grid are generally closer to the published values, indicating that 10 X 10 estimates of the basins may be either including or excluding significant areas of the basin, hence producing unrepresentative basin average estimates. The 2 X 2 data also appears to provide more realistic estimates of area of lakes and swamps which, considering the larger scale of the maps from which $2 \times 2$ are extracted, is as would be expected.

As to the comparison of effects on regression equations of various physiographic parameters, the results do appear to depend upon hydrologic region. Except for the Dry Interior Region, $2 \times 2$ descriptors explain more of the variance, i.e. they provide better predictors than do 10 X 10 descriptors. 10 X 10 locators improve regression equations in terms of reducing the standard errors. In the Dry Interior Region only the 10 X 10 basin locators were significant. These results clearly show the value of the 2 X 2 descriptors and the 10 X 10 locators to hydrologic regionalization by regression equations.

Recommendations arising from this study are:

1) That the importance of slope be further investigated, particularly with emphasis on the estimation of slope by
contour crossings. It may be that contour crossings do not have to be extracted from the unprocessed mapsheets since estimates from $2 \times 2$ elevations are very close to 1/2 estimates by contour crossing.
2) That the development of location parameters from $2 \times 2$ data be examined since location parameters do make a contribution to regional equations.
3) That Digital Elevation Models be prepared for several basins, particularly the ones with large differences between $2 \times 2$ and $10 \times 10$ estimates. This would allow a more extensive examination of the assumption that closely sampled data provide better basin estimates.
4) That air photos used to develop the DEMs in Recommmendation 3 be used to estimate areas of lakes and swamps and forest cover for the basins.
5) That the addition of land cover data such as forest be considered for the 2 X 2 data bank.

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Figure 2 Location of Basins (black silhouettes) and Delineation of Hydrologic Regions

Basins are indicated in black.


Dry Interior Region
$\because \because \because$ Vancouver Island and
Lower Mainland
Wet Interior Region

- remaining study region

Figure 3 Shawnigan Creek Basin on Vancouver Island and Sample of Coded Grid Square


Upper Right
Upper Left
Lower Left

Lower Right

Elevation in Feet
Number of Contour Crossings
Number of Dots for Area of Lakes and Swamps

Number of Divider Intervals for Measuring Channel Length

Figure 4 Estimation of Basin Slope From $2 \times 2$ Elevations Arrow Creek near Erickson

21031829
16461875
$1676 \quad 1737$
176815241341
160011891692
143310971737
15099751798
1463975

Slope in upper corner is $\frac{(2103-1646)}{2 \mathrm{~km}}=229 \mathrm{~m} / \mathrm{km}$.
The direction is toward the square with elevation 1646 m , i.e., south or $180^{\circ}$.

From $10 \times 10$ Elevation Data

| 10 km |  |  |
| :---: | :---: | :---: |
|  | $\int \begin{aligned} & \text { Northwest Corner } \\ & 1070\end{aligned}$ |  |
|  | Southwest Corner $1500$ | Southeast Corner $1200$ |
|  | 10 km |  |
|  | Slope $=\sqrt{\left(\frac{1500-1070}{10}\right)^{2}+\left(\frac{1500-1200}{10}\right)^{2}}$ |  |
| $=52 \mathrm{~m} / \mathrm{km}$ |  |  |

## Figure 5 Square by Square Aspect Values for Arrow Creek

 near Erickson BasinPhysiographic parameter number 8 Aspect (degrees clockwise from North)

| 180 | 225 |  |  |  |
| ---: | ---: | ---: | ---: | ---: |
| 180 | 270 |  |  |  |
|  | 135 | 135 |  |  |
|  | 90 | 135 | 180 |  |
|  |  | 135 | 180 | 225 |
|  |  | 135 | 180 | 225 |
|  |  | 90 | 180 | 270 |
|  |  | 45 | 0 |  |



Distribution of aspects found in the basin

Direction Tally of $2 \mathrm{~km} \times 2 \mathrm{~km}$ squares

| N | $X$ |
| ---: | :--- |
| $N E$ | $X$ |
| $E$ | $X X$ |
| $S E$ | $X X X X X$ |
| $S$ | $X X X X X X$ |
| $S W$ | $X X X$ |
| $W$ | $X X$ |
| $N W$ |  |

Mode of physiographic parameters over basin




Figure 10 Standard Errors versus Number of Basins in Wet Interior Equation Building

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Table 1 Summary of Comparison of Drainage Areas

| Range of Published Drainage Areas Square Kilometres | Number of Basins | $2 \mathrm{~km} \dot{\mathrm{X}} 2 \mathrm{~km}$ |  | $10 \mathrm{~km} \times 10 \mathrm{~km}$ |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | Mean Absolute Difference | Range of Differences | Mean Absolute Difference | Range of Differences |
| 0-40 | 6 | 13\% | -11 to 28 |  |  |
| 41-100 | 20 | 4\% | -8 to 9 | 18\% | -62 to 20 |
| 101-200 | 13 | 5\% | -4 to 12 | 3\% | -4 to 8 |
| 201-400 | 19 | 2\% | -7 to 4 | 8\% | -5 to 46 |
| 401-1000 | 10 | 4\% | -4 to 28 | 8\% | -3 to 18 |
| Over 1000 | 3 | 3\% | -5 to 2 | 3\% | -4 to -1 |

Table 2 Summary of Differences in Mean Basin Elevation Estimates

| Drainage Area <br> Square <br> Kilometres | Number of <br> Basins | Mean Unsigned <br> Difference <br> 2 km by 2 km <br> 10 km by 10 km | Range of <br> Differences |
| :---: | :---: | :---: | :---: |
| $0-40$ | 6 | $41 \%$ | $14 \%-68 \%$ |
| $40-100$ | 20 | $13 \%$ | $1 \%-48 \%$ |
| $101-200$ | 13 | $12 \%$ | $0 \%-37 \%$ |
| $201-400$ | 19 | $13 \%$ | $3 \%-47 \%$ |
| $401-1000$ | 10 | $10 \%$ | $0 \%-25 \%$ |
| 0 ver 1000 | 3 | $6 \%$ | $1 \%-8 \%$ |

Table 3 Basins with Large Differences between Estimates of Mean Basin Elevation from $2 \times 2$ and $10 \times 10$ Data

Water Survey Station Number

Basin Name
Drainage Area Difference in ( $\mathrm{km}^{2}$ ) Percent
$08 H A 016 \quad$ Bings Creek near the Mouth 15.542

08HA001
Chemainus River near Westholme 355 31

08 HA034 Craigflower Creek below Burnside Road, Vict. 13.534
08 MH076 Kanaka Creek near Webster Corners 47.7
08HB063 Kitsucksus Creek above Cherry Creek 68
08 KA003 Koksilah River at Cowichan Station 20947
08HB029 Little Qualicum River near Qualicum Beach 23739
08NN021 Moody Creek near Christina 46
08MH105 Nicomekl River below Murray Creek 64.5
08 NH058 Norrish Creek near Dewdney 11722
08HA010 San Juan River near Port Renfrew 580
08HA033 Shawnigan Creek near Mill Bay 91.9
08MH056 Slesse Creek near Vedder Crossing 162
08NN016 Sutherland Creek near Fife 88.1 22
08HB024 Tsable River near Fanny Bay $113 \quad 37$
08MG021 Twenty-One Mile Creek at 670 m Contour 28.26

Table 4 Summary of Mean Basin Elevation, Estimates

| Water Survey Station Number | Mean Basin Elevation Estimate in Feet |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  | Drainage Area Sq. Km. | $1 \times 1$ | $\begin{gathered} 2 \times 2 \\ 4 \text { corner } \\ \text { points } \\ \text { (ft.) } \end{gathered}$ | $\begin{aligned} & 2 \mathrm{X} 2 \\ & \text { centre } \\ & \text { point } \end{aligned}$ | $10 \times 10$ |
| OBHA016 | 15.5 | 828 | 797 | 825 | 480 |
| 08 HA 001 | 355 |  | 2013 | 1995 | 1371 |
| 08HA034 | 13.5 | 316 | 344 | 313 | 435 |
| 08MH076 | 47.7 | 1069 | 1014 | 1091 | 563 |
| 08 HB 063 | 15.5 | 600 | 697 | 475 | 795 |
| 08 HA 003 | 209 | 1516 | 1504 | 1544 | 2272 |
| 08HB029 | 237 | 1907 | 1995 | 1885 | 1150 |
| 08 MHI 05 | 64.5 | 189 | 185 | 187 | 128 |
| 08 MH 058 | 117 | na | 2494 | 2589 | 2030 |
| 08 HA 010 | 580 | na | 1748 | 1826 | 1366 |
| 08 HA 033 | 91.9 | 649 | 684 | 641 | 832 |
| 08MH056 | 162 | na | 4292 | 4257 | 5439 |
| 08 HB 024 | 113 | 2427 | 2387 | 2491 | 1629 |

## Table 5 Basin Slope Estimates for 08LG016 Pennask Creek above Quilchena

| Method | Slope (m/km) |
| :---: | :---: |
| $2 \mathrm{~km} \times 2 \mathrm{~km}$ Contour Crossings | 126 |
| $2 \mathrm{~km} \times 2 \mathrm{~km}$ Centre Point Elevations (max. gradient) | 51 |
| $2 \mathrm{~km} \times 2 \mathrm{~km}$ Four Corner Point Elevation (max. gradient) | 80 |
| $10 \mathrm{~km} \times 10 \mathrm{~km}$ | 22 |
| Digital Elevation Model 100 m x 100 m grid (max. gradient) | 186 |


| Basin otting lumber | Table 6 Basins with Large Differences between the Estimates of Mean Basin Slope ( $2 \mathrm{~km} \times 2 \mathrm{~km}$ contour crossings and $10 \mathrm{~km} \times 10 \mathrm{~km}$ contour crossings) |  |  |  | . |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  | Water Survey Station Number | Basin Name Dr | Drainage Area ( $\mathrm{km}^{2}$ ) | Difference in Mean Basin Elevation $2 \times 210 \times 10$ | Difference in Drainage Area |
| 4 | O8NE008 | Beaton Creek near Beaton | 99.5 | 7\% | 60\% |
| 5 | 08 NMO 035 | Bellevue Creek near Okanagan Mission | 73.3 | 7\% | 23\% |
| 13 | 08GA060 | Chapman Creek above Sechelt Diversion | 64.5 | 5\% | NA |
| 15 | 08MH103 | Chilliwack River above Slesse Creek | 645 | 15\% | 50\% USA |
| 16 | 08 MH 016 | Chilliwack River at Outlet of Chilliwack Lake | 329 | 18\% | part of basin USA |
| 17 | 08 MM142 | Coldstream River above Municipal Intake | 58.5 | 13\% | NA |
| 18 | 08LG048 | Coldwater River near Brookmere | 316 | 3\% | 2\% |
| 23 | 08 NE087 | Deer Creek at Deer Park | 80.5 | 18\% | 4\% |
| 29 | 08NE120 | Hill Creek at Mill Road Bridge | 28.2 | 14\% | NA |
| 38 | 08MG005 | Lillooet River near Pemberton | 2160 | -1\% | glaciers |
| 51 | 08HA020 | San Juan River near Port Renfrew | 580 | 25\% | NA |
| 54 | 08NL070 | Similkameen River above Goodfellow Creek | k 407 | 6\% | USA |
| 55 | 08MH056 | Slesse Creek near Vedder Crossing | 162 | 28\% | USA |
| 57 | 08NNO19 | Trapping Creek near the Mouth | 144 | 4\% | 4\% |
| 62 | 08HC002 | Ucona River at the Mouth | 185 | 18\% | $3 \%$ |

Table 7 Relative Area of Lakes and Swamps Expressed as Percent of Basin Area

| Identifier on plot (Fig. 8) | Water <br> Survey <br> Number | Basin Name | $2 \times 2$ | $\begin{array}{r} 10 \\ \text { Lakes } \end{array}$ | $10$ <br> Swamps |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 1 | $08 \mathrm{NH084}$ | Arrow Creek near Erickson | 0.0 | 0.0 | 0.0 |
| 2 | 08NL004 | Ashnola River near Keremeos | 0.475 | 0.0 | 0.0 |
| 3 | 08 NE077 | Barnes Creek near Needles | 0.160 | 0.0 | 0.0 |
| 4 | 08NE008 | Beaton Creek near Beaton | 1.63 | 2.0 | 0.0 |
| 5 | 08NM035 | Bellevue Creek near Okanagan Mission | 0.0 | 2.0 0.0 | 0.0 |
| 6 | 08NE039 | Big Sheep Creek near Rossland | 0.023 | 0.0 | 0.0 |
| 7 | 08HAO16 | Bings Creek near the Mouth | 0.0 | 2.0 | 0.0 |
| 8 | 08LE094 | Bolean Creek near the Mouth | 1.174 | 1.0 | 0.0 |
| 10 | 08NH032 | Boundary Creek near Porthill | 0.308 | 0.0 | 0.0 |
| 10 | 08NM133 | Bull Creek near Crump | 0.269 | 0.0 | 0.0 |
| 11 | 08NN023 | Burrell Creek near Gloucester Creek | 0.161 | 6.0 | 0.0 |
| 12 | 08NH131 | Carney Creek below Pambrun Creek | 0.428 | 1.0 | 0.0 |
| 13 | 08GA060 | Chapman Creek above Sechelt Diversion | 1.512 | 1.0 | 0.0 |
| 14 | 08HA001 | Chemainus River near Westholme | 0.241 | 1.0 | 0.0 |
| 15 | 08MH103 | Chilliwack River above Slesse Creek | 2.164 | 2.0 | 0.0 |
| 16 | $08 \mathrm{MH016}$ | Chilliwack River at Outlet of Chilliwack Lake | 3.886 | 2.0 | 0.0 |
| 17 | 08NM142 | Coldstream Creek above Municipal Intake | 0.317 | 0.0 | 0.0 |
| 18 | 08LG048 | Coldwater River near Brookmere | 0.459 | 0.0 | 0.0 |
| 19 | 08MF003 | Coquihalla River near Hope | 0.154 | 0.0 | 0.0 |
| 20 | 08MF062 | Coquihalla River below Needle Creek | 0.543 | 0.0 | 0.0 |
| 21 | 08NH068 | Corn Creek near Creston | 0.0 | 0.0 | 1.0 |
| 22 | 08HA034 | Craigflower Creek below Burnside Road, Vict. | 3.73 | 2.0 | 1.0 |
| 23 | 08NE087 | Deer Creek at Deer Park | 0.0 | 2.0 | 0.0 |
| 24 | 08NM176 | Ewer Creek near the Mouth | 0.403 | 2.0 | 0.0 |
| 25 | 08 NH 130 | Fry Creek below Carney Creek | 0.219 | 0.0 | 0.0 |
| 26 | 08NJ112 | Goose Creek near Crescent Valley | 0.018 | 2.0 | 0.0 |
| 27 | 08LG056 | Guichon Creek above Tunkwa Lake Diversion | 0.220 | 0.0 | 0.0 |
| 28 | 08NL050 | Hedley Creek near the Mouth | 0.584 | 0.0 | 1.0 |
| 29 | 08NE120 | Hill Creek at Mill Road Bridge | 0.224 | 4.0 | 0.0 |
| 30 | 08 NE110 | Inonoaklin Creek above Valley Creek | 0.281 | 0.0 | 0.0 |
| 31 | 08 MH 076 | Kanaka Creek near Websters Corners |  | 9.0 |  |
| 32 | 08NH005 | Kaslo River below Kemp Creek | 0.355 | 4.0 | 0.0 0.0 |
| 33 | 08 NH 132 | Keen Creek below Kyawats Creek | 0.719 | 0.0 | 0.0 |
| 34 | 08HB063 | Kitsucksus Creek above Cherry Creek | 0.101 | 3.0 | 0.0 |
| 35 | 08HA003 | Koks ilah River at Cowichan Station | 0.909 | 0.0 | 0.0 |
| 36 | 08NE006 | Kuskanax Creek near Nakusp | 0.291 | 0.0 | 0.0 |
| 37 | 08 NE117 | Kuskanax Creek at 1040 m Contour | 0.149 | 0.0 | 0.0 |
| 38 | 08MG005 | Lillooet River near Pemberton | 0.272 | 0.0 | 0.0 |
| 40 | 08HB029 | Little Qualicum River near Qualicum Beach | 2.259 | 2.0 | 0.0 |
| 40 | 08GA054 | Mamquam River above Mashiter Creek | 0.321 | 0.0 | 0.0 |

Table 7 (cont'd)

Water Survey<br>Station Number

Basin Number
$2 \times 2$
$10 \times 10$
Lakes Swamps

| 08NN021 | Moody Creek near Christina |
| :--- | :--- |
| 08NH120 | Moyie Creek above Negro Creek |
| 08MO65 | Nahatlatch River below Tachewana |
| 08MH105 | Nicomek River below Murray Creek |
| 08MH058 | Norrish Creek near Dewdney |


| 0.0 | 2.0 | 0.0 |
| :--- | ---: | ---: |
| 0.451 | 0.0 | 0.0 |
| 0.667 | 1.0 | 0.0 |
| 0.202 | 4.0 | 1.0 |
| $0.848:$ | 0.0 | 0.0 |
|  |  |  |
| 1.613 | 1.0 | 0.0 |
| 0.192 | 0.0 | 0.0 |
| 1.565 | 1.0 | 0.0 |
| 2.364 | 1.0 | 2.0 |
| 0.110 | 0.0 | 0.0 |
|  |  |  |
| 0.450 | 0.0 | 0.0 |
| 0.767 | 2.0 | 0.0 |
| 6.250 | 5.0 | 0.0 |
| 0.224 | 3390 | 0.0 |
| 0.182 | 0.0 | 0.0 |
|  |  |  |
| 0.017 | 2.0 | 0.0 |
| 0.520 | -1.0 | 0.0 |
| 1.584 | 1.0 | 0.0 |
| 0.297 | 0.0 | 0.0 |
| 0.521 | 0.0 | 1.0 |
|  |  |  |
| 0.857 | 2.0 | 0.0 |
| 3.432 | 2.0 | 0.0 |
| 0.224 | 0.0 | 0.0 |
| 0.583 | 0.0 | 0.0 |
| 0.441 | 0.0 | 0.0 |
| 0.987 | 1.0 | 0.0 |
| 0.151 | 1.0 | 4.0 |
| 0.403 | 0.0 | 0.0 |
| 0.040 | 2.0 | 0.0 |
| 1.251 | 1.0 | 0.0 |
| 0.560 | 1.0 | 0.0 |
|  |  |  |

Table 8 . Mode of Basin Aspect in Degrees Clockwise from North

| Water Survey |  | From | From |
| :--- | :---: | :---: | :---: |
| Station Number | Basin Name | $2 \times 2$ | $10 \times 10$ |


| 1 | 08NH084 | Arrow Creek near Erickson | 180 | 208 |
| :---: | :---: | :---: | :---: | :---: |
| 2 | 08NL004 | Ashnola River near Keremeos | 1815 | 208 |
| 3 | 08 NE077 | Barnes Greek near Needles | 315 | 351 |
| 4 | 08NE008 | Beaton Creek near Beaton | 135 45 | 156 |
| 5 | 08NM035 | Bellevue Creek near Okanagan Mission | 45 315 | 28 319 |
| 6 | 08NE039 | Big Sheep Creek near Rossland | 135 |  |
| 7 | 08HA016 | Bings Creek near the Mouth | 135 135 | 164 |
| 8 | 08LE094 | Bolean Creek near the Mouth | 135 | 164 |
| 9 | 08 NH 032 | Boundary Creek near Porthill | 225 45 | 216 |
| 10 | 08NM133 | Bull Creek near Crump | 45 45 | 147 83 |
| 11 | 08NN023 | Burrell Creek above Gloucester Creek | 135 | 118 |
| 12 | 08NH131 | Carney Creek below Pambrun Creek | 135 45 | 118 |
| 13 | 08GA060 | Chapman Creek above Sechelt Diversion | 180 | 180 |
| 14 | 08 HA 001 | Chemainus River near Westholme | 180 90 | 119 |
| 15 | 08 MH 103 | Chilliwack River above Slesse Creek | 315 | 337 |
| 16 | 08MH016 | Chilliwack River at Outlet of Chilliwack Lake | 315 | 337 |
| 17 | 08NM142 | Coldstream Creek above Municipal Intake | 180 | 190 |
| 18 | 08LG048 | Coldwater River near Brooknere | 45 | 18 |
| 19 | 08MF003 | Coquihalla River near Hope | 315 | 279 |
| 20 | 08MF062 | Coquihalla River below Needle Creek | 315 | 199 |
| 21 | 08NH068 | Corn Creek near Creston | 135 | 321 |
| 22 | 08 HA 034 | Craigflower Creek below Burnside Road, Victoria | 180 | 124 |
| 23 | 08NE087 | Deer Creek at Deer Park | 225 | 252 |
| 24 | 08NM176 | Ewer Creek near the Mouth | 90 | 164 |
| 25 | 08NH130 | Fry Creek below Carney Creek | 315 | 191 |
| 26 | 08NJ112 | Goose Creek near Crescent Valley | 135 | 141 |
| 27 | 08LG056 | Guichon Creek above Tunkwa Lake Diversion | 45 | 81 |
| 28 | 08NL050 | * Hedley Creek near the Mouth | 225 | 174 |
| 29 | 08NE 120 | Hill Creek at Mill Road Bridge | 315 | 208 |
| 30 | 08NE 110 | * Inonoaklin Creek above Valley Creek | 45 | 155 |
| 31 | 08MH076 | Kanaka Creek near Websters Corners | 225 | 172 |
| 32 | 08NH005 | Kaslo River below Kemp Creek | 135 | 172 |
| 33 | 08NH132 | Keen Creek below Kyawats Creek | 270 | 185 |
| 34 | 08 HB 063 | Kitsucksus Creek above Cherry Creek | 180 | 320 |
| 35 | 08 HA 003 | Koksilah River at Cowichan Station | 180 45 | 320 32 |
| 36 | 08NE006 | Kuskanax Creek near Nakusp | 225 | 190 |
| 37 | 08NE117 | Kuskanax Creek at 1040 m Contour | 225 | 186 |
| 38 | 08MG005 | * Lillooet River near Pemberton | 45 | 141 |
| 39 | 08 HB 029 | Little Qualicum River near Qualicum Beach | 45 | 141 |
| 40 | 08GA054 | Mamquam River above Mashiter Creek | 225 | 223 |

[^0]Table 8 (cont'd)
Water Survey
Station Number
Basin Name
From
From
$2 \times 2$
$10 \times 10$

| 41 | 08NN021 | Moody Creek near Christina | 135 | 133 |
| :---: | :---: | :---: | :---: | :---: |
| 42 | 08NH120 | * Moyie River above Negro Creek | 135 | 65 |
| 43 | 08MF065 | * Nahatlatch River below Tachewana | 45 | 176 |
| 44 | 08MH105 | Nicomekl River below Murray Creek | 315 | 260 |
| 45 | 08MH058 | Norrish Creek near Dewdney | 135 | 153 |
| 46 | 08NL060 | Otter Creek below Spearing Creek | 45 | 189 |
| 47. | 08NL069 | Pasayten River above Calcite Creek | 315 | 203 |
| 48 | 08NM172 | Pearson Creek near the Mouth | 225 | 250 |
| 49 | 08LG016 | Pennask Creek near Quilchena | 0 | 352 |
| 50 | 08NE074 | Salmo River near Salmo | 225 | 284 |
| 51 | $08 \mathrm{HA010}$ | San Juan River near Port Renfrew | 225 | 208 |
| 52 | 08 HB 014 | Sarita River near Bamfield | 315 | 242 |
| 53 | 08HA033 | Shawnigan Creek near Mill Bay | 0 | 142 |
| 54 | 08NL070 | Similkameen River above Goodfellow Creek | 135 | 203 |
| 55 | 08MH056 | Slesse Creek at Vedder Crossing | 45 | 12 |
| 56 | 08NN016 | Sutherland Creek near Fife | 225 | 185 |
| 57 | 08NN019 | Trapping Creek near the Mouth | 180 | 195 |
| 58 | $08 \mathrm{HB024}$ | Tsable River near Fanny Bay | 45 | 330 |
| 59 | 08NL067 | * Tulameen River near Tulameen | 315 | 151 |
| 60 | 08NL071 | Tulameen River below Vuich Creek | 315 | 207 |
| 61 | 08MG021 | Twenty-One Mile Creek at 670 m Contour | 135 | 19 |
| 62 | 08HC002 | Ucona River at the Mouth | 130 | 219 |
| 63 | 08LC040 | Vance Creek below Deafies Creek | 135 | 94 |
| 64 | 08NM015 | Vaseux Creek above Dutton | 225 | 280 |
| 65 | 08NM171 | Vaseux Creek above Solco Creek | 315 | 201 |
| 66 | 08NN015 | West Kettle River near McCulloch |  |  |
| 67 | 08NL036 | Whipsaw Creek below Lamont Creek | 135 | 285 |
| 68 | 08NM174 | * Whiteman Creek above Bouleau Creek | 13 | 118 |
| 69 | 08NJ120 | Winlaw Creek near Winlaw | 270 | 184 |
| 70 | 08LG009 | Witches Brook near Merritt | 45 | 189 95 |
| 71 | 08NL041 | Wolfe Creek at Outlet of Issitz Lake | 315 | 229 |

[^1]Table 9 Regional Regression Equations for Wet Interior Region 26 Basins Mean Basin Runoff: 846 mm


## Table 11 Regional Regression Equations for Coastal Region: Vancouver Island and Lower Mainland 19 Basins Mean Basin Runoff: 1912 mm



Table 12 Frequency of 0ccurrence of Variables in Equation Building for Wet Interior Region

| Variable |  | Total |
| :--- | :--- | :---: |
| ELEV.2 | $X$ | 1 |
| RALS.2 | $X X$ | 2 |
| DRNGD2 |  |  |
| SLPCC2 | XXXXXX | 6 |
| SLPEL2 |  |  |
| ASPCT2 | $X X X X$ | 4 |
| DIS.NW |  |  |
| DIS.N | $X X X X$ | 4 |
| DIS.SW |  |  |
| RALK10 | $X X X X X$ |  |
| RAFOR |  | 2 |
| RASWMP | $X X$ | 1 |
| RAGLAC | $X$ | 1 |
| BH.N |  |  |
| BH.NW |  |  |
| BH.W |  |  |
| BH.SW |  |  |
| SE.N |  |  |
| SE.WW |  |  |
| SE.SW |  |  |

APPENDIX 1.
DESCRIPTION OF 10 X 10 PHYSIOGRAPHIC DATA

## ABBREVIATIONS

| AREAl0 | The 10 km by 10 km basin area |
| :--- | :--- |
| ELEVIO | The 10 km by 10 km mean basin elevation |
| SLP.10 | The 10 km by 10 km mean basin slope |
| ASPTIO | The lo km by 10 km mean basin aspect |
| DIS.N | Distance to sea to North |
| DIS.NW | Distance to sea to North-West |
| DIS.SW | Distance to sea to South-West |
| RALK10 | The 10 km by 10 km relative area of lakes |
| RAFOR | Relative area of forest |
| RASWP | Relative area of swamps |
| RAGLC | Relative area of glaciers |
| RAURB | Relative area of urbanization |
| BH.N | Barrier height to North |
| BH.NW | Barrier height to North-West |
| BH.W | Barrier height to West |
| BH.SW | Barrier height to South-West |
| SE.N | Shield effect to North |
| SE.NW | Shield effect to North-West |
| SE.W | Shield effect to West |
| SE.SW | Shield effect to South-West |

The 10 km by 10 km grid was overlaid on $1: 250,000$ scale topographic maps. Elevation sampling consisted of extracting elevations at the four corner points and at the centre of each $10 \times 10$ square. The calculation of slope for each square is described in Section 3.2.4. The basin slope is the average value for the squares included in the basin.

The aspect of a basin is described in Section 2.6.
Distance to Sea

Distance from centre of gravity of basin to the sea in the

| North | DSN |  |
| :--- | :--- | :--- |
| Northwest | DSNW |  |
| West | DSW | Kms. |
| Southwest | DSSW |  | north, the northwest, west, and southwest directions

Southwest DSSW

The relative areas of lakes, forest, glacier and urban land cover were taken from the maps by counting the number of dots on a random array which covered the appropriate classification.

| Lake | RALKE |  |
| :---: | :---: | :---: |
| Forest | RAFOR |  |
| Swamp | RASWP |  |
| Glacier | RAGLC | Dimension- |
| Urban | RAURB | less |
| Barrier Height |  |  |
| North | BHN |  |
| Northwest | BHNW |  |
| West | BHW | Feet |
| Southwest | BHSW |  |

For each $10 \times 10$ square the value of barrier height is weighted by barrier heights of 5 squares to the north and 5 squares to the south.

Shield Effect

| North | SEN |
| :--- | :--- |
|  |  |
| Northwest | SENW |
| West | SEW Feet |
| Southwest | SESW |

Sum of elevation differential of all ascending stretches of terrain encountered when travelling from ocean shore at north, northwest, west, southwest directions to corresponding point.

APPENDIX 2. DESCRIPTION OF $2 \times 2$ PHYSIOGRAPHIC DATA

The $2 \times 2$ data were extracted from 1:50,000 scale Universal Transverse Mercator Projection maps. A one kilometre square grid with northing and easting coordinates is printed on the maps.

AREA. 2
ELEV. 2
RALS. 2

DRNGD2
SLPCC2 The 2 km by 2 km mean basin slope by contour crossings
SLPEL2 The 2 km by 2 km mean basin slope by elevation differences
ASPT. 2 The 2 km by 2 km mode of the aspect of the squares

Mean Elevation

This is the elevation at the centre of the 2 X 2 square which is assumed to represent the average elevation of the square.

Average Land Slope

The contour lines on the maps provide a basis for determining the slope of the land by the intersection-line method as outlined by Horton and described in the U.S. Geological Survey [1947] Water Supply Paper. The number of contour lines crossing the north-south and east-west centre lines of the square are counted and recorded. By the use of the following formula the mean slope of the area can be determined: slope $=N h \sec (c c), \sec (c c)=1.571$,
the constant for averaging the angle at which contours cross
the intersection line, $h$ is the contour interval or difference in elevation in feet, $N$ is the number of contours crossed, and $l$ is the sum of lengths of the centre lines.

## Stream Density

This is defined as the total length of the streams divided by the area. Thus the total length of streams in the square is the desired parameter to be identified. This is accomplished by setting a pair of dividers a specified distance (in this study scaled at 0.2 miles before metric conversion) and recording the number of intervals required to trace the length of the streams in the square.

Area of Lakes and Swamps

This area is determined with the aid of a random dot overlay, then counting and recording the number of dots found within the lake or swamp boundaries in each square. By a suitable conversion, the area of lakes and swamps may then be computed.

To estimate the area of lakes and swamps from 1:50,000 maps:

```
number of dots = area in square inches
    100
```

Maps are 1:50,000 or 1 inch on map $=50,000$ inches on the ground. There are $5,280 \mathrm{X} 12=63,360$ inches in one mile. One inch on mapsheet $=50,000=0.7891$ miles on the ground
or 1 square inch on mapsheet equals $(0.7891)=0.62$ miles on the ground.


[^0]:    * large differences in results

[^1]:    * large differences in results

